

The 2010 Haiti earthquake disaster: The ShakeMap that could have been...

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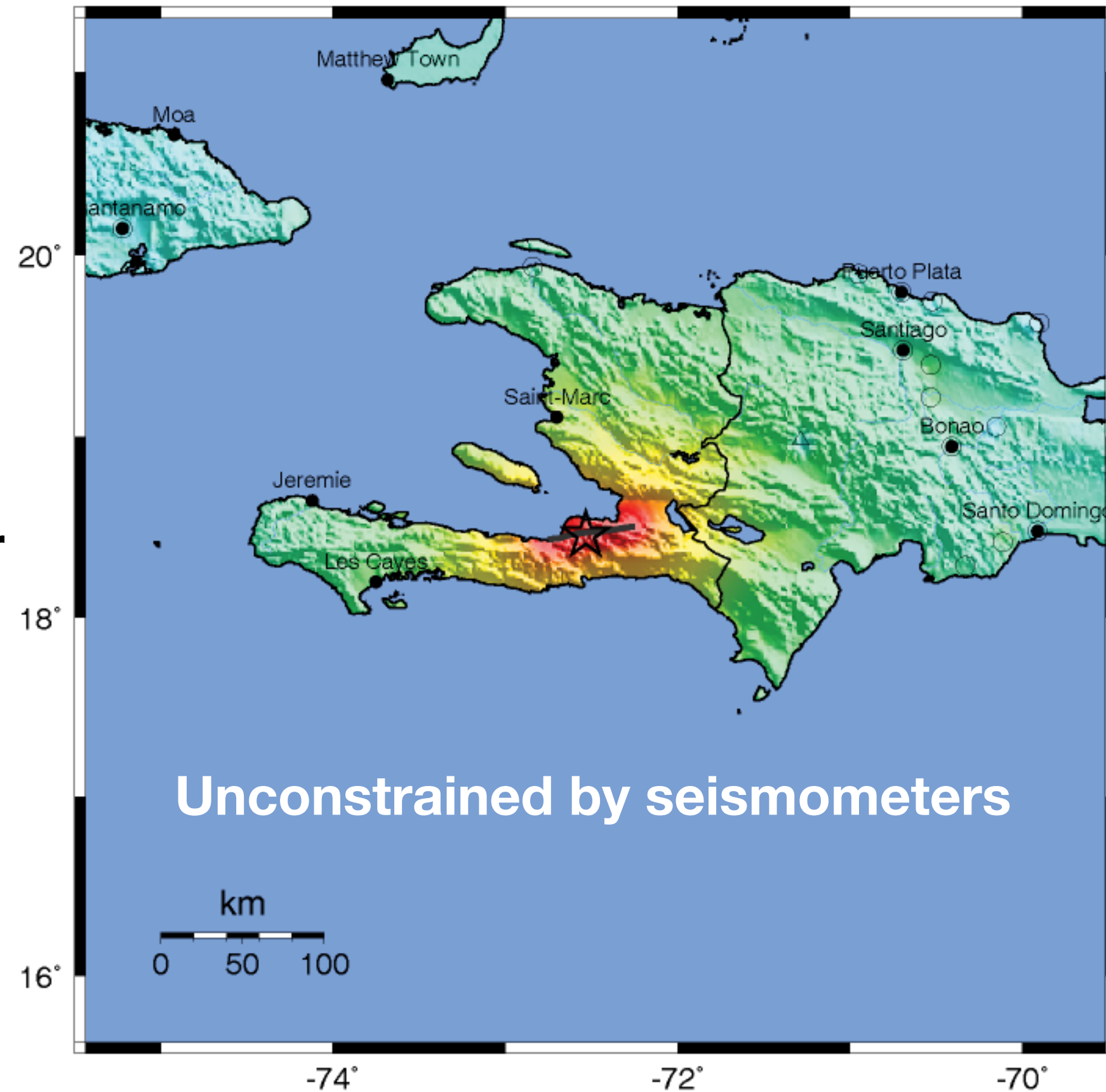
The January 12, 2010 M_W 7 Earthquake Disaster in Haiti

Due to the lack of seismometers during the 2010 Haiti earthquake:

Overestimated shaking intensity near the epicenter

Underestimated the extent of potentially damaging shaking

(a) USGS ShakeMap : HAITI REGION
Tue Jan 12, 2010 21:53:10 GMT M 7.0 N18.46 W72.53 Depth: 13.0km ID:2010rja6

