



Climatology of the Arctic Ocean based on NEMO results

SU Jie (sujie@ouc.edu.cn), LI Xiang, ZHANG Yang
Key Lab of Polar Oceanography and Global Ocean Change
Ocean University of China, Qingdao, China

Cooperator: Youyu Lu, Zeliang Wang, Frederic Dupont
Bedford Institute of Oceanography, Dartmouth, Canada



Outline

- Motivation
- Model description
- Basic results of 10-y's standard run (global and Arctic)
- Scientific interest
 - Pacific inflow's effect on sea ice distribution and upper layer ocean
 - How to get a reasonable middle layer water from Atlantic with a coarse resolution model?



Motivation

- Realistic modelling Arctic conditions (climatology and variability) and its influences on global ocean climate
- Understanding the Pacific inflow and Atlantic inflow's influences on Arctic Ocean



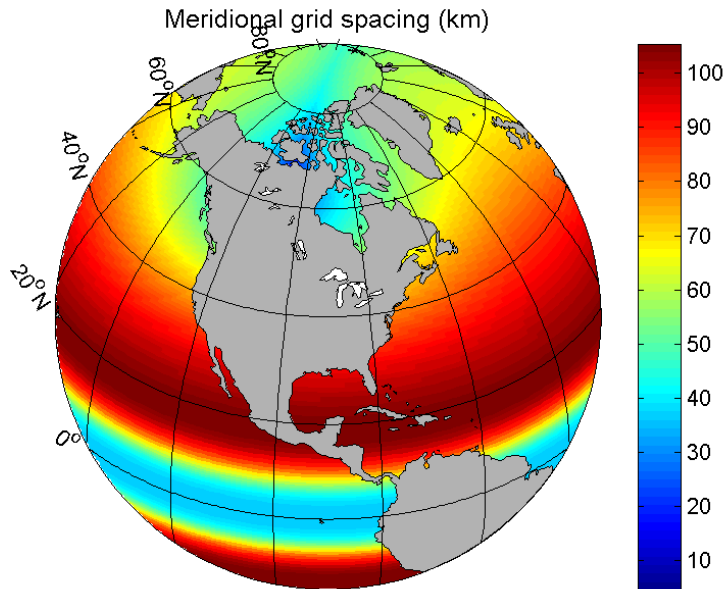
Model Description

- A coupled ice-ocean model under the NEMO framework
 - the Ocean model: OPA (Madec, et al., 1998) ;
 - The ice model: LIM2 (Fichefet and Morales, 1997; Christian and Gurvan)
- Initial field:
 - January climatology of temperature and salinity based on **PHC3.0**
 - Sea ice thickness: 3m in northern hemisphere; 1m in southern hemisphere
- Surface forcing:
 - OMIP daily climatology (from 1957 -2002 ERA-40) of wind stress, air temperature, cloud cover, relative humidity, precipitation, wind speed;
 - SSS restored
- Boundary forcing
 - Arctic model takes boundary forcing from global model

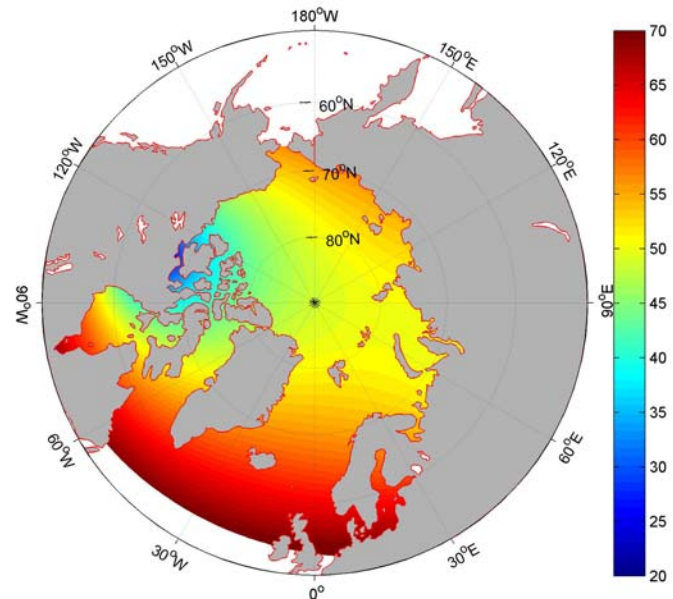
Model's Grid

- Tri-polar grids
 - The horizontal resolution: 1 degrees Mercator mesh with finest resolution in CAA is 25km
 - The vertical resolution is 46 levels (interval: 6-250m)
 - The time step is 1800s.
- Standard run include both global and arctic domain
- Model is integrated for 10 years

ORCA1 glob grids



ORCA1 actc grids

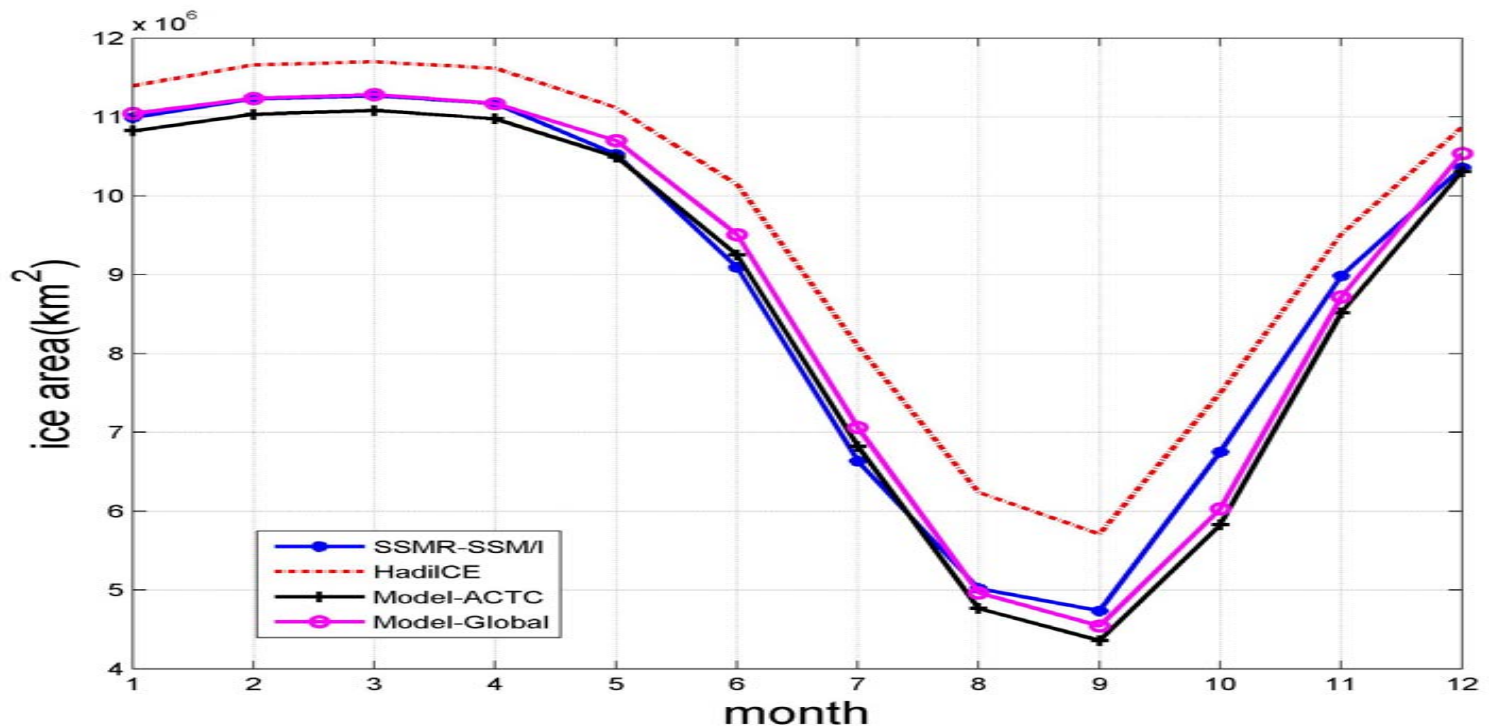




Basic results of 10-y's global and Arctic standard run

- Ice area
- Ice concentration
- Ice motion
- Ocean current (upper 100m and 450m)
- Sea temperature field (10m and 450m)
- Vertical temperature profile

Total ice area of Arctic



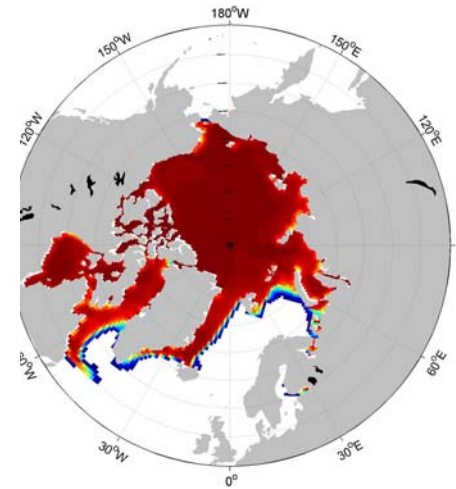
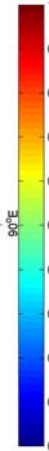
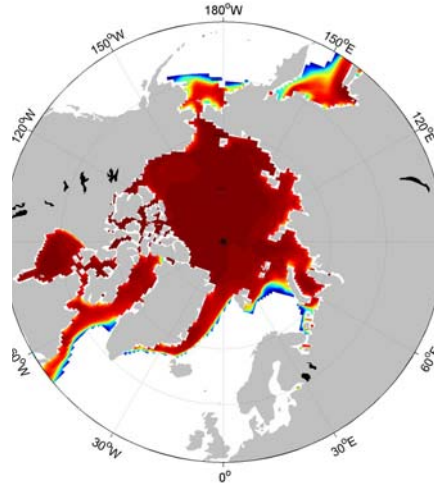
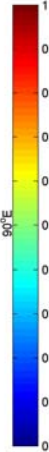
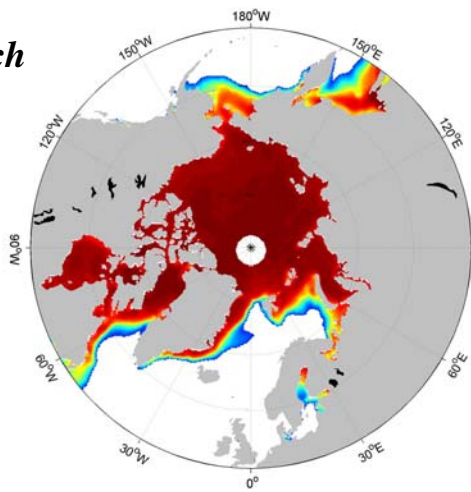
Sea ice concentration

