

*Controlled vehicles, en route assimilation,  
conclusions and connections*

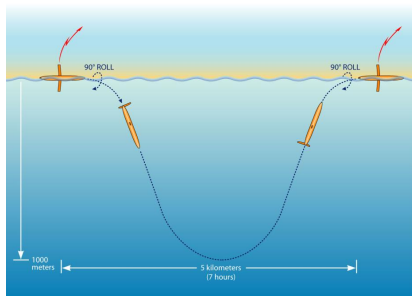
Elaine Spiller, Adam Mallen

Marquette University

## Instrument



## Profile

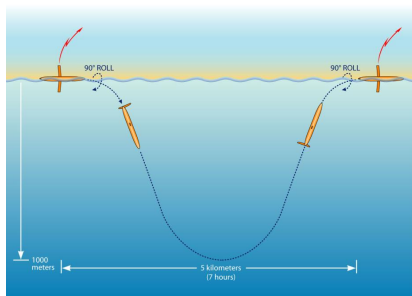


- collect standard (salinity, temperature, pressure) data
- can be fit with chemical/optical/biological sensors
- guided/steered by manipulation of buoyancy and battery location on pre-programmed flight plan,  $\mathbf{u}_c$

Instrument



Profile

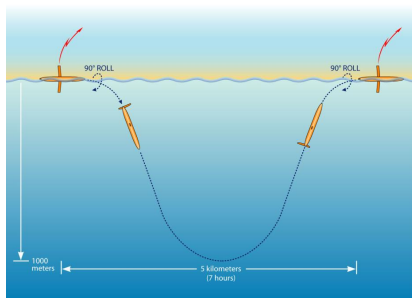


$$\begin{pmatrix} dx/dt \\ dy/dt \\ dz/dt \end{pmatrix} = \mathbf{u}(x, y, z) + \mathbf{u}_c$$

## Instrument



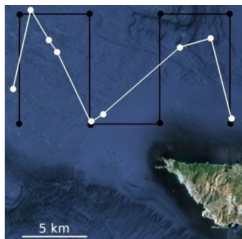
## Profile



- good kinematic model of motion (Leonard & others)  
based on balancing forces during equilibrium glide
- velocity field uncertainty  $\gg$  glider model uncertainty
- path estimates available via dead reckoning – small velocity errors get integrated into large position errors

# Challenges and objectives

## Targeted vs actual surfacings



(Smith *et al* 2011)

## Challenges

- glider path semi-Lagrangian
- ocean velocity,  $\mathbf{u}(x, y, z)$ , needed to plan flight
- data collected along unknown/poorly estimated paths

## Data assimilation objectives:

- probabilistic description of glider paths
  - better estimates of glider paths and uncertainty
- update estimates of  $\mathbf{u}(x, y, z)$  for next flight planning
  - parameterization of model
  - model discovery

# Test problem – Ekman Layer

- $d$  – Ekman layer depth (unknown, a priori)
- $\bar{\mathbf{u}}$ , – background ocean velocity
- $\tau/\rho$  – surface wind stress
- $f$  – Coriolis parameter

$$u(x, y, z) = \bar{u}(x, y) + \frac{\sqrt{2}}{fd} e^{z/d} \left[ \frac{\tau^x(x, y)}{\rho} \cos\left(\frac{z}{d} - \frac{\pi}{4}\right) - \frac{\tau^y(x, y)}{\rho} \sin\left(\frac{z}{d} - \frac{\pi}{4}\right) \right]$$
$$v(x, y, z) = \bar{v}(x, y) + \frac{\sqrt{2}}{fd} e^{z/d} \left[ \frac{\tau^x(x, y)}{\rho} \sin\left(\frac{z}{d} - \frac{\pi}{4}\right) + \frac{\tau^y(x, y)}{\rho} \cos\left(\frac{z}{d} - \frac{\pi}{4}\right) \right]$$

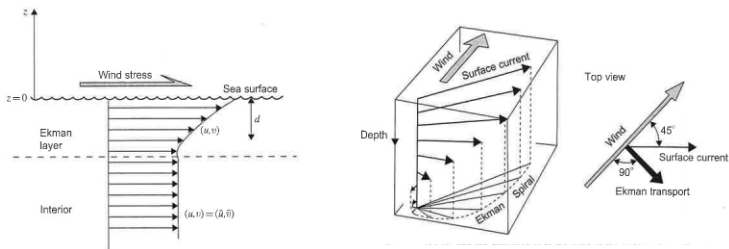


Figure: (Taken from Cushman-Roison 'Introduction to Geophysical Fluid Dynamics')

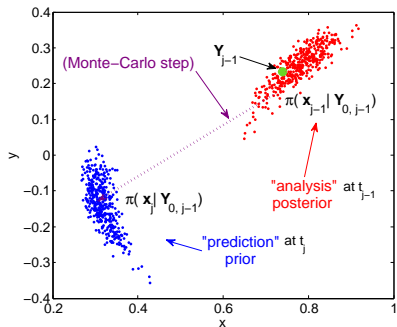
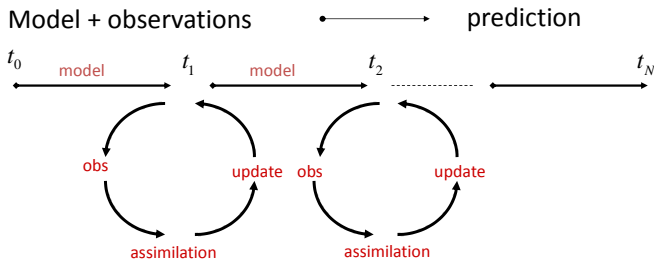
## Experiment:

- twin-experiments, observe “true” surfacing location
- assimilate using particle filter
- use assimilated parameterization of velocity field to guide next flight

## Goals:

- posterior distributions of parameters in velocity model
- probabilistic description of 3D glider paths
- improve glide path and velocity estimates
- reduce and accurately describe uncertainty in those estimates

# Particle filters: from $t_{j-1}$ to $t_j$



discrete approx:

Particles are the support of the discrete approximations to these distributions

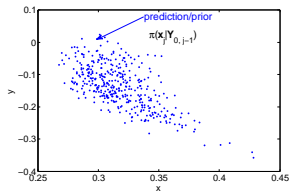
Each particle is associated with a weight,  $w_j(x_j)$



## Particle filters: update/analysis at $t = t_j$

Know (discrete approximation):

$$\pi(x_j | Y_{0,j-1}) \text{ (from last page)}$$



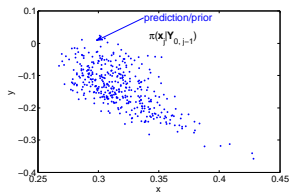
# Particle filters: update/analysis at $t = t_j$

Know (discrete approximation):

$$\pi(x_j | Y_{0,j-1}) \text{ (from last page)}$$

Bayes:

$$\pi(x_j | Y_{0,j}) \propto g(Y_j | x_j) \pi(x_j | Y_{0,j-1})$$



# Particle filters: update/analysis at $t = t_j$

Know (discrete approximation):

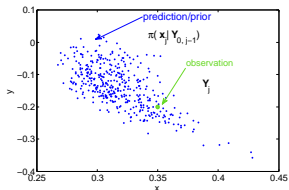
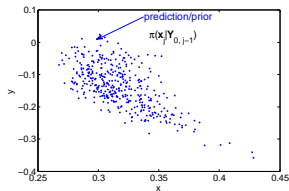
$$\pi(x_j | Y_{0,j-1}) \text{ (from last page)}$$

Bayes:

$$\pi(x_j | Y_{0,j}) \propto g(Y_j | x_j) \pi(x_j | Y_{0,j-1})$$

Likelihood:

$$g(Y|x) = \exp\left[\frac{H(x) \cdot Y}{\theta^2} - \frac{|H(x)|^2}{2\theta^2}\right]$$



# Particle filters: update/analysis at $t = t_j$

Know (discrete approximation):

$$\pi(x_j | Y_{0,j-1}) \text{ (from last page)}$$

Bayes:

$$\pi(x_j | Y_{0,j}) \propto g(Y_j | x_j) \pi(x_j | Y_{0,j-1})$$

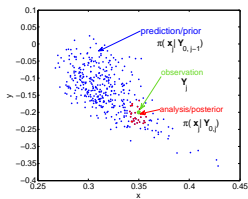
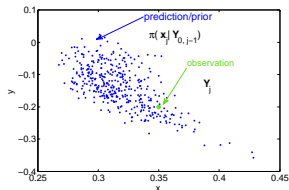
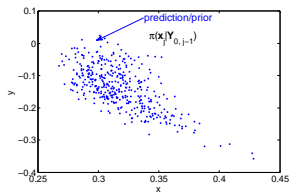
Likelihood:

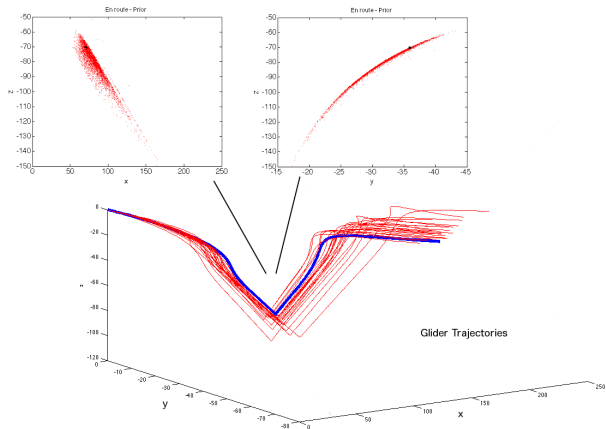
$$g(Y|x) = \exp\left[\frac{H(x) \cdot Y}{\theta^2} - \frac{|H(x)|^2}{2\theta^2}\right]$$

Update (discrete Bayes):

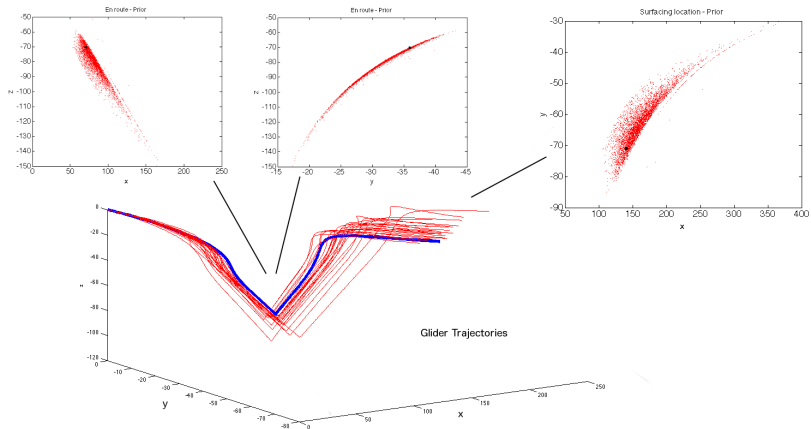
$$w_j(x_j) \propto g(Y_j | x_j) w_j^p(x_j)$$