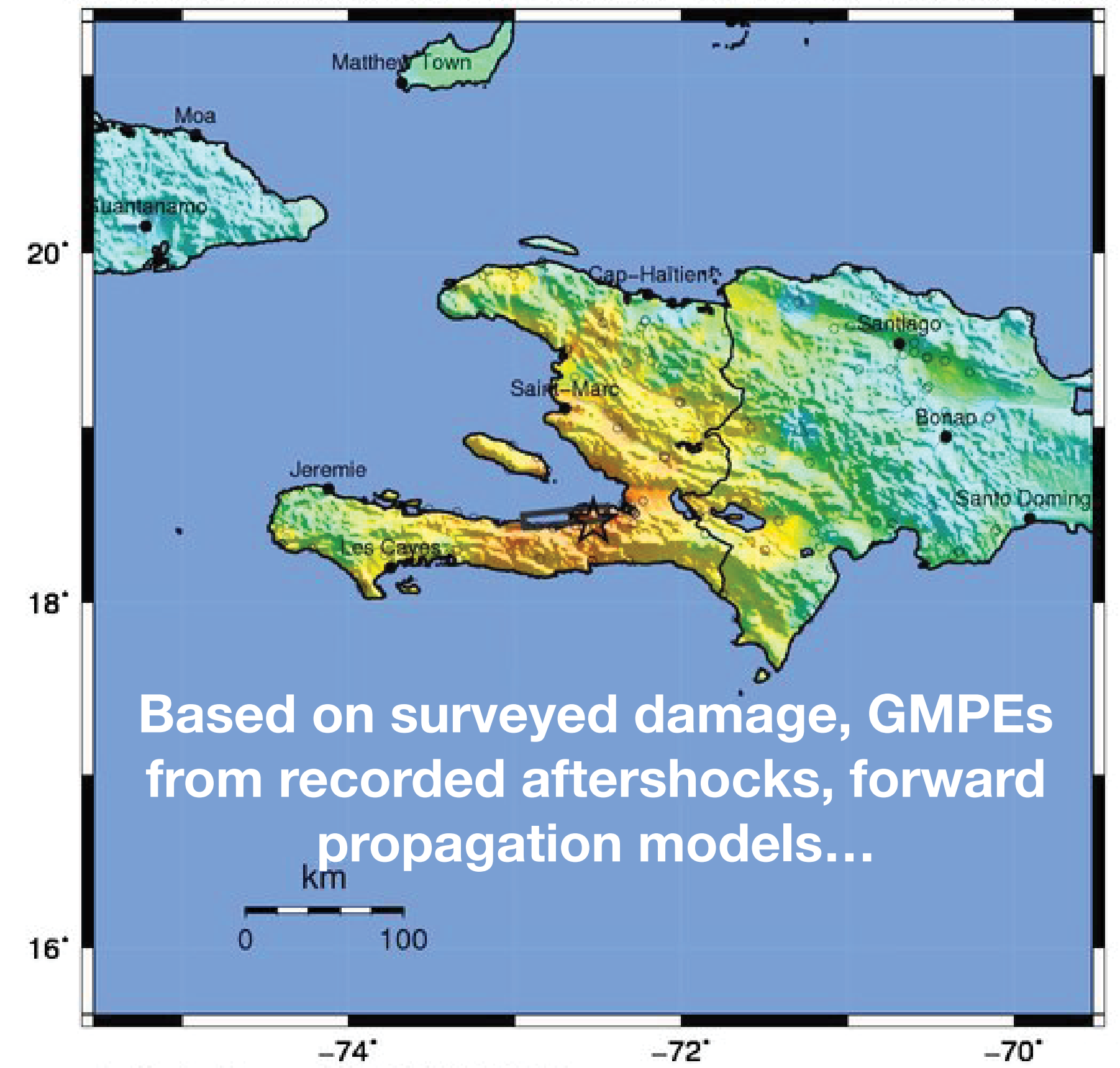


Map Version 7 Processed Wed Jan 13, 2010 06:53:11 PM MST -- NOT REVIEWED BY HUMAN

PERCEIVED SHAKING	Not felt	Weak	Light	Moderate	Strong	Very strong	Severe	Violent	Extreme
POTENTIAL DAMAGE	none	none	none	Very light	Light	Moderate	Moderate/Heavy	Heavy	Very Heavy
PEAK ACC.(%g)	<0.17	0.17-1.4	1.4-3.9	3.9-9.2	9.2-18	18-34	34-65	65-124	>124
PEAK VEL.(cm/s)	<0.1	0.1-1.1	1.1-3.4	3.4-8.1	8.1-16	16-31	31-60	60-116	>116
INSTRUMENTAL INTENSITY	I	II-III	IV	V	VI	VII	VIII	IX	X+

ShakeMap generated more than a day after the earthquake

(b) USGS ShakeMap : Haiti
Jan 12, 2010 21:53:10 UTC M 7.0 N18.46 W72.53 Depth: 13.0km ID:20100112215310