SSM/I 1979-2002

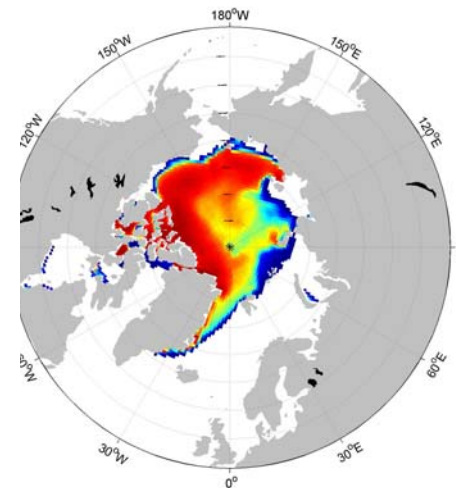
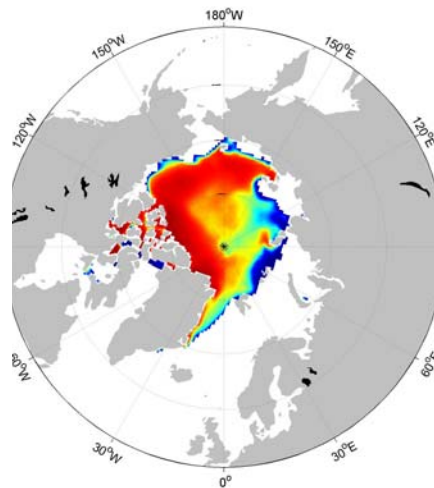
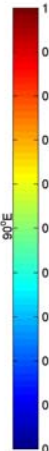
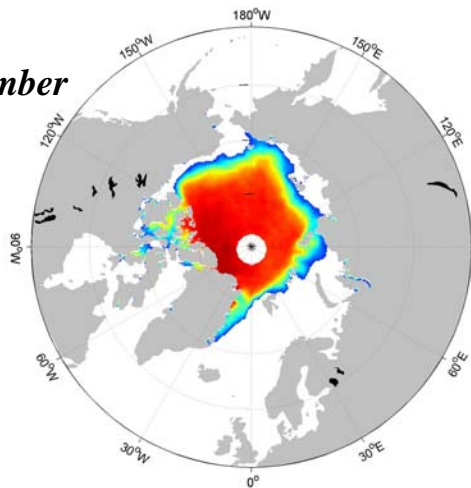
Modeled-glob

Modeled-actc

March



September



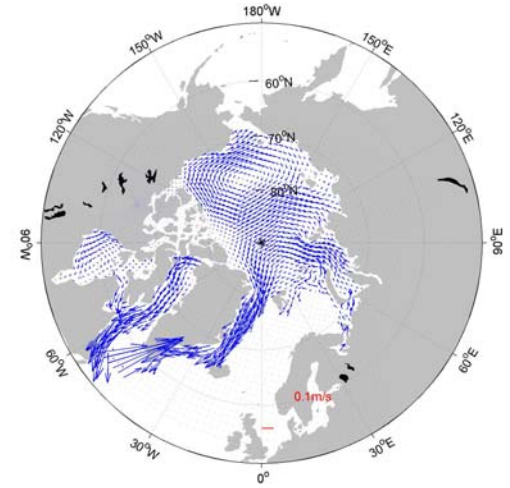
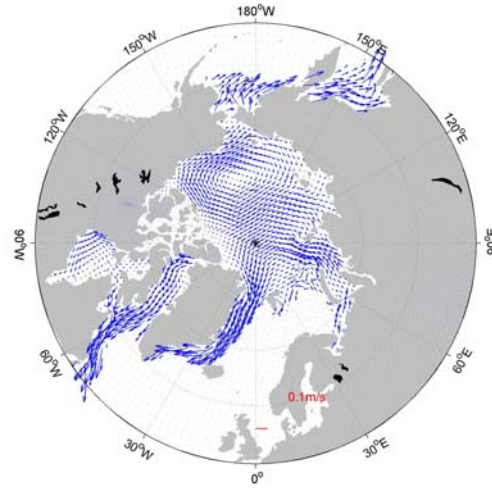
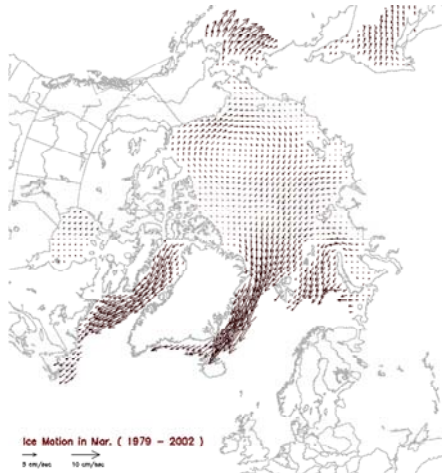
Sea ice motion

NSIDC 1979-2002

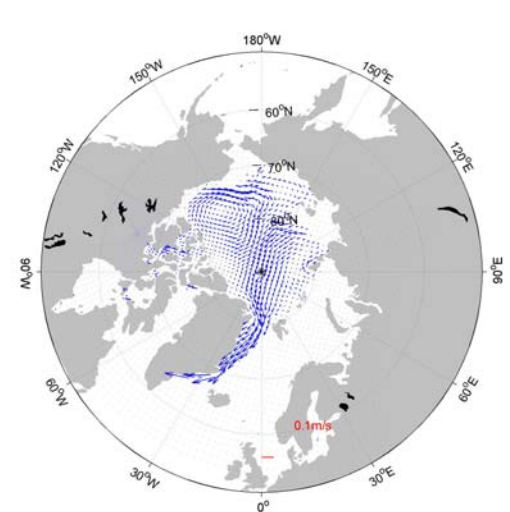
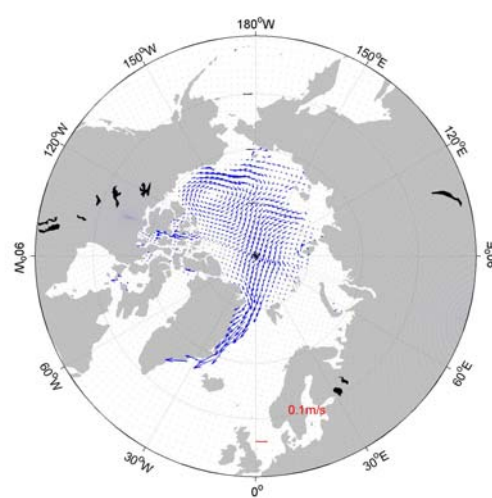
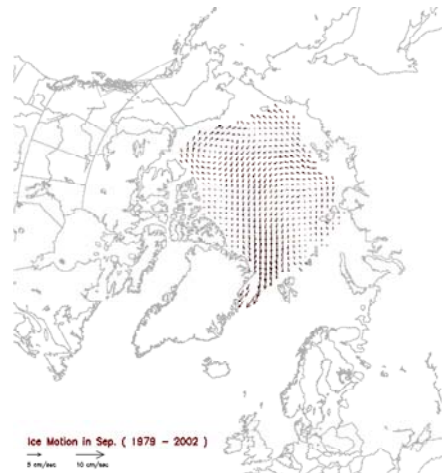
Modeled-glob