$$\pi(x_j | Y_{0,j}) = \{x_j, w_j(x_j)\}$$

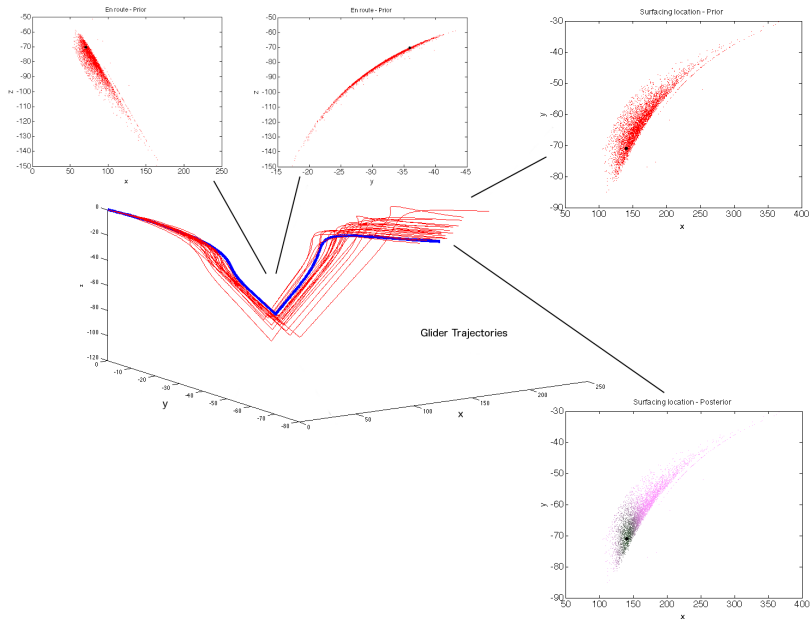


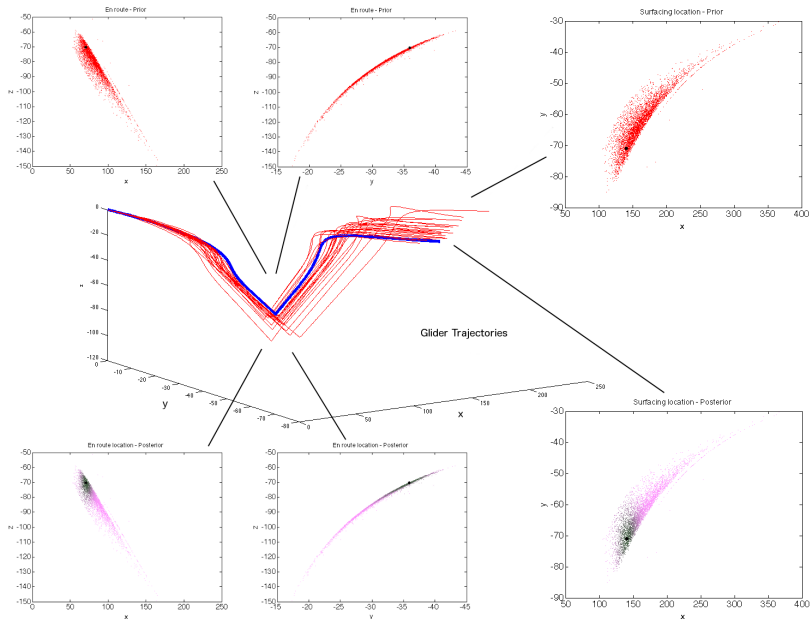


# Glider DA



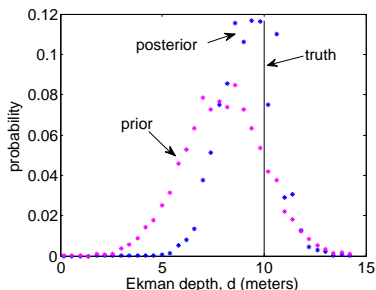
# Glider DA







# Assimilated Ekman depth



- quantify uncertainty in parameterization of  $\mathbf{u}(x, y, z)$
- iterate, posterior of first dive becomes prior of second
- reduced uncertainty/improved accuracy of glider paths in subsequent dives
- use estimates of  $\mathbf{u}$  to help choose  $\mathbf{u}_c$

- twin-experiments
  - observe “true” surfacing location
  - observe vertical,  $z(t)$ , paths
- spatially or time varying background velocity and/or wind stress
- use LADA with model-discovery to determine distribution of most-likely velocity fields
- probabilistic description of 3D glider paths

# En-route data assimilation

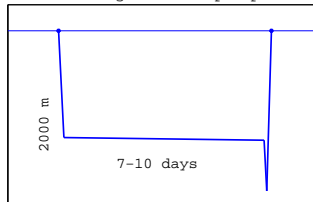
## Argo float



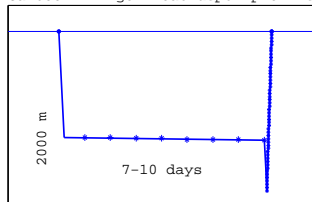
- Lagrangian instruments collect data en route (salinity, temperature, pressure)
- Observations depend on unknown drifter paths
- What to do with that data?

# Float depth profile

Cartoon: Argo float depth profile



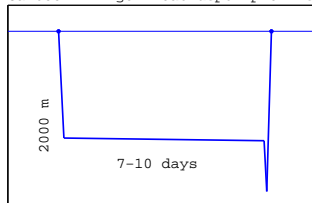
Cartoon: Argo float depth profile



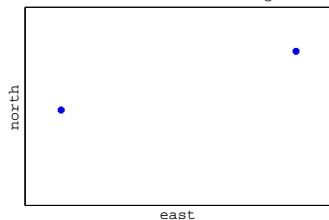
- 7-10 day float results in O(10)-O(100) km traveled
- high frequency data in dive/ascent just before surfacing in water column beneath “surfacing location”
- low frequency en-route measurements at depth, no latitude/longitude information
- en-route measurements averaged, not used in assimilation

# Float depth and overview

Cartoon: Argo float depth profile



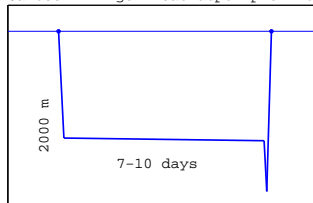
Cartoon: Overview of surfacing locations



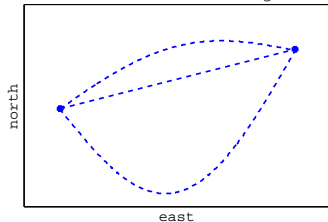
- Lagrangian DA can help ascertain velocities w/o averaging

# Float depth and overview

Cartoon: Argo float depth profile



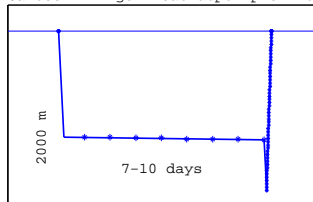
Cartoon: Overview of surfacing locations



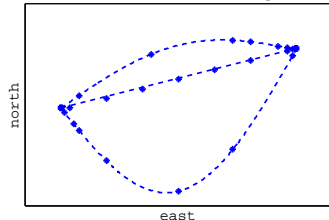
- Some possible Lagrangian paths

# Float depth and overview

Cartoon: Argo float depth profile



Cartoon: Overview of surfacing locations

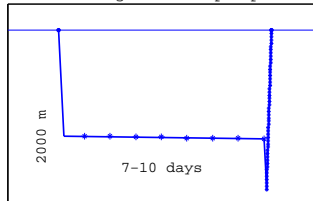


- need path & speed for subsurface observation locations

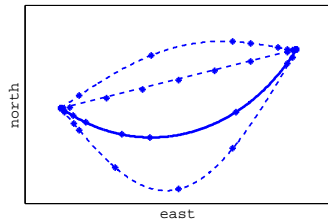


# Float depth and overview

Cartoon: Argo float depth profile



Cartoon: Overview of surfacing locations



- Can en-route observations help Lagrangian DA?