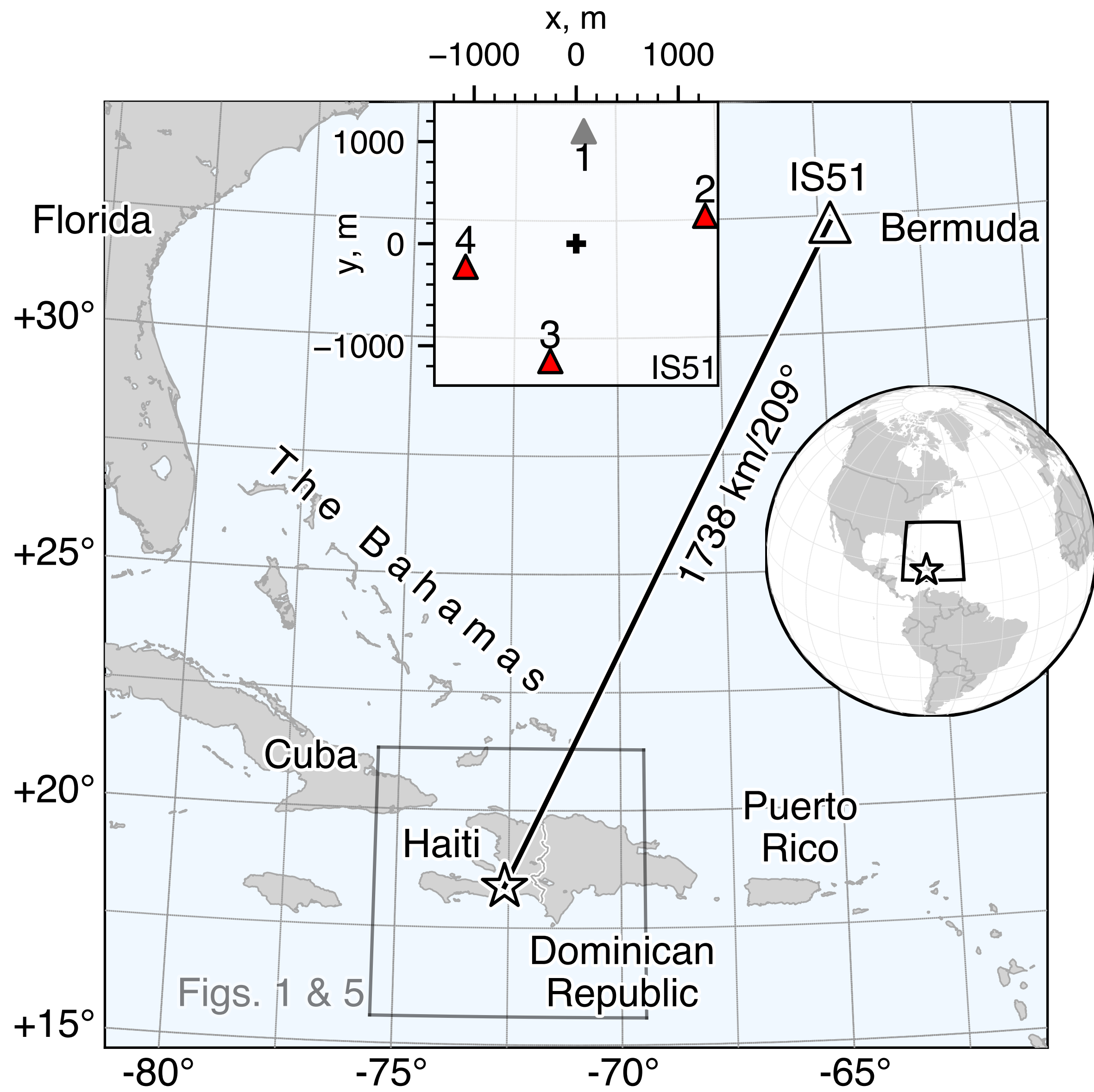
Map Version 1 Processed 2017-01-27 03:39:48 UTC

PERCEIVED SHAKING	Not felt	Weak	Light	Moderate	Strong	Very strong	Severe	Violent	Extreme
POTENTIAL DAMAGE	none	none	none	Very light	Light	Moderate	Mod./Heavy	Heavy	Very Heavy
PEAK ACC.(%g)	<0.05	0.3	2.8	6.2	12	22	40	75	>139
PEAK VEL.(cm/s)	<0.02	0.1	1.4	4.7	9.6	20	41	86	>178
INSTRUMENTAL INTENSITY	I	II-III	IV	V	VI	VII	VIII	IX	X+