Modeled-actc

March



September

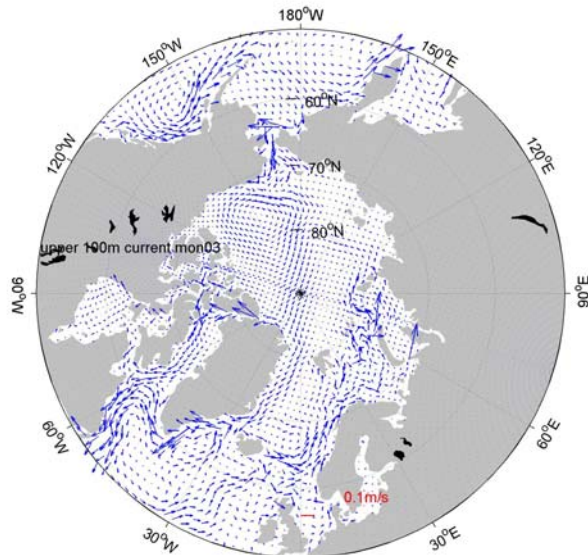


Ocean current

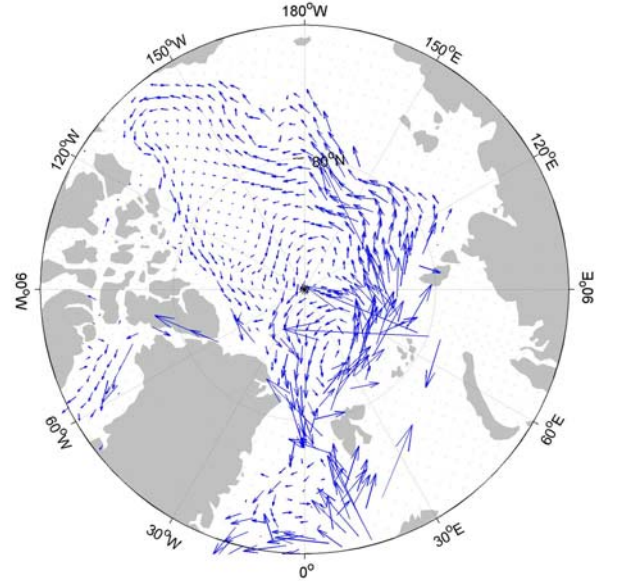
upper 100m

450m

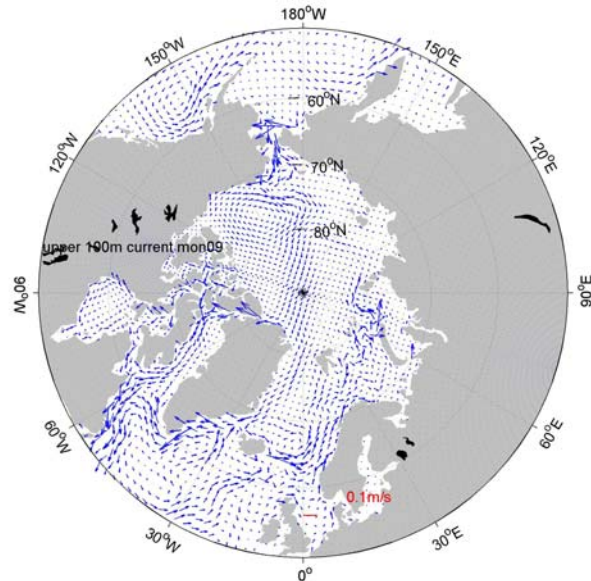
March



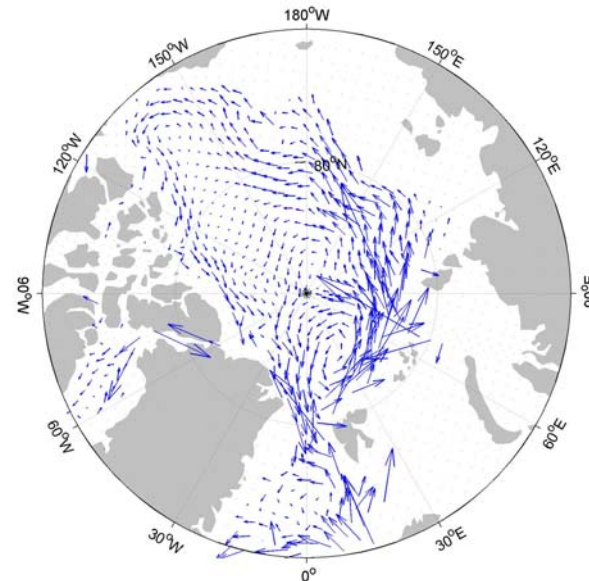
Im03data15



September



Im09data15



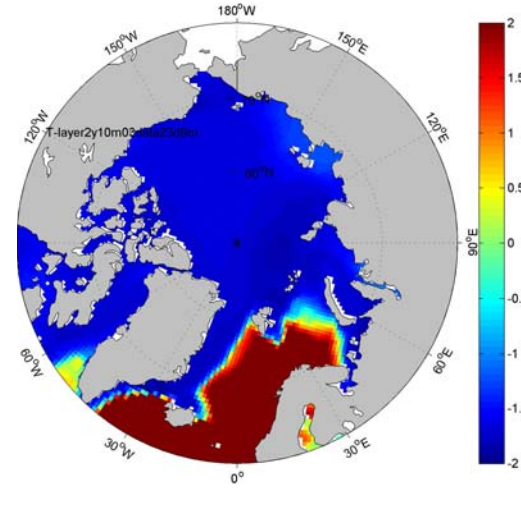
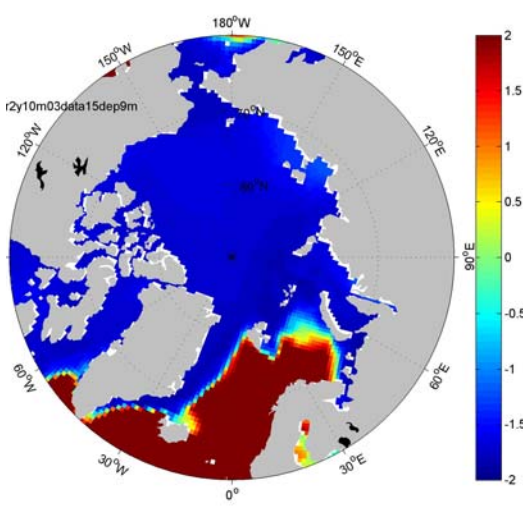
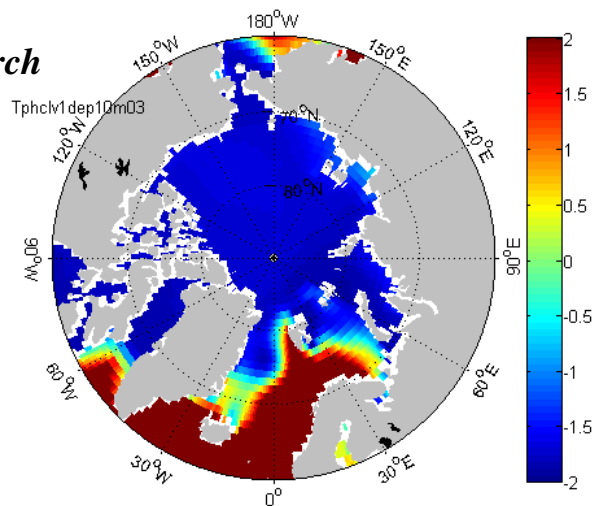
10m sea temperature

PHC3.0

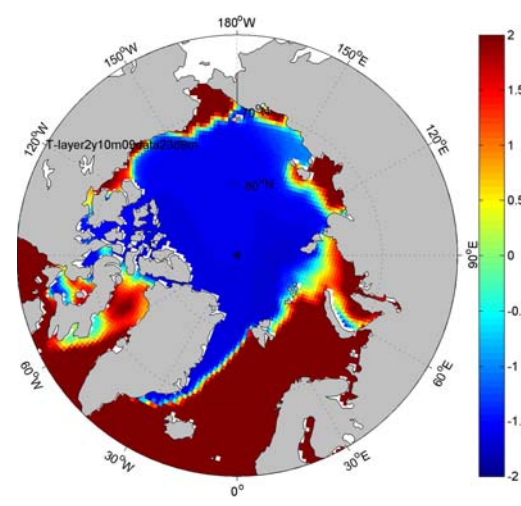
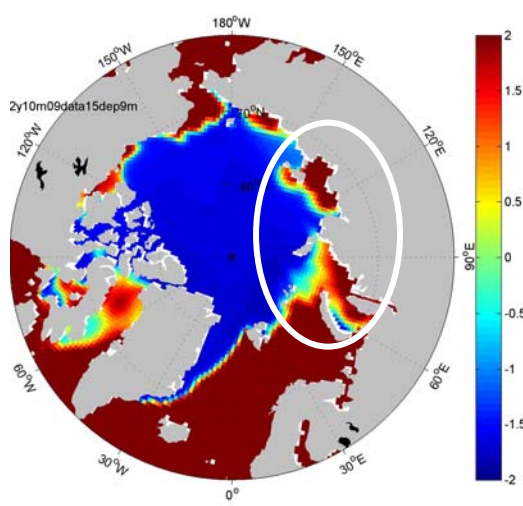
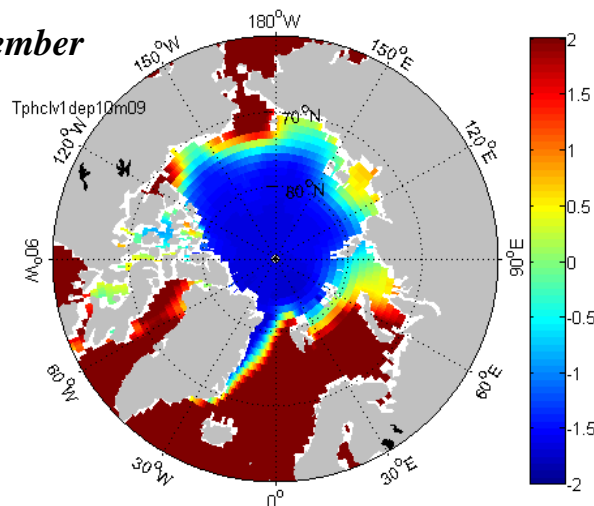
Modeled-glob

Modeled-actc

March



September



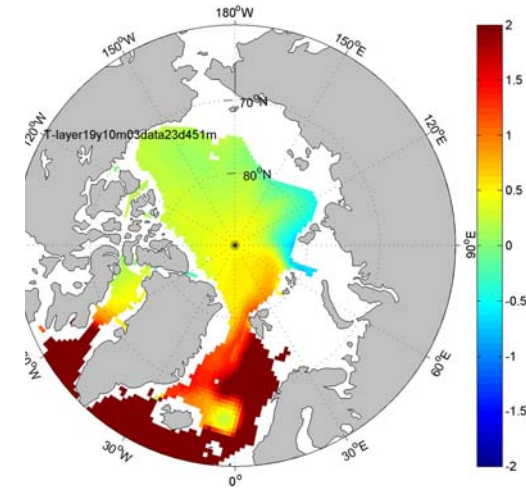
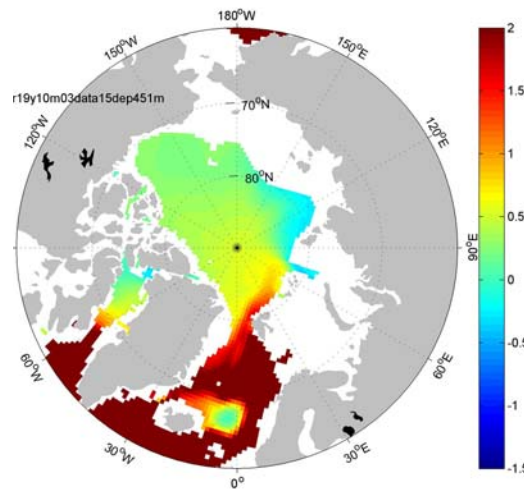
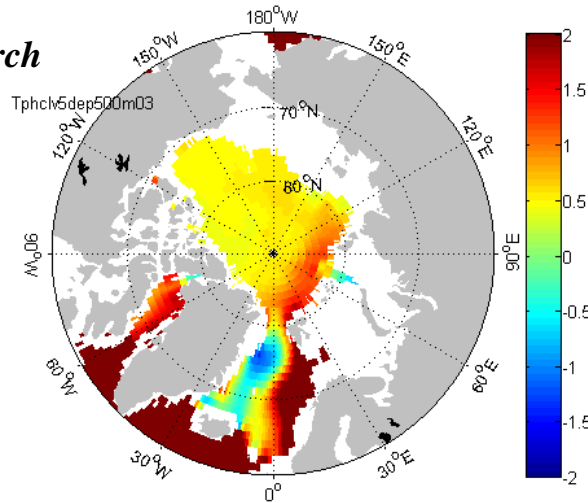
450m sea temperature

PHC3.0

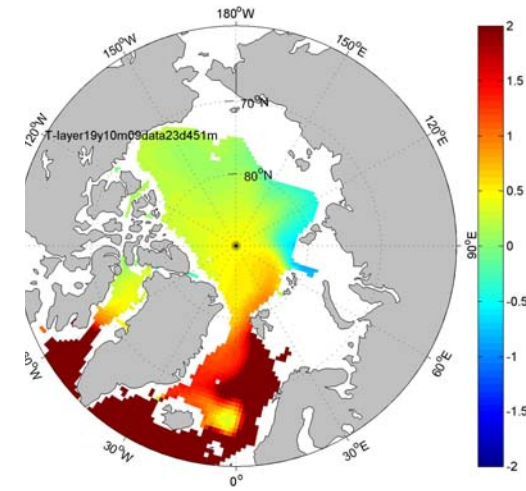
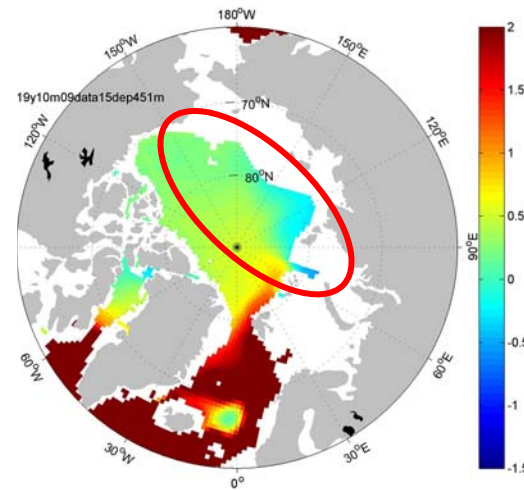
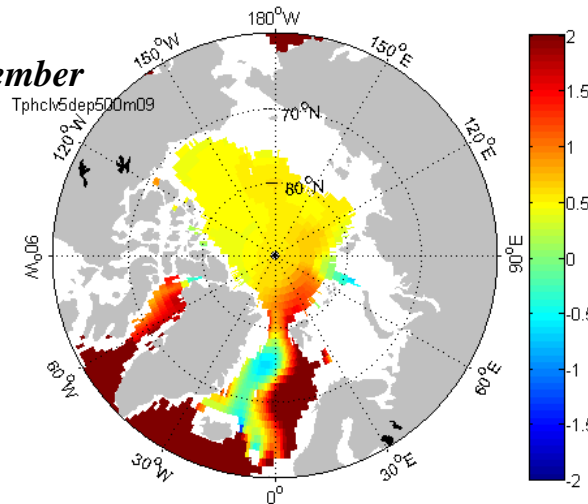
Modeled-glob