- aid in resolving Lagrangian structures
- assimilating data into high resolution models
- avoiding averaging via determining en-route data collection locales along paths which cross multiple grid cells

## Non-dimensional velocity fields

$$\begin{aligned}\frac{\partial u}{\partial t} &= v - \frac{\partial h}{\partial x} \\ \frac{\partial v}{\partial t} &= -u - \frac{\partial h}{\partial y} \\ \frac{\partial h}{\partial t} &= -\frac{\partial u}{\partial x} - \frac{\partial v}{\partial y}\end{aligned}$$

## Lagrangian trajectories

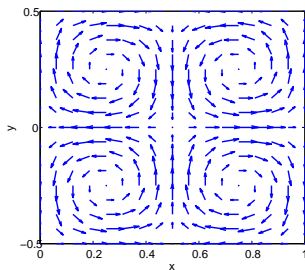
$$\begin{aligned}\dot{x}(t) &= u[x(t), y(t), t] \\ \dot{y}(t) &= v[x(t), y(t), t]\end{aligned}$$

## Decomposition into Fourier Modes

$$\begin{aligned}u(x, y, t) &= -2\pi \sin(2\pi x) \cos(2\pi y) u_0 + \cos(2\pi y) u_1(t) \\ v(x, y, t) &= 2\pi \cos(2\pi x) \sin(2\pi y) u_0 + \cos(2\pi y) v_1(t) \\ h(x, y, t) &= \sin(2\pi x) \sin(2\pi y) u_0 + \sin(2\pi y) h_1(t)\end{aligned}$$

# Cellular flow field

If  $u_1 = v_1 = h_1 = 0$ ,  
flow field is constant & tracers  
stay w/in cells



# Cellular flow field

If  $u_1 = v_1 = h_1 = 0$ ,  
flow field is constant & tracers  
stay w/in cells

otherwise,

$$\dot{u}_o = 0$$

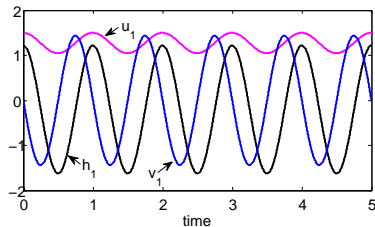
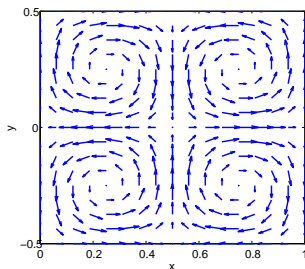
$$\dot{v}_1 = -u_1 - 2\pi h_1$$

$$\dot{u}_1 = v_1$$

$$\dot{h}_1 = 2\pi v_1$$

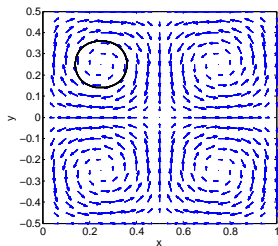
w/initial conditions

$$[u_o(0), u_1(0), v_1(0), h_1(0)]$$



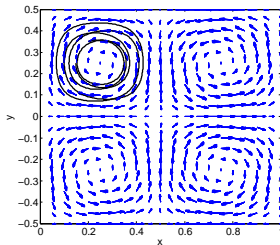
# A few trajectories

Steady



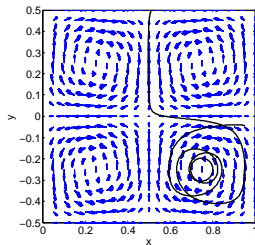
$$\begin{aligned}u_1(0) &= 0 \\v_1(0) &= 0 \\h_1(0) &= 0 \\x(0) &= 0.2 \\y(0) &= 0.3\end{aligned}$$

Center



$$\begin{aligned}u_1(0) &= 0.5 \\v_1(0) &= 0.5 \\h_1(0) &= 0.5 \\x(0) &= 0.2 \\y(0) &= 0.3\end{aligned}$$

Saddle



$$\begin{aligned}u_1(0) &= 0.2 \\v_1(0) &= 1.3 \\h_1(0) &= 1.4 \\x(0) &= 0.51 \\y(0) &= 0.498\end{aligned}$$

## Test problem:

- $u_1(0) = v_1(0) = h_1(0) = 0.5$ ,  $x(0) = .2$ ,  $y(0) = .3$
- broad priors on  $(u_1, v_1, h_1)$ , tight on  $(x, y)$  at  $t=0$
- run to  $t = T$  (1 period of coefficients)
- 5 noisy observations of drifter

## Test problem:

- $u_1(0) = v_1(0) = h_1(0) = 0.5$ ,  $x(0) = .2$ ,  $y(0) = .3$
- broad priors on  $(u_1, v_1, h_1)$ , tight on  $(x, y)$  at  $t=0$
- run to  $t = T$  (1 period of coefficients)
- 5 noisy observations of drifter