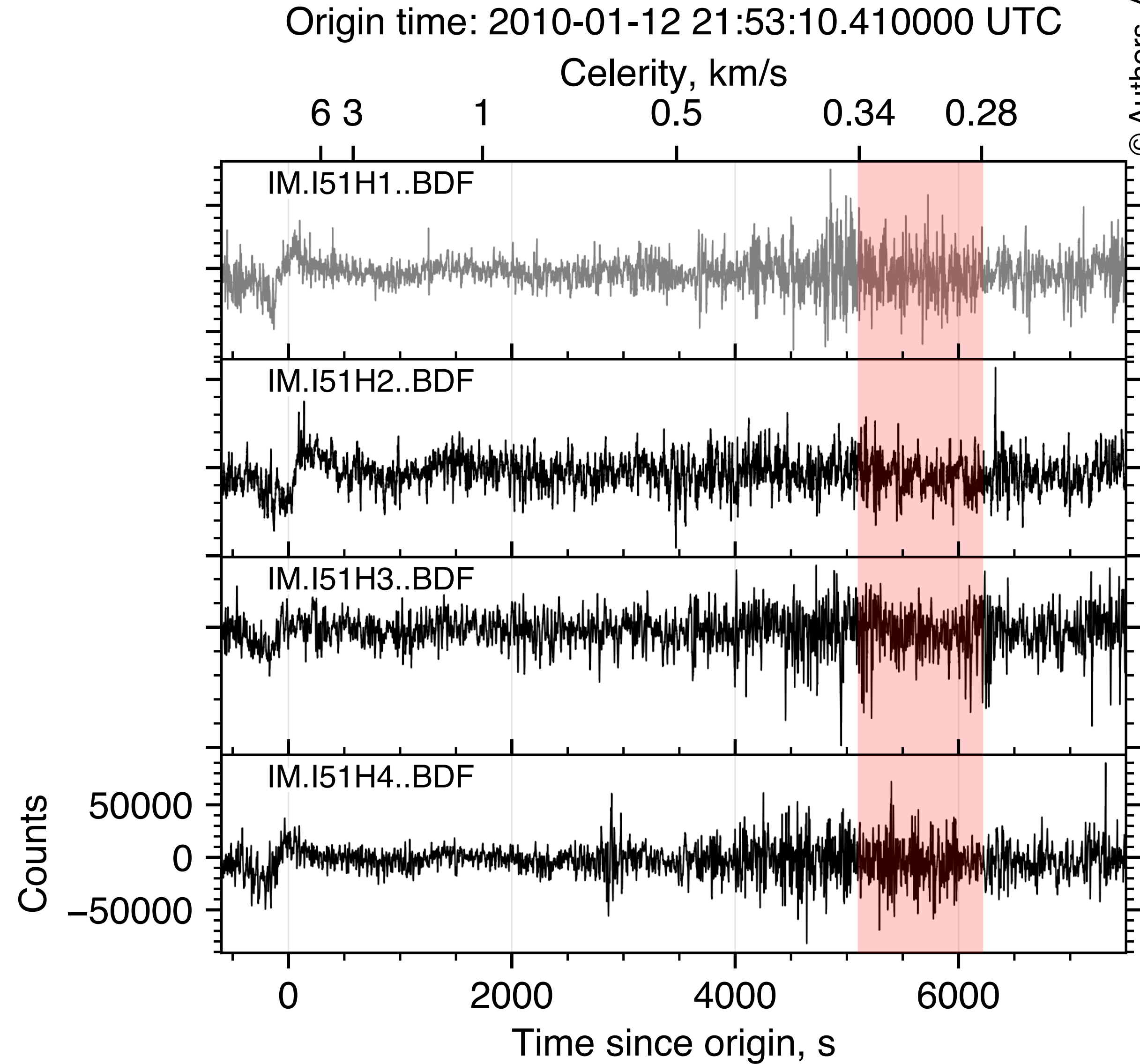