Modeled-actc

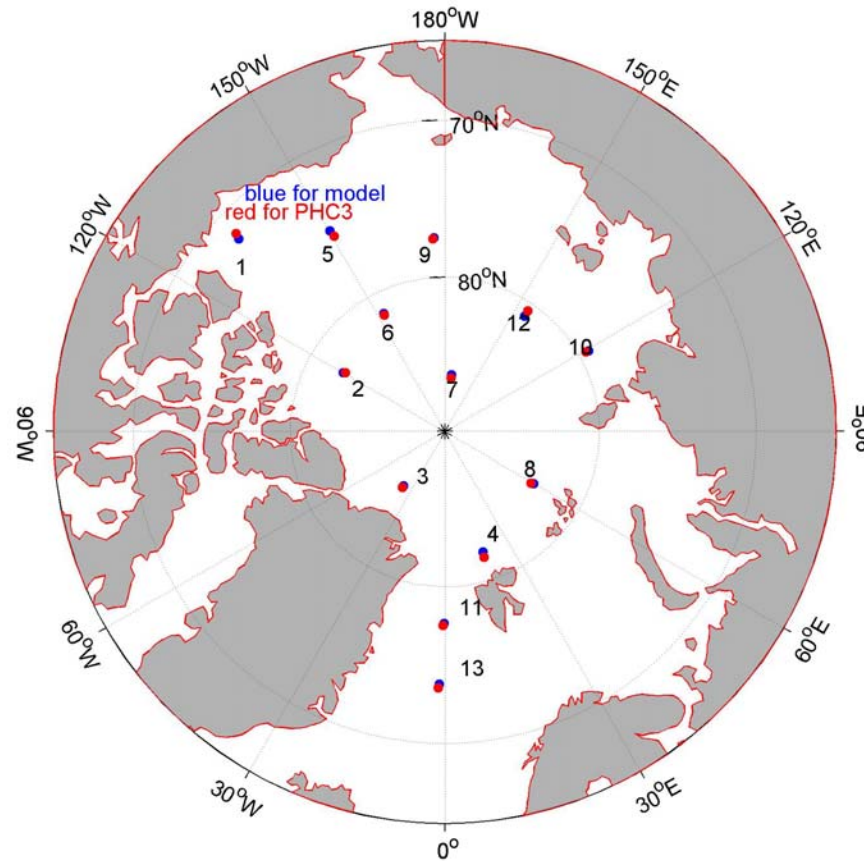
March



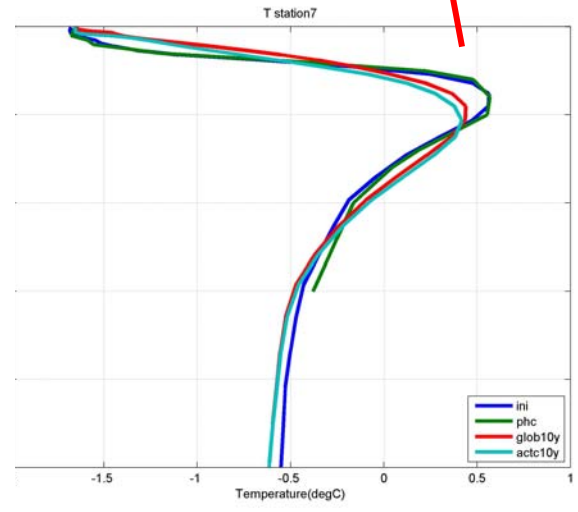
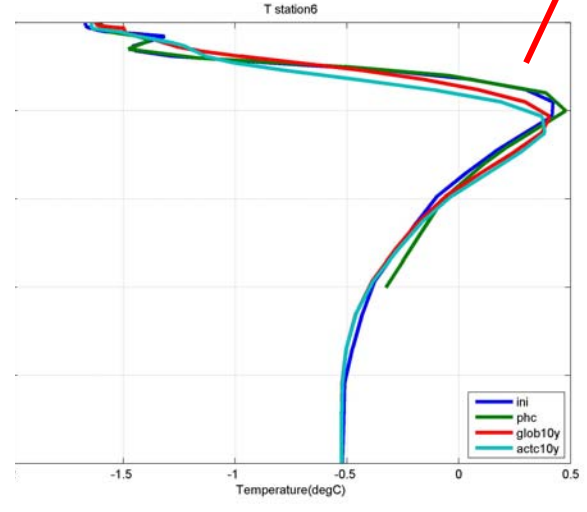
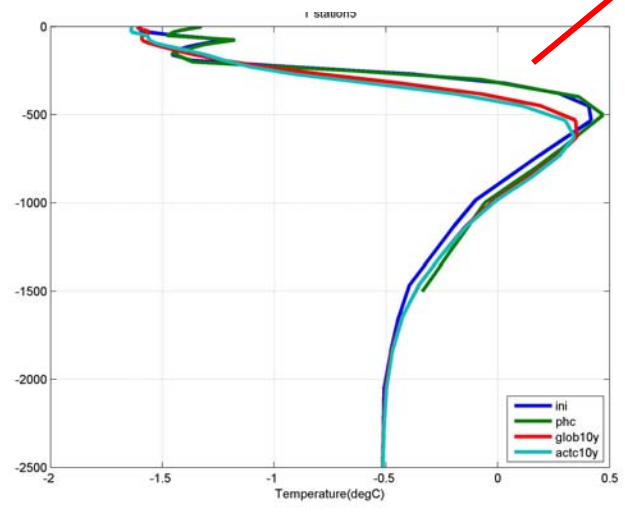
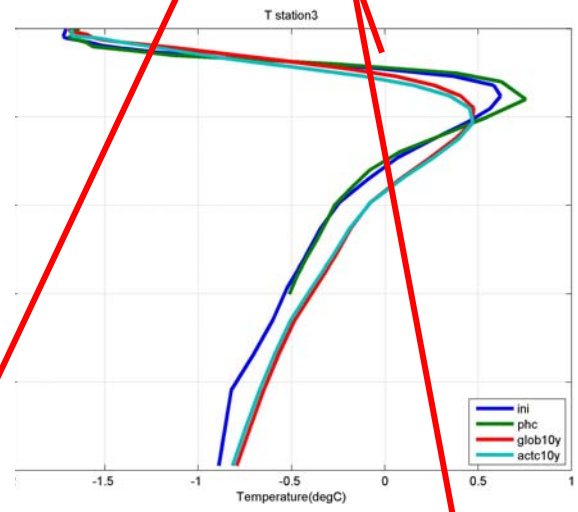
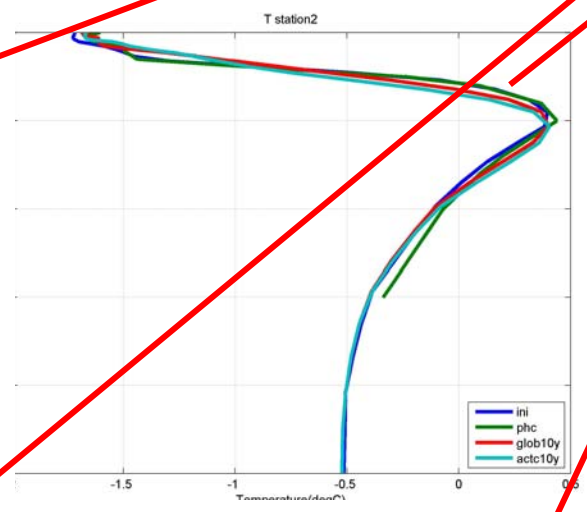
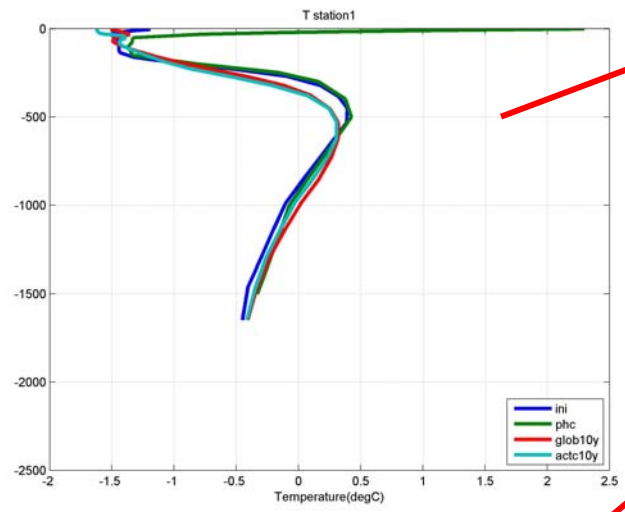
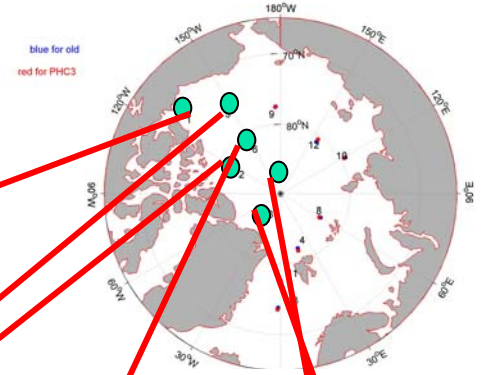
September



Vertical temperature profile

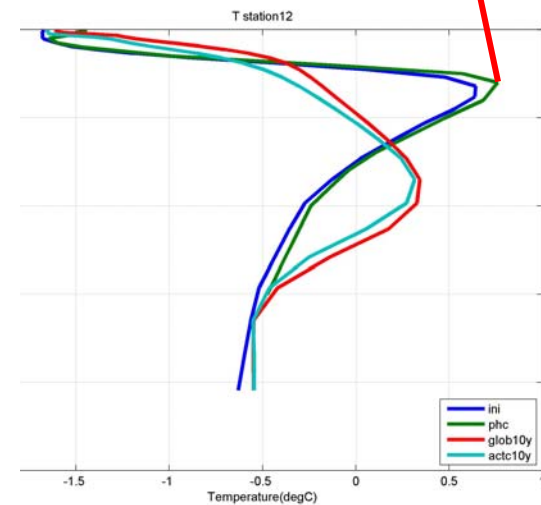
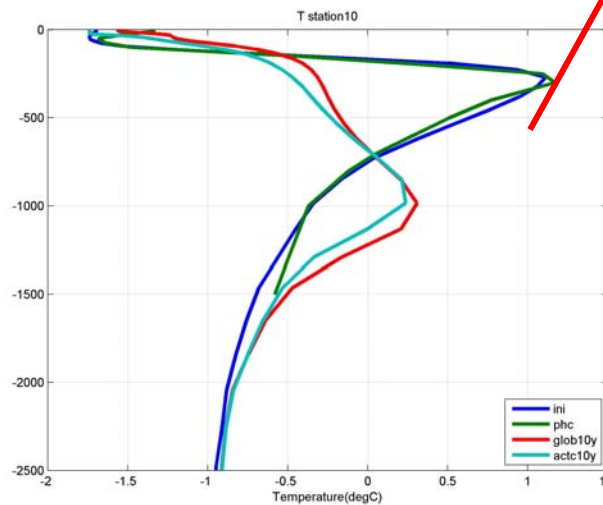
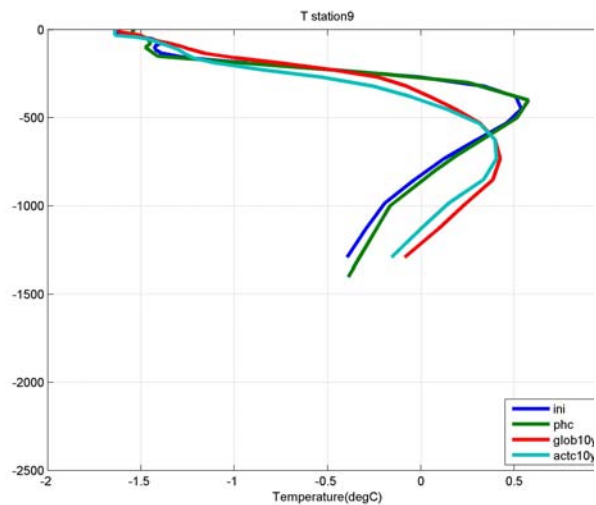
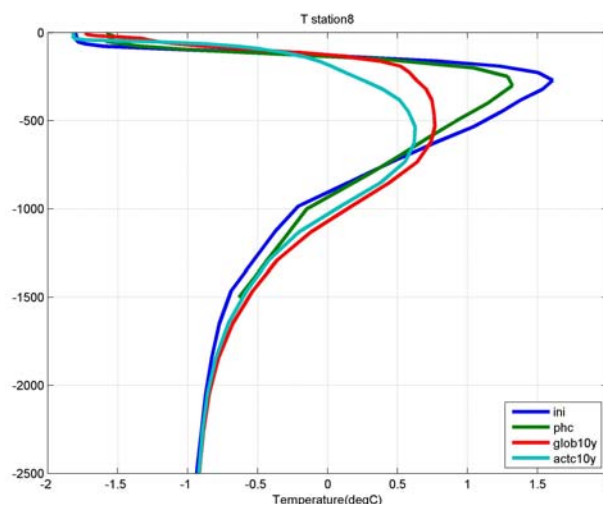
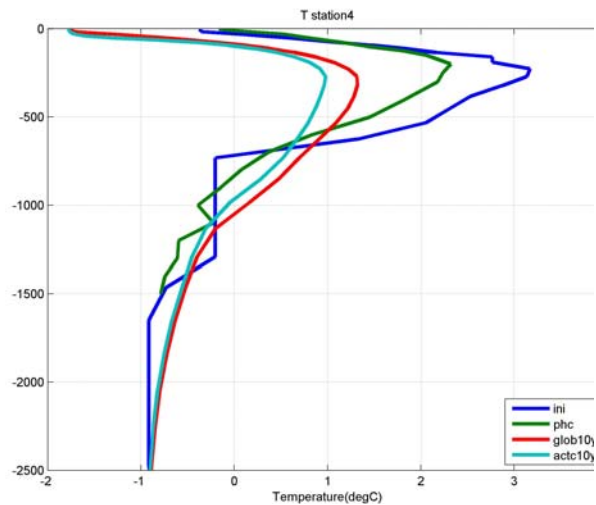
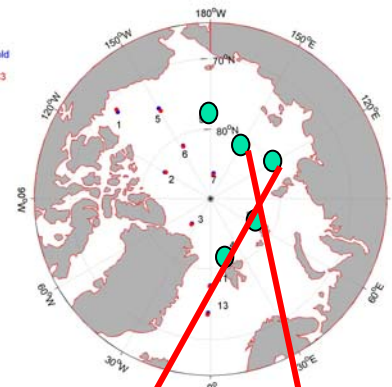