## Goal:

- learn about  $u_1(0)$ ,  $v_1(0)$ ,  $h_1(0)$  from Lagrangian observations



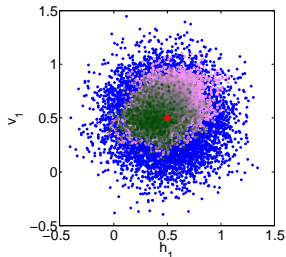
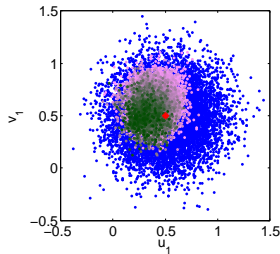
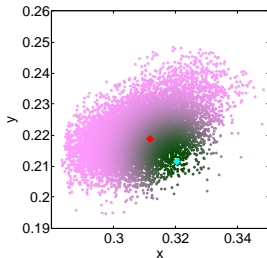
# Particle filter for standard LADA

## Test problem:

- $u_1(0) = v_1(0) = h_1(0) = 0.5$ ,  $x(0) = .2$ ,  $y(0) = .3$
- broad priors on  $(u_1, v_1, h_1)$ , tight on  $(x, y)$  at  $t=0$
- run to  $t = T$  (1 period of coefficients)
- 5 noisy observations of drifter

## Goal:

- learn about  $u_1(0), v_1(0), h_1(0)$  from Lagrangian observations



## *En route Lagrangian data – a test problem*

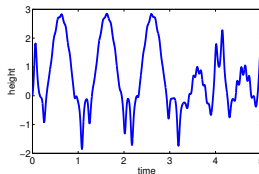
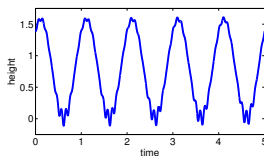
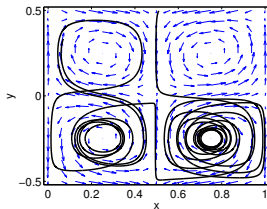
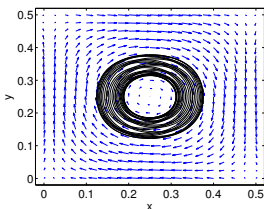
- **Idea** treat height,  $h(x, y, u_1, v_1, h_1)$ , as proxy for temperature – typical quantity measured en route

## *En route Lagrangian data – a test problem*

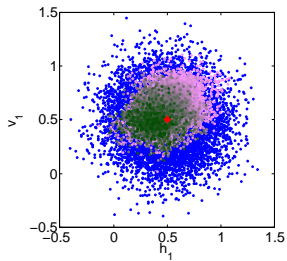
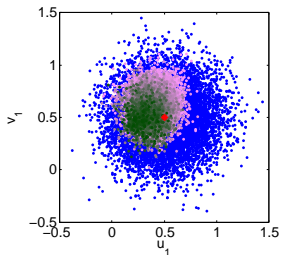
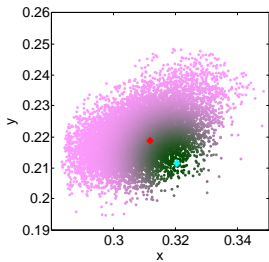
- **Idea** treat height,  $h(x, y, u_1, v_1, h_1)$ , as proxy for temperature – typical quantity measured en route
- **Sample** height,  $\hat{h}(t) = h(x^d(t), y^d(t), t) + \textit{noise}$  between “surfacing”, e.g. traditional observation instants  $t_j$

# En route Lagrangian data – a test problem

- **Idea** treat height,  $h(x, y, u_1, v_1, h_1)$ , as proxy for temperature – typical quantity measured en route
- **Sample** height,  $\hat{h}(t) = h(x^d(t), y^d(t), t) + \textit{noise}$  between “surfacings”, e.g. traditional observation instants  $t_j$

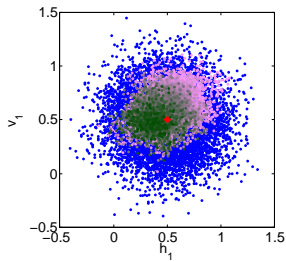
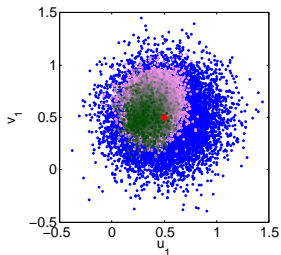
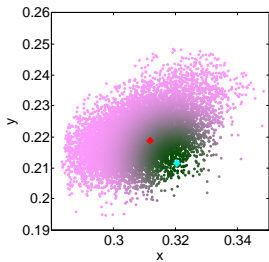


## "traditional" LADA:

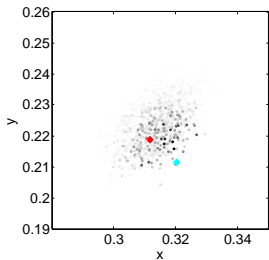


# Particle filter w/en route observations

“traditional” LADA:

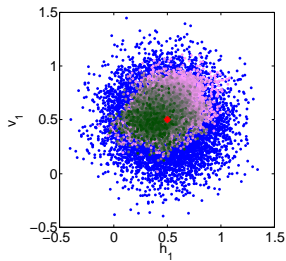
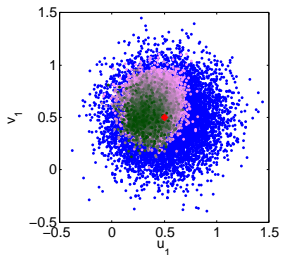
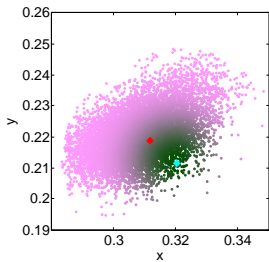


en route LADA:

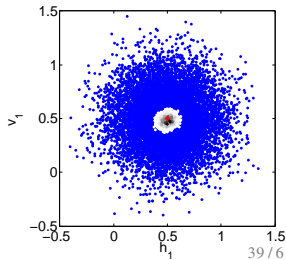
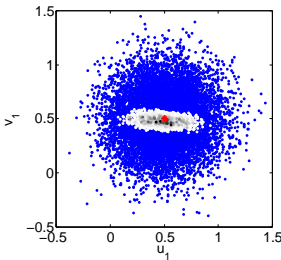
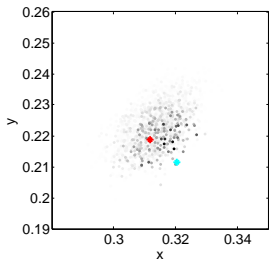