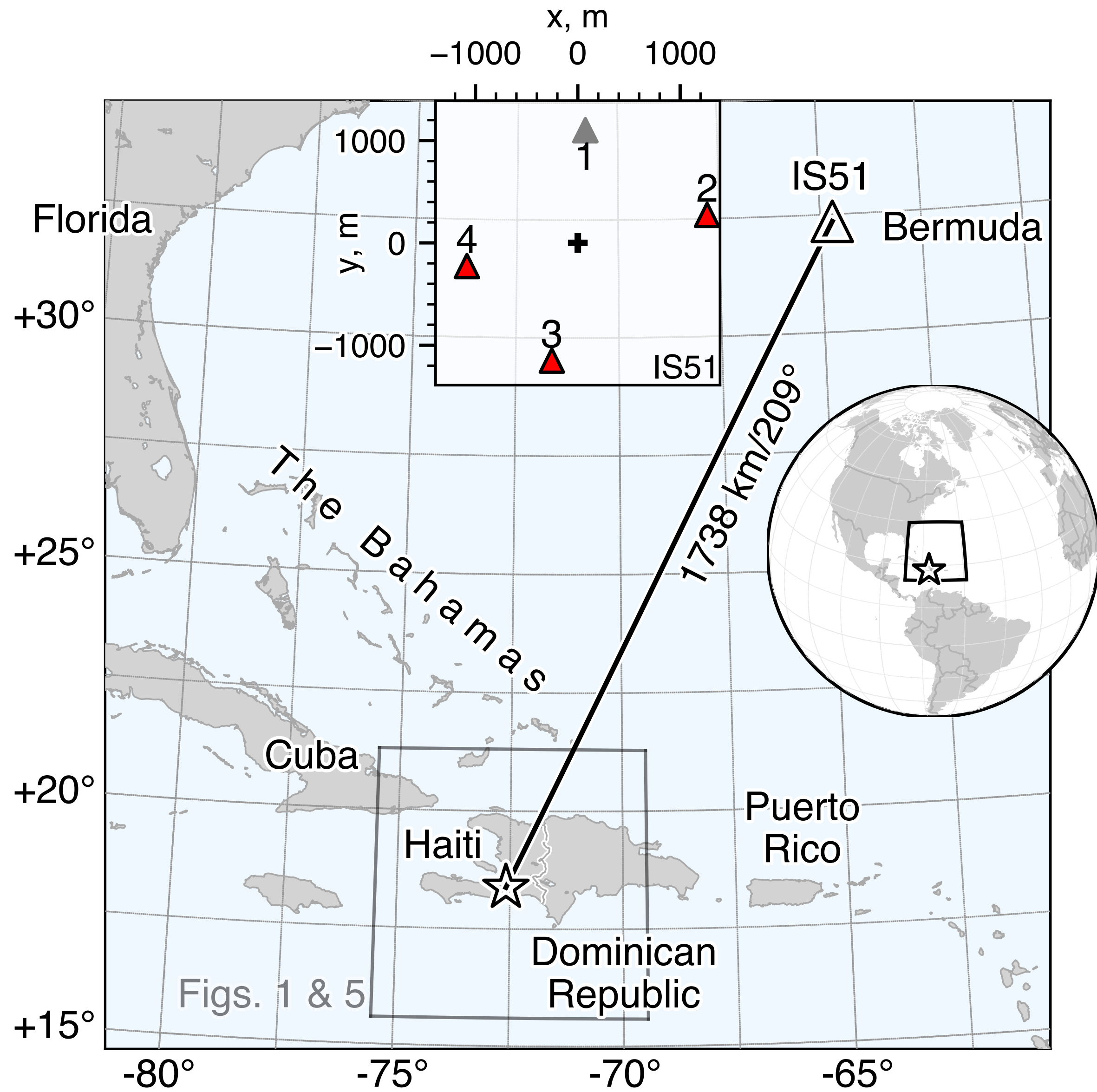
Scale based upon Worden et al. (2012)