Station 1,2,3,5,6,7

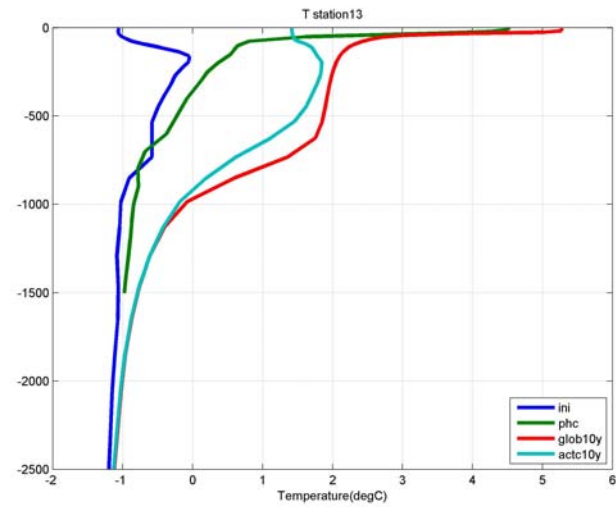
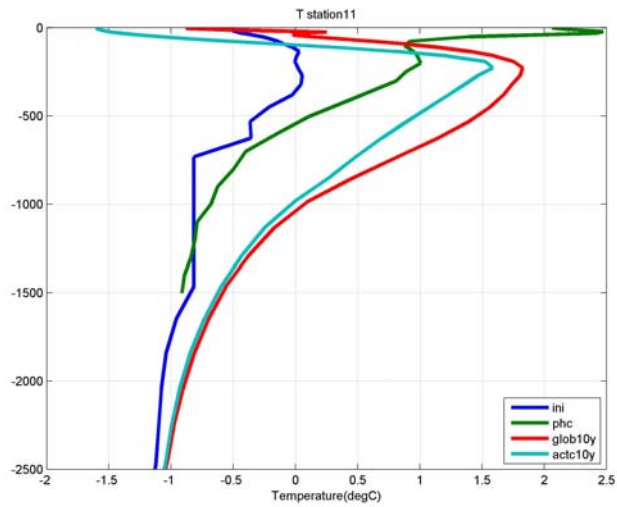
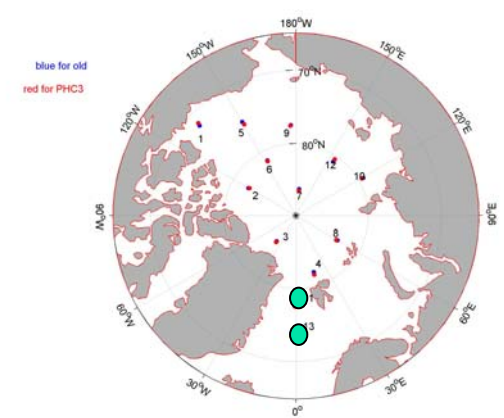


Station 4,8,9,10,12

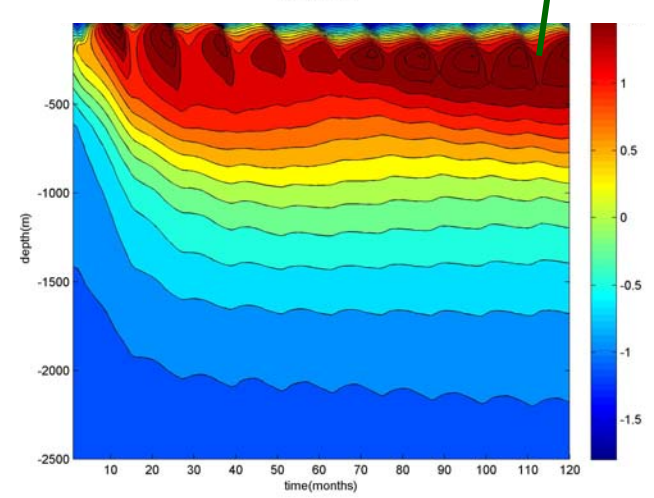
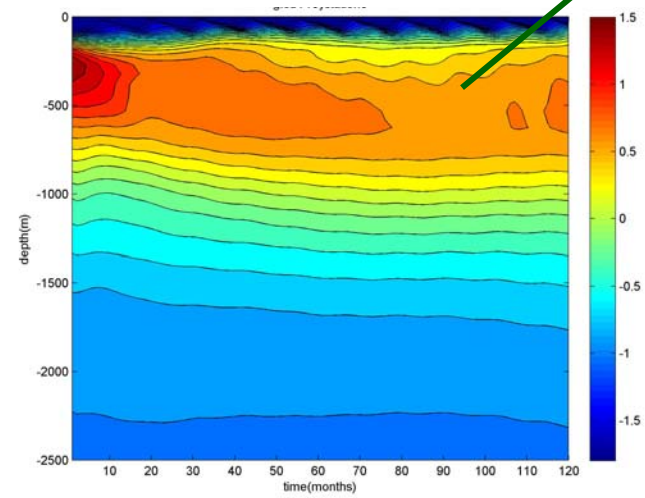
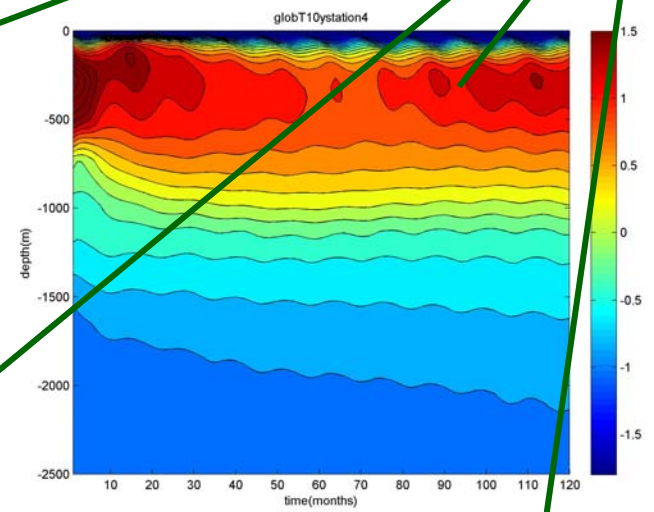
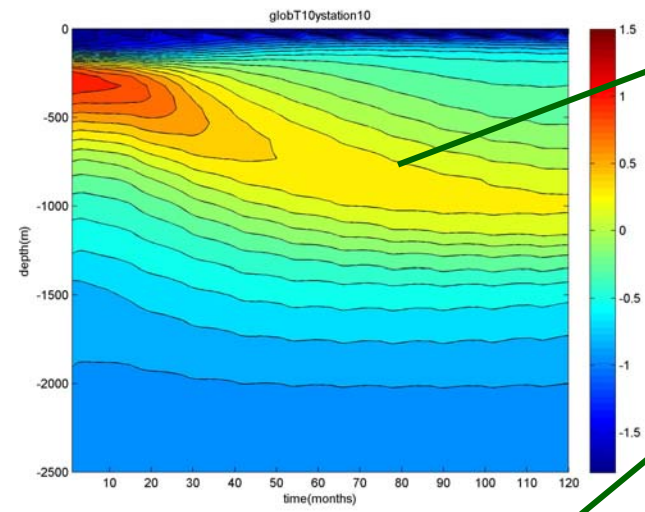
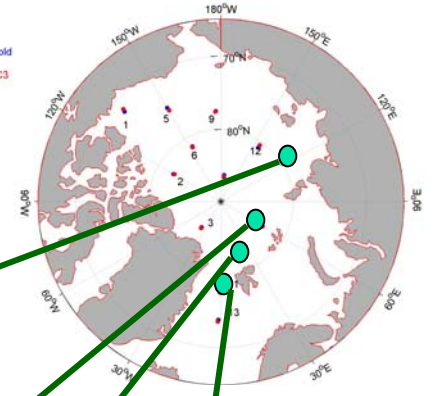
blue for old
red for PHC3