# Particle filter w/en route observations

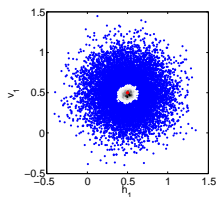
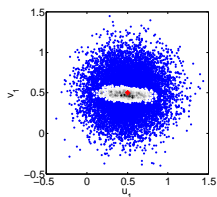
“traditional” LADA:



en route LADA:



# Characterizing improvement



Characterizing improvement:

compare covariance matrices of prior and posterior distribution

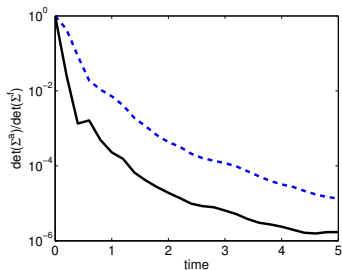
$$d_s(t) = \text{tr}[\mathbf{I} - \Sigma_F^a(t)(\Sigma_F^f)^{-1}] \quad (\text{Zupanski, 2007})$$

$$r(t) = \det(\Sigma_F^a(t)) / \det(\Sigma_F^f)$$

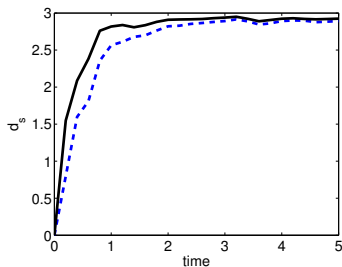
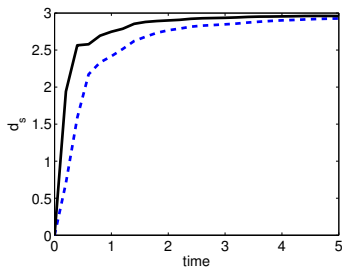
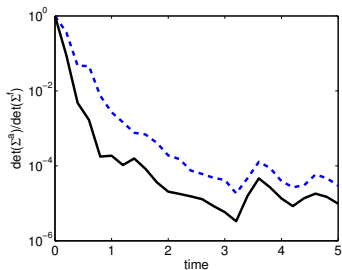


# Improvements w/assimilating en-route data

left column: center



right column: saddle



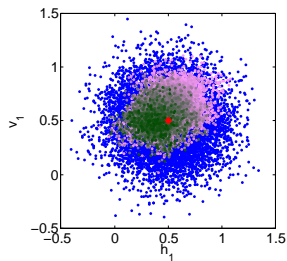
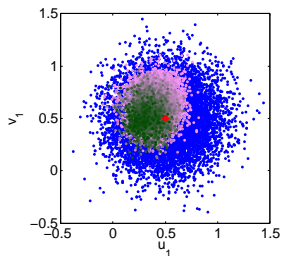
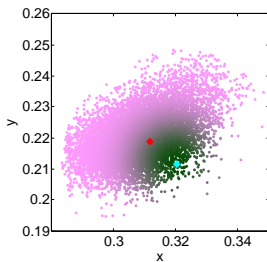
(S, Apte, Jones, submitted 2012)

- Consider two-layer or 3D model w/observable-at-depth, spatially dependent variable
  - collect en-route observations on bottom layer
  - traditional Lagrangian observations, less frequent
- Improvements from en-route assimilation?
- Can we estimate Lagrangian paths at depth?
- How dependent is this on coupling strength?

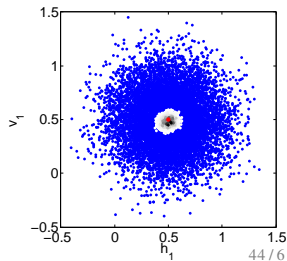
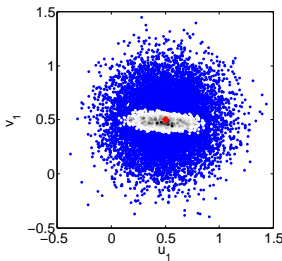
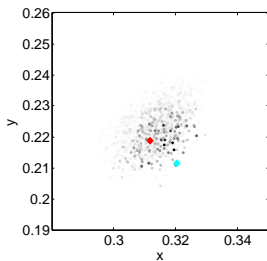
# Conclusions and Connections

# Observing unknown locations: en-route assimilation

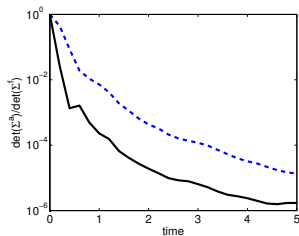
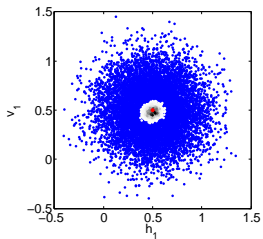
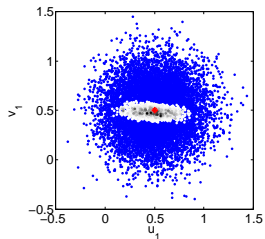
“traditional” LADA:



en route LADA:

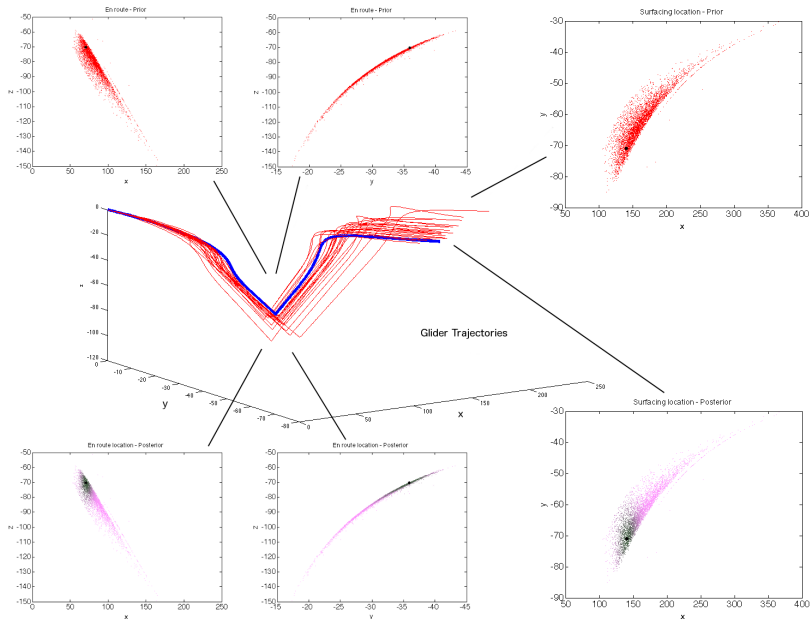


# Observing unknown locations: en-route assimilation

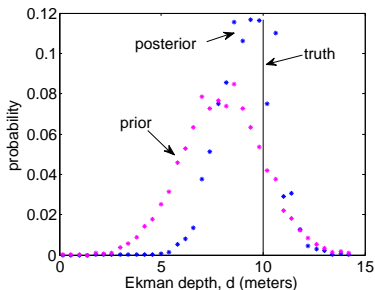


- developed strategies which **fully exploit data collected at unknown locations**
- marked improvement over assimilating only instrument location data

# Subsurface data via controlled vehicles: gliders

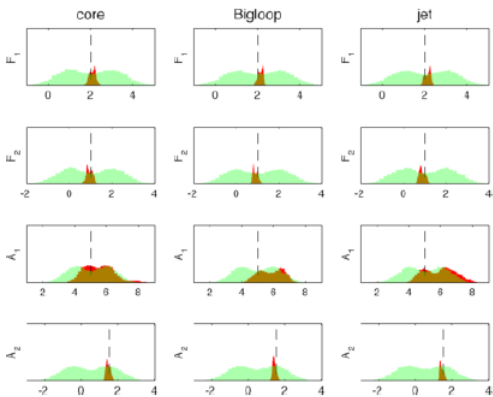


# Subsurface data via controlled vehicles: gliders



- probabilistic description of 3D glider paths
- improve glide path and z-dependent velocity estimates
- reduce and accurately describe uncertainty in those estimates

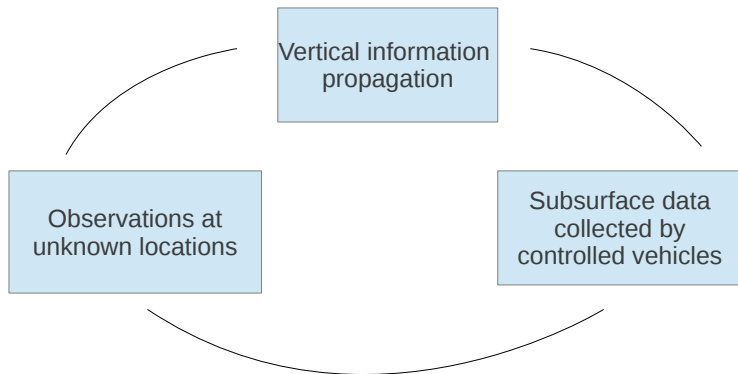
# Vertical information propagation



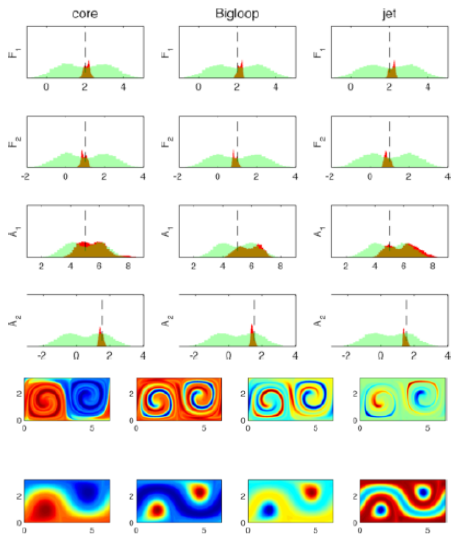
- Ambiguity resolved by more observations and, in some cases, uncertainty, but generally not inaccuracy
- Baroclinic perturbations to barotropic flow are not easily resolved by observations on more layers



## *Connections: between DA projects*

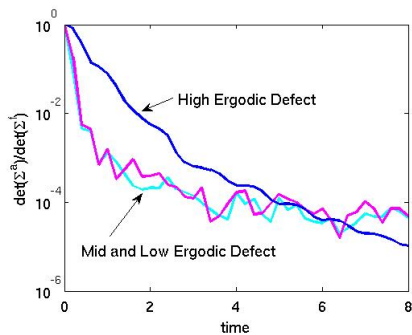


# Connections: vertical information and Koopman operator



- Identified relationship between accuracy/uncertainty and complexity of coherent sets as encoded in Koopman eigenfunctions
- Need to further explore and understand this relationship

# Connections: complexity measures & Lagrangian DA



three different trajectories, each w/different ergodicity defect

- lower values mean greater information content through DA
- how does complexity of trajectories relate to information gain through assimilation?

## Two models:

- **kinematic**  
3D velocity non-divergent but no dynamics
- **CFD**  
nonlinear numerical model, Navier-Stokes equations

## Two models:

- **kinematic**  
3D velocity non-divergent but no dynamics
- **CFD**  
nonlinear numerical model, Navier-Stokes equations

## Strategy:

- treat **CFD** as “reality” and take observations
- treat **kinematic** model as “model”

## Two models:

- **kinematic**  
3D velocity non-divergent but no dynamics
- **CFD**  
nonlinear numerical model, Navier-Stokes equations

## Strategy:

- treat **CFD** as “reality” and take observations
- treat **kinematic** model as “model”

## Problem:

- kinematic model can't reproduce CFD trajectories even with “best” choice of parameters

## Two models:

- **kinematic**  
3D velocity non-divergent but no dynamics
- **CFD**  
nonlinear numerical model, Navier-Stokes equations

## Strategy:

- treat **CFD** as “reality” and take observations
- treat **kinematic** model as “model”