ShakeMap generated more than 7 years after the earthquake

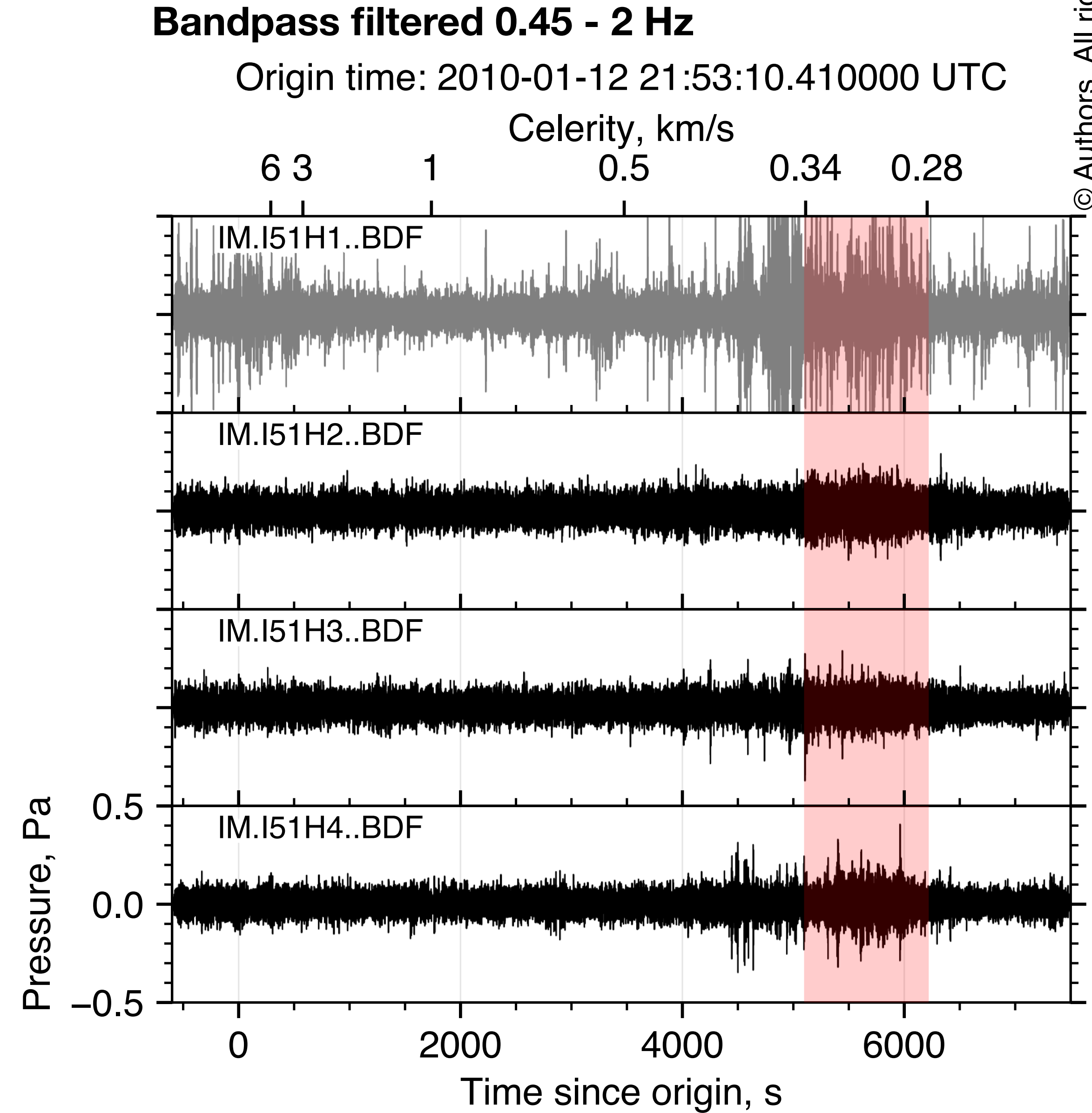