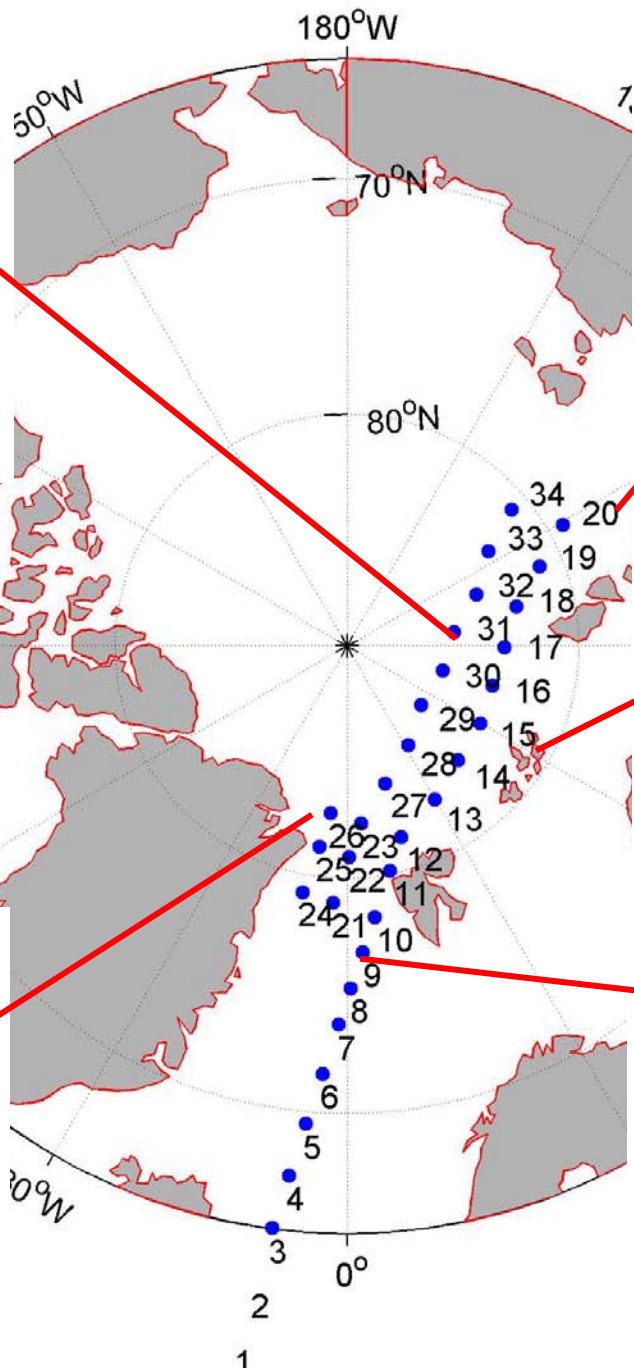
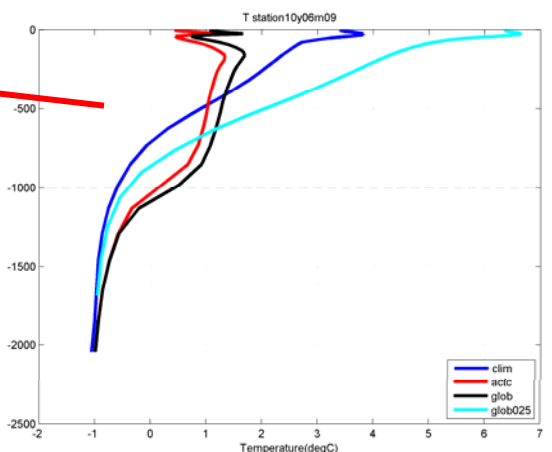
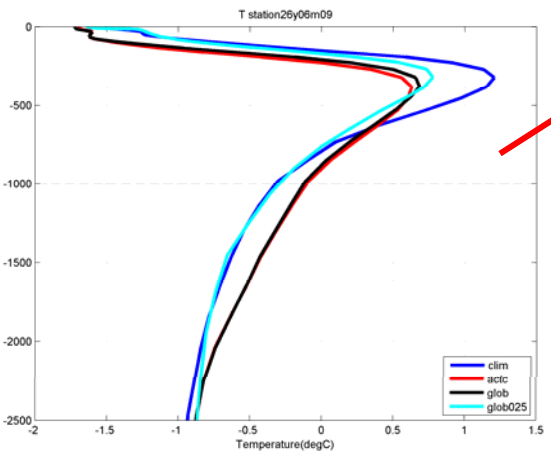
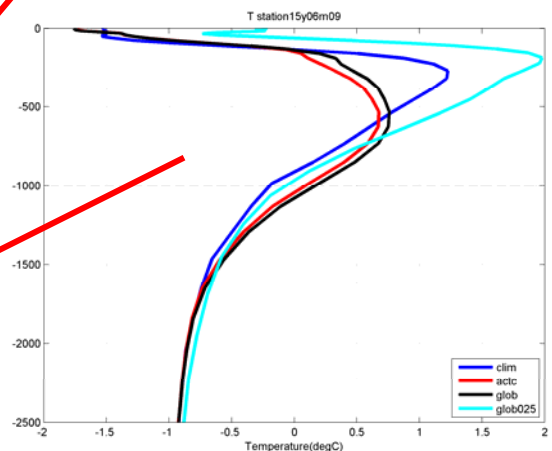
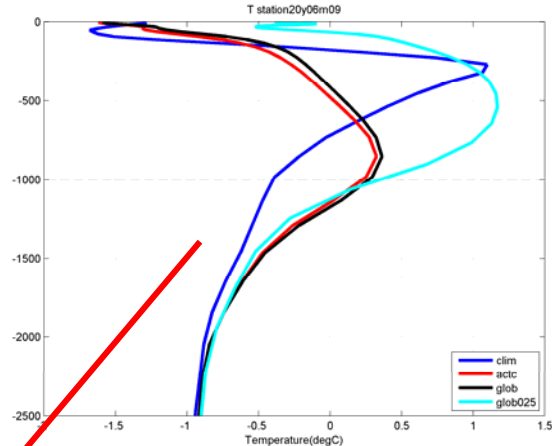
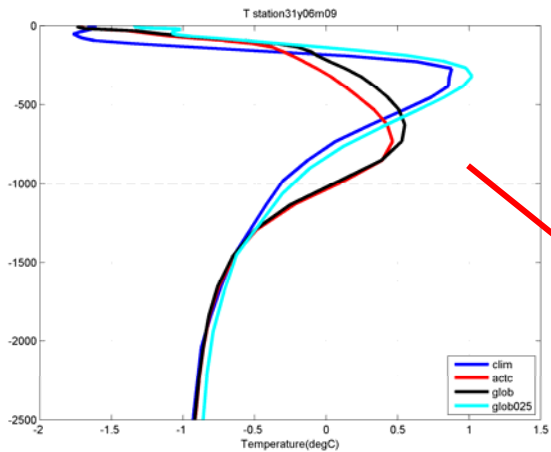
Station 11,13



How to get a reasonable middle layer water from Atlantic in a coarse resolution model?



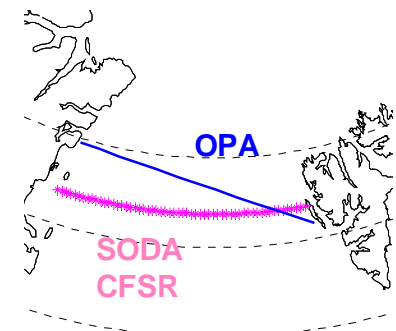
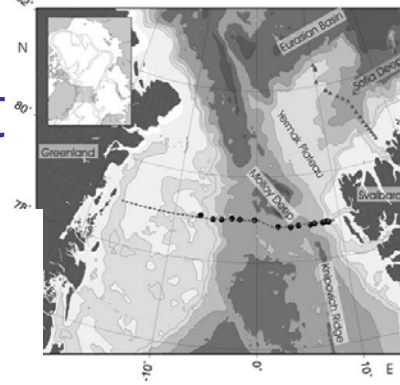
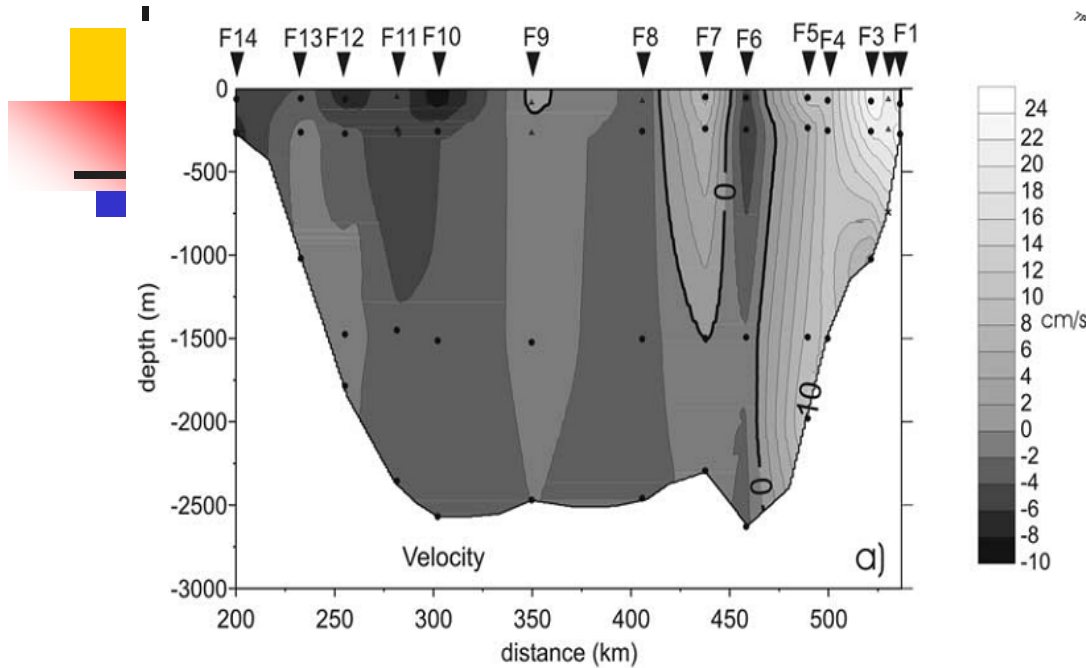
||



¼ degree resolution

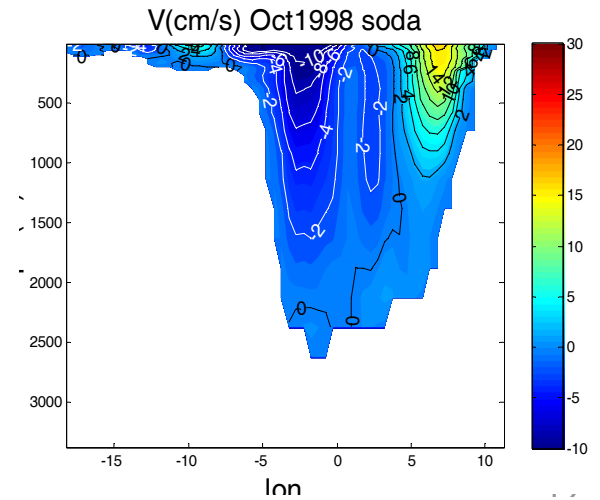
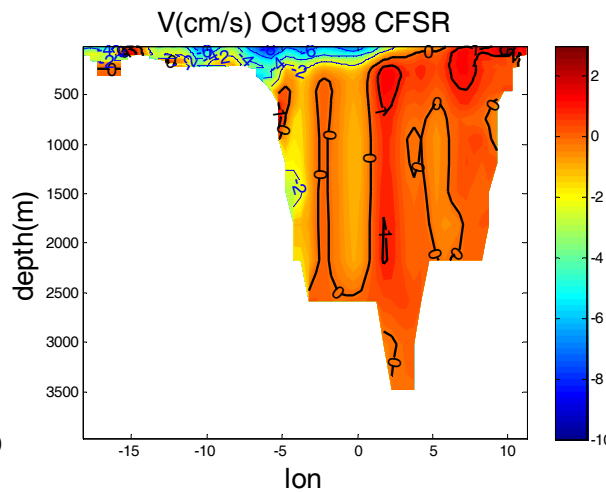
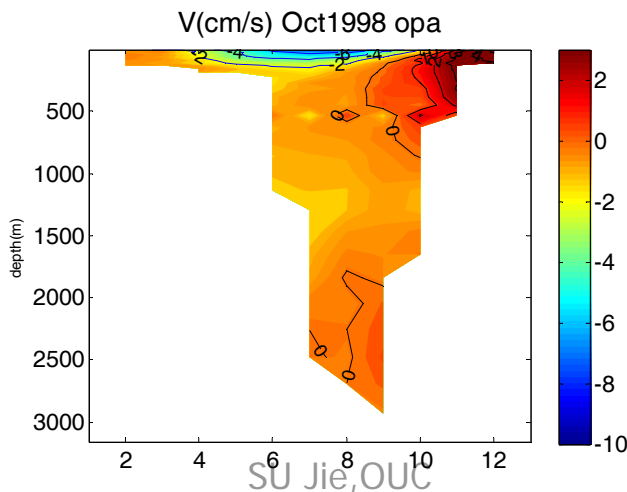
M₀06