## Problem:

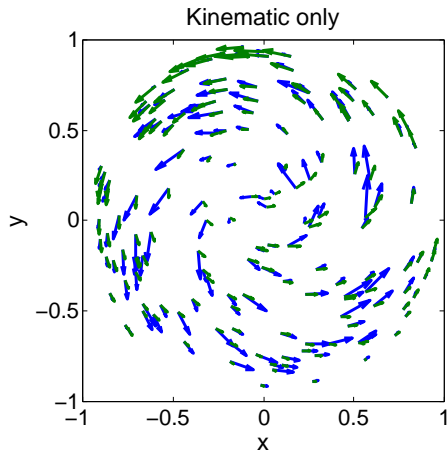
- kinematic model can't reproduce CFD trajectories even with “best” choice of parameters

## Assimilate:

- data from CFD to “discover” model bias

$$\mathbf{u}_{CFD}(x, y, z) = \mathbf{u}_{KM}(x, y, z) + \mathbf{u}_B(x, y, z)$$

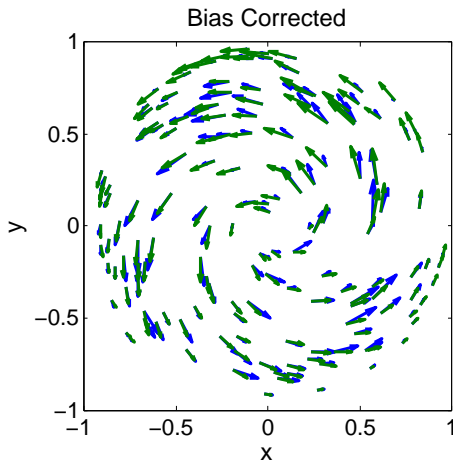
# Rotating can: comparison of velocity fields



- CFD field and "best fit" Kinematic field
- $(u, v)$  projection of top of can,  $z > 0.8$



## Rotating can: comparison of velocity fields



- CFD field and bias corrected kinematic field
- $(u, v)$  projection of top of can,  $z > 0.8$

- Kinematic model

$$\frac{dx}{dt} = u_{KM}, \quad \frac{dy}{dt} = v_{KM}, \quad \frac{dz}{dt} = w_{KM}$$

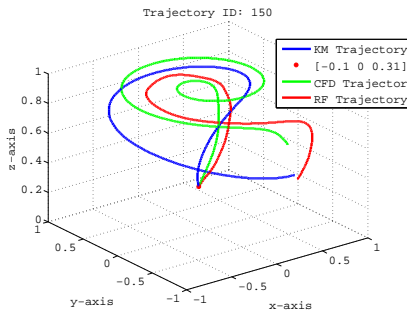
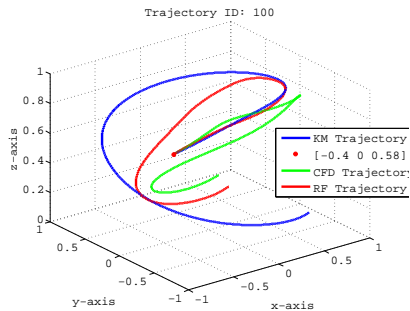
- CFD model

$$\frac{dx}{dt} = u_{CFD}, \quad \frac{dy}{dt} = v_{CFD}, \quad \frac{dz}{dt} = w_{CFD}$$

- Bias corrected kinematic model (labeled RF)

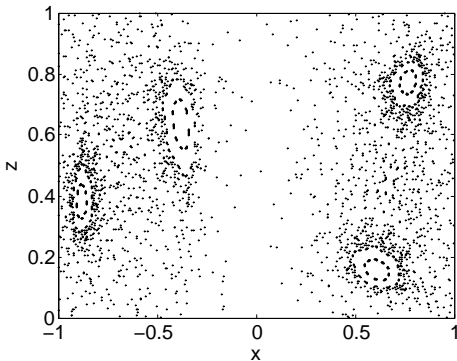
$$\frac{dx}{dt} = u_{KM} + \hat{u}_B, \quad \frac{dy}{dt} = v_{KM} + \hat{v}_B, \quad \frac{dz}{dt} = w_{KM} + \hat{w}_B$$

# Rotating can: comparison of trajectories

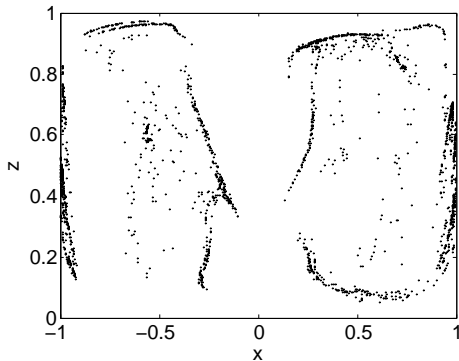


# *Proof of concept: Poincaré Sections, bias corrected model*

kinematic model

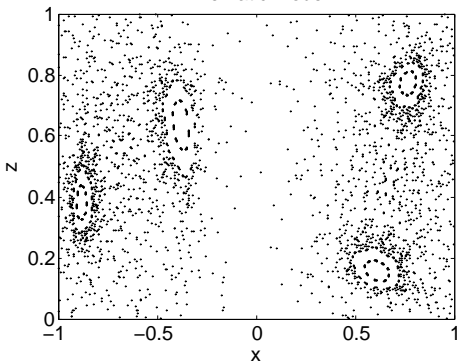


bias corrected

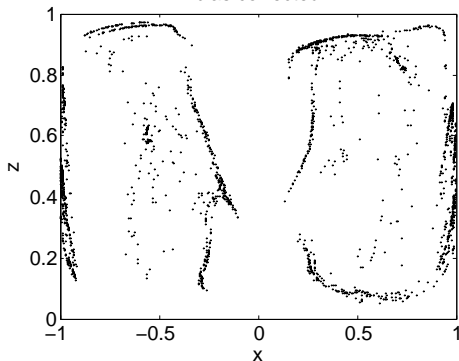


# *Proof of concept: Poincaré Sections, bias corrected model*

kinematic model



bias corrected



- what are we seeing on the right?
- lots of computational experiments to be done
- what is we move to non-steady case?