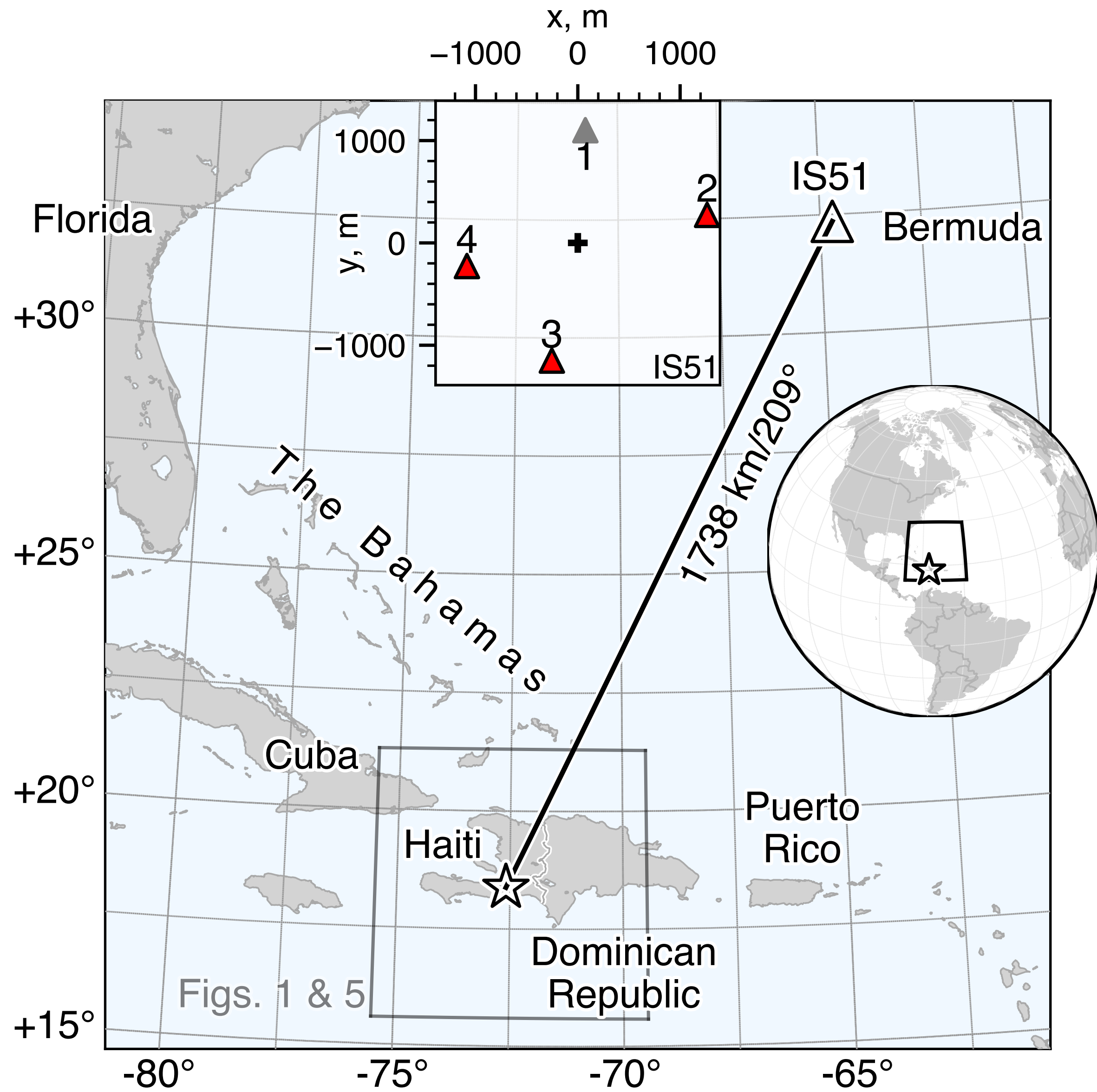
Data Processing



Data Processing

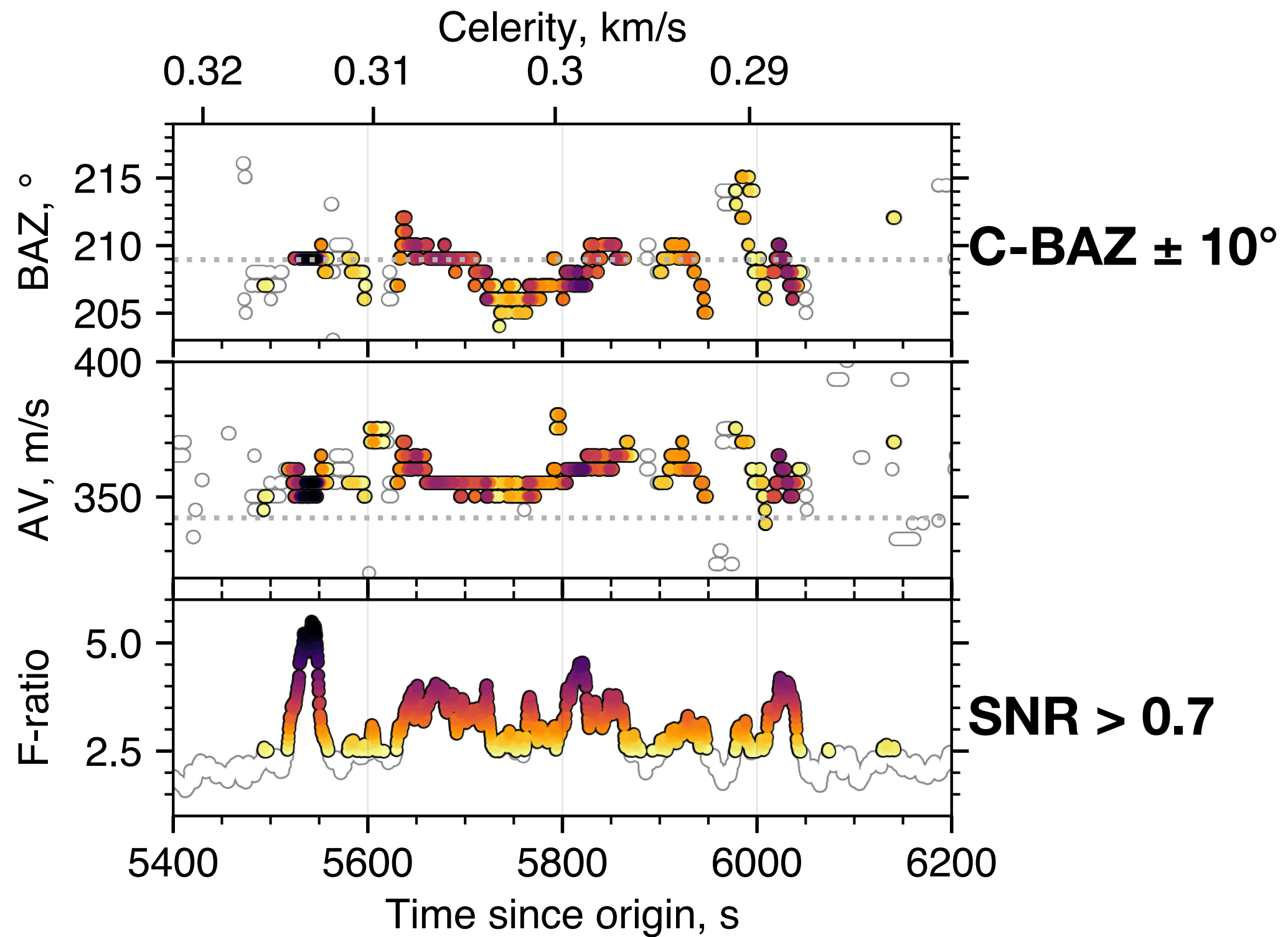
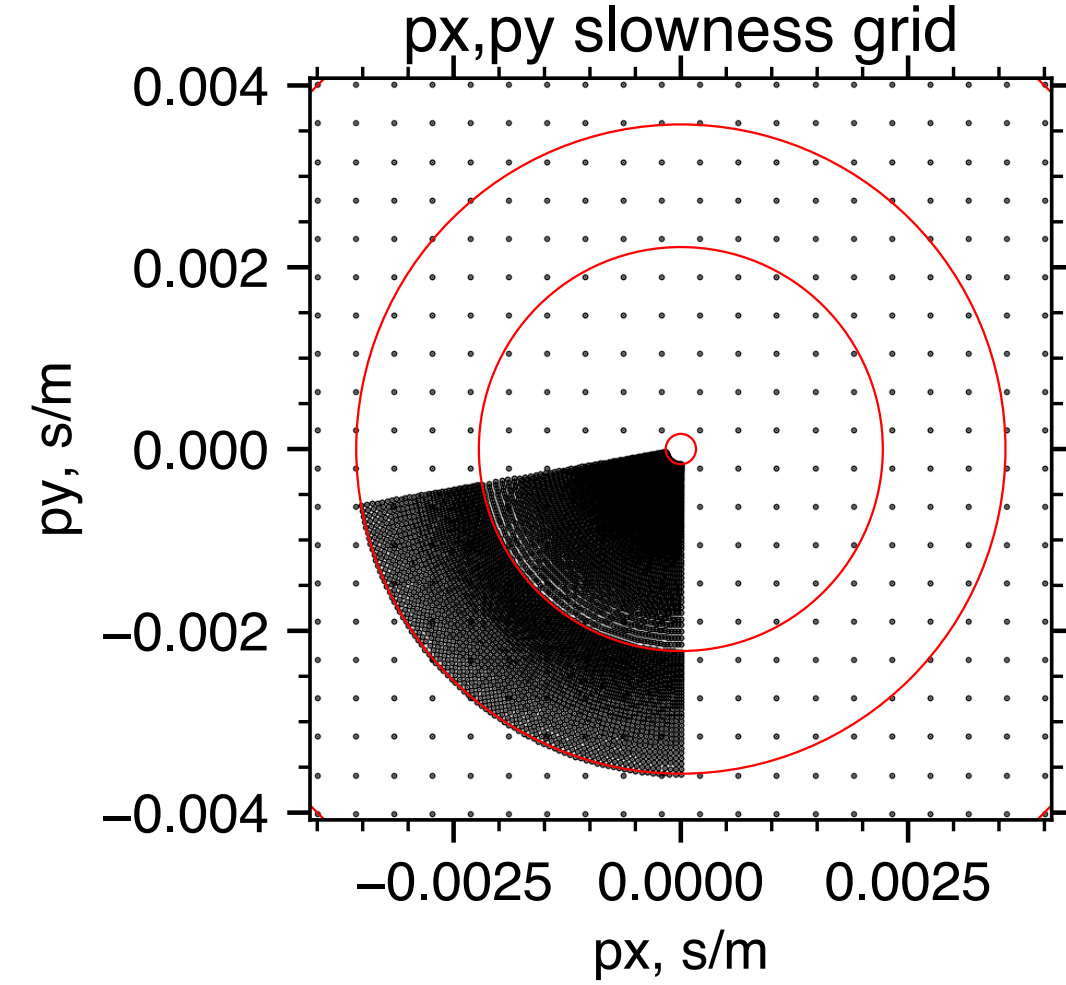