Velocity section of Fram Strait

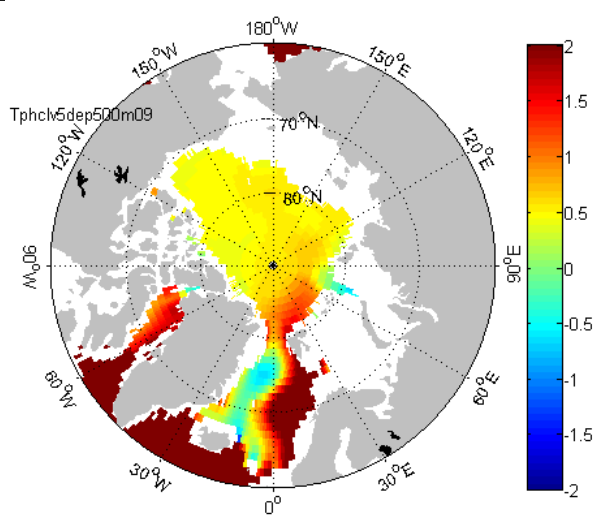


Schauer et al.(2004)
Oct. 1998 (Obs.)

Adding Neptune effect
(Hollway,1992) did help to increase
the current of Atlantic inflow, but
still work not well for the vertical
Temp. profile

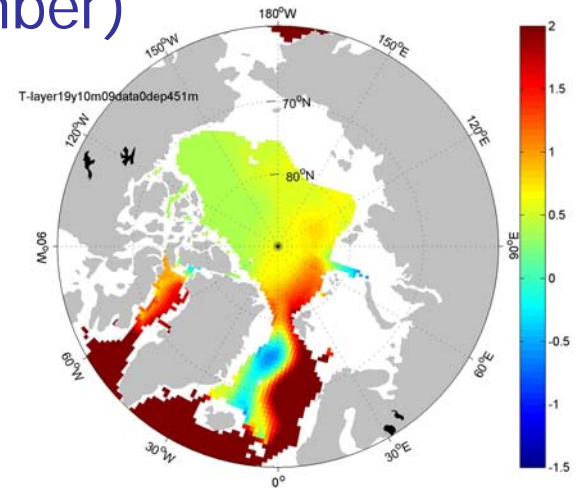


450m sea temperature (September)

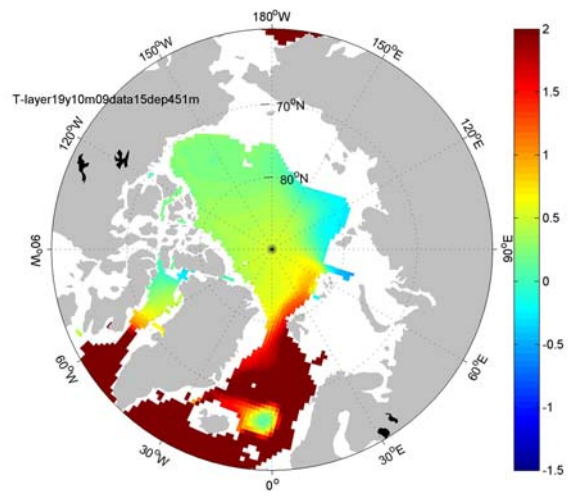


PHC3.0

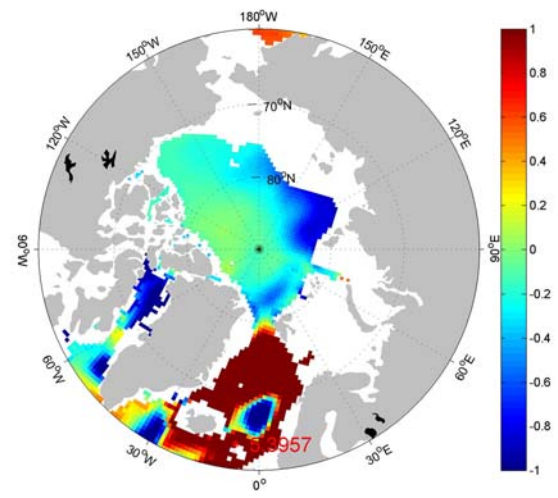
(September)



initial

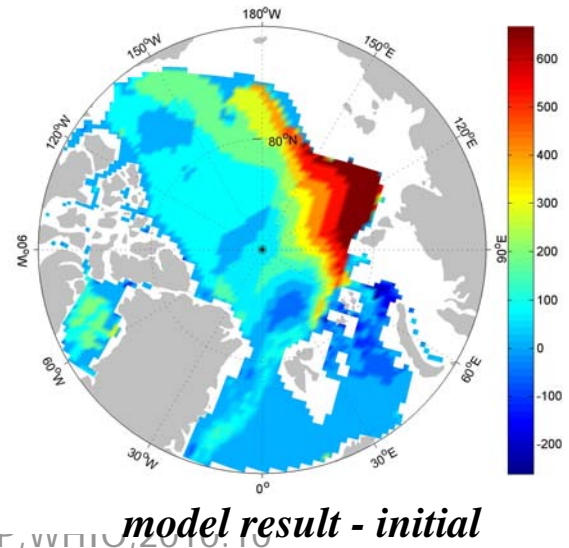
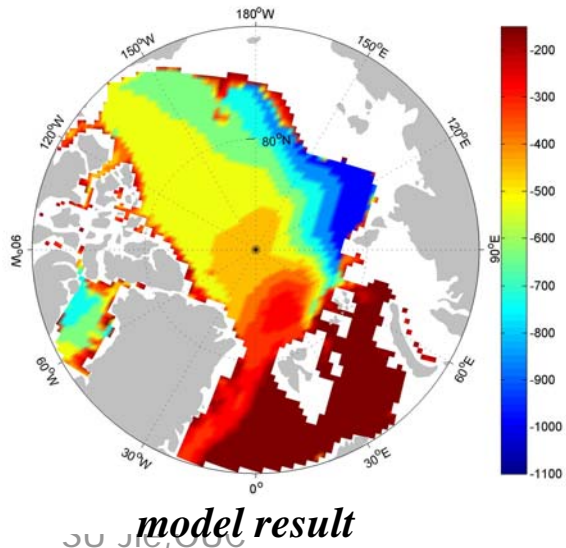
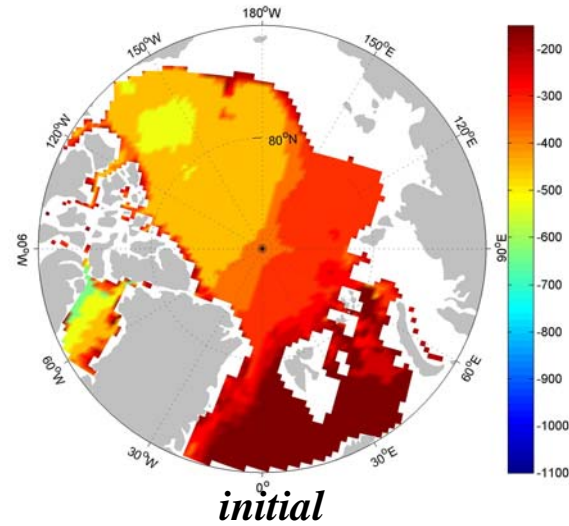
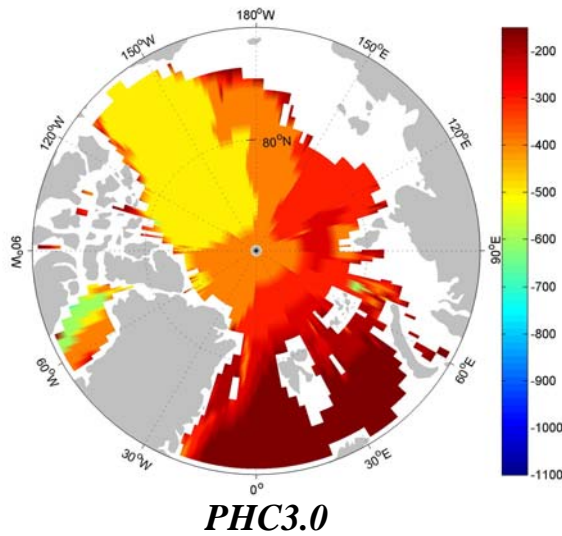


SU JI, 06 *model result*



AOMIP,WHIO,2019.10 *model result - initial*

Max Temp. depth of Middle layer water (September)



Experiments on vertical turbulence problem

- The vertical turbulent coefficient include three items:

- TKE (turbulent closure scheme)

$$A^{vT} = A^{vm} / P_n \quad A^{vm} = C_k l_k \sqrt{e}$$

$$\frac{\partial \bar{e}}{\partial t} = \frac{A^{vm}}{e_3} \left[\left(\frac{\partial u}{\partial k} \right)^2 + \left(\frac{\partial v}{\partial k} \right)^2 \right] - A^{vT} N^2 + \frac{1}{e_3} \frac{\partial}{\partial k} \left[\frac{A^{vm}}{e_3} \frac{\partial \bar{e}}{\partial k} \right] - c_\varepsilon \frac{\bar{e}^{-3/2}}{l_\varepsilon}$$

- DDM (double diffusion mixing)
- Item caused by choosing horizontal mixing scheme

$$D^{lt} = \nabla \cdot (A^{lt} \mathfrak{R} \nabla T)$$

iso-level

$$\mathfrak{R} = \begin{pmatrix} 1 & 0 & -r_1 \\ 0 & 1 & -r_2 \\ -r_1 & -r_2 & r_1^2 + r_2^2 \end{pmatrix}$$

iso-neutral

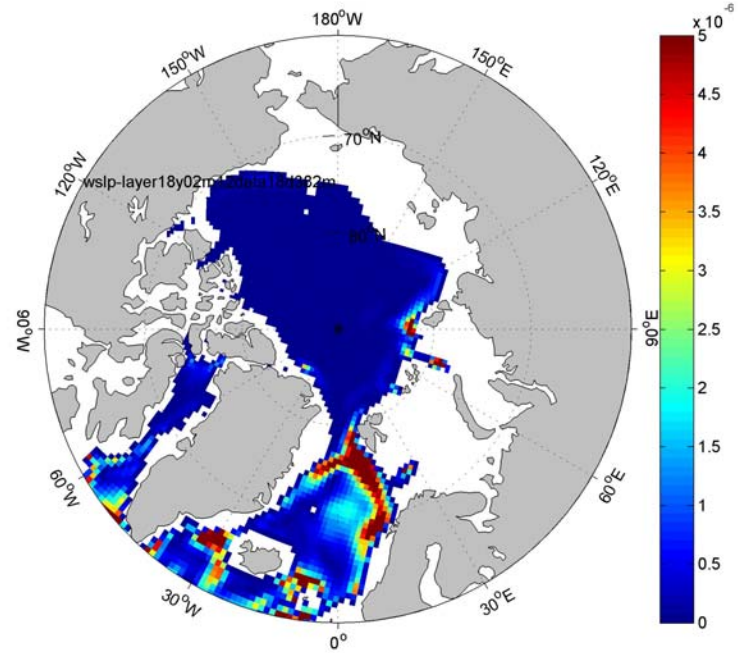
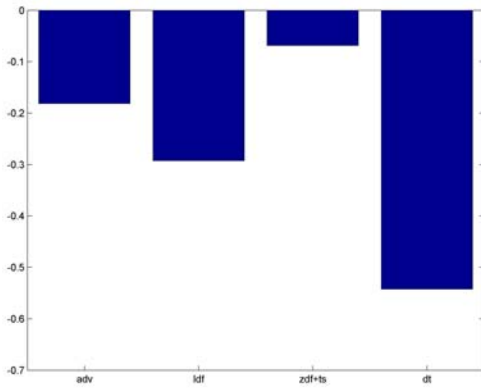
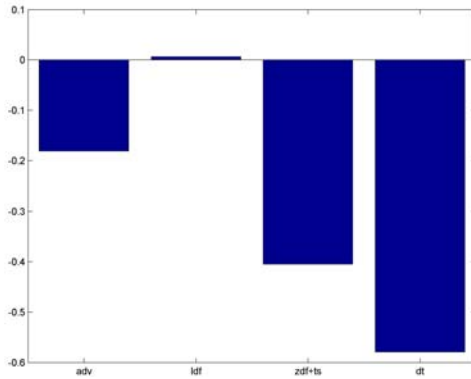
$$\frac{\partial}{\partial k} \left(A^{lt} (r_1^2 + r_2^2) \frac{\partial T}{\partial k} \right)$$

$$r_1 = \left(\frac{\partial \rho}{\partial i} \right) \left(\frac{\partial \rho}{\partial k} \right)^{-1}$$

$$r_2 = \left(\frac{\partial \rho}{\partial j} \right) \left(\frac{\partial \rho}{\partial k} \right)^{-1}$$

Experiments on vertical turbulence problem

1 2 3 all

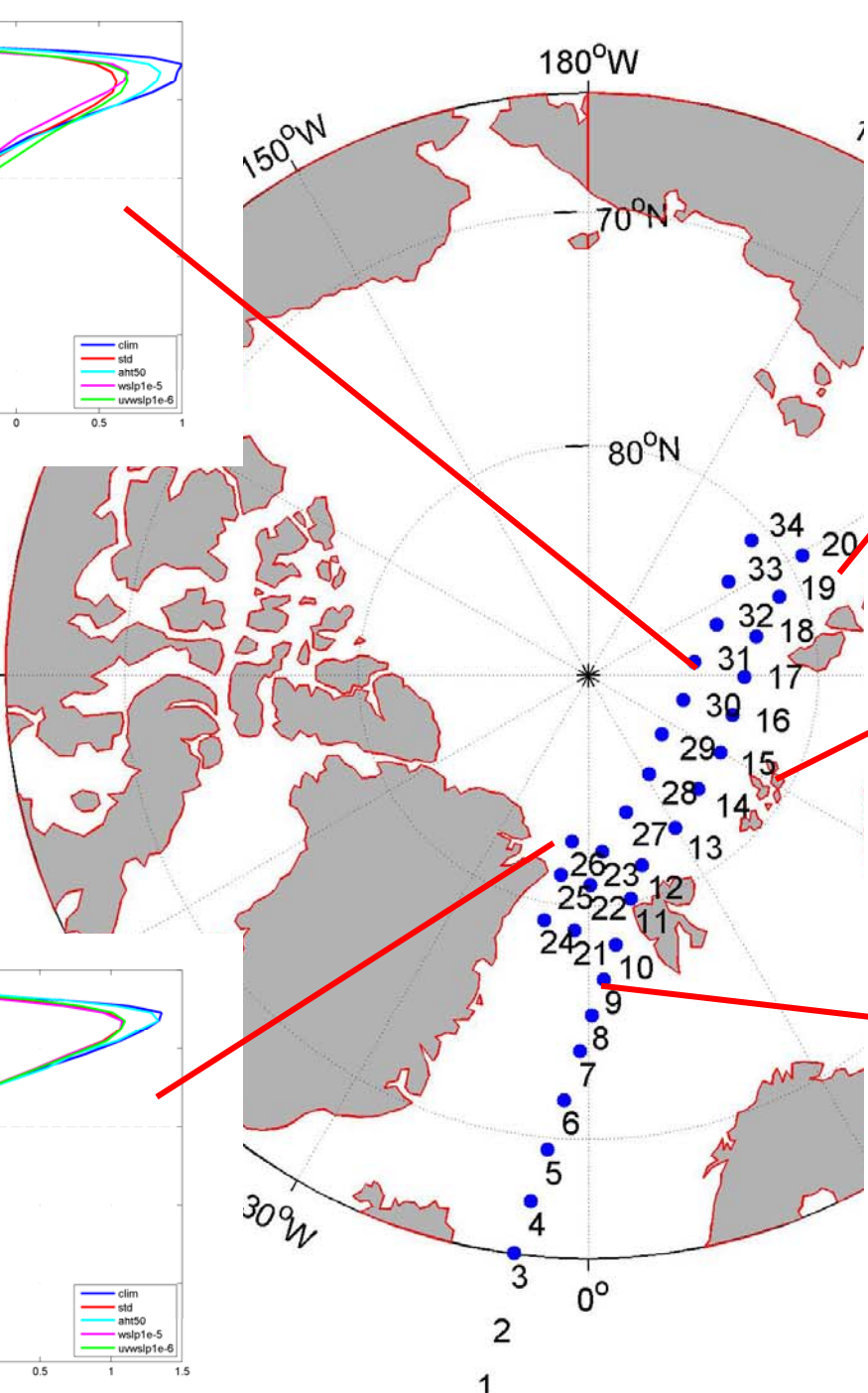
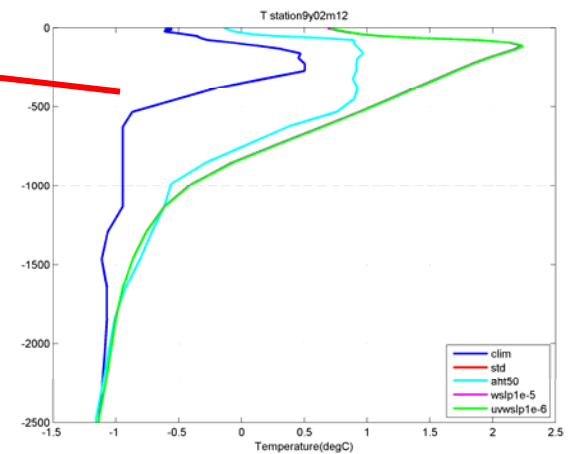
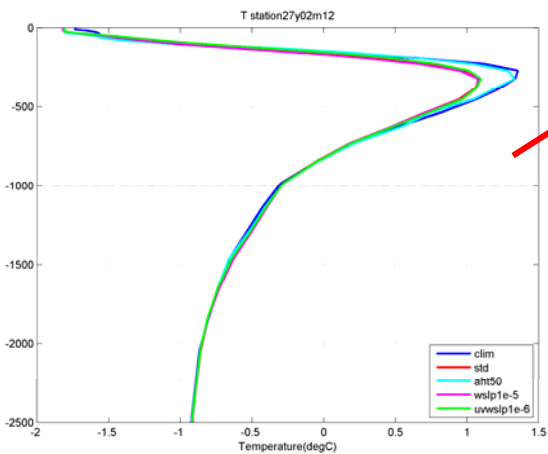
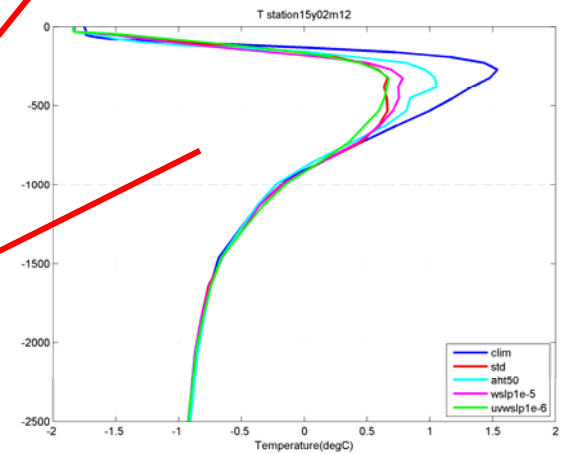
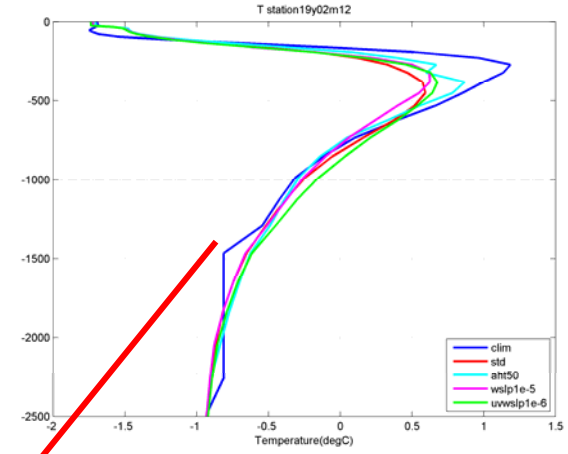
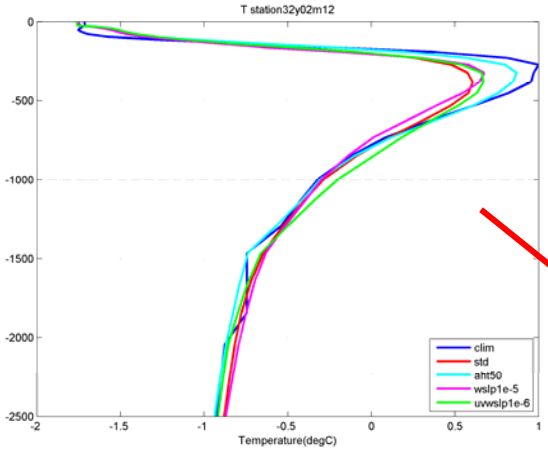


$$\mathfrak{R} = \begin{pmatrix} 1 & 0 & -r_1 \\ 0 & 1 & -r_2 \\ -r_1 & -r_2 & r_1^2 + r_2^2 \end{pmatrix}$$

Exp1: limit wslp to 1e-5

Exp2: limit r1, r2 to 1e-6

Exp3: limit Alt (under ice) to 1e-7



Two years experiments



Summary and discussion

- The modeled sea ice results can represent the seasonal cycle of Arctic sea ice in basic pattern. Ice area agrees well with those calculated by SSMR-SSMI ice concentration.
- The Beaufort gyre and the transpolar drift are showed obviously in the modeled ice drift and upper 100m current field.
- 10y run fails to reproduce the Atlantic layer in ice covered region, especially in Euro-Asian Basin. This problem can be solved by enhance model resolution.
- Using a coarse model, the problem of Atlantic layer is possibly caused by too weak transport though Fram Strait and too strong vertical diffusion which caused by iso-neutral horizontal mixing scheme. Limit the advection diffusion coefficient can partly improve the results.



Thank you
for your
attention!
Any
Question?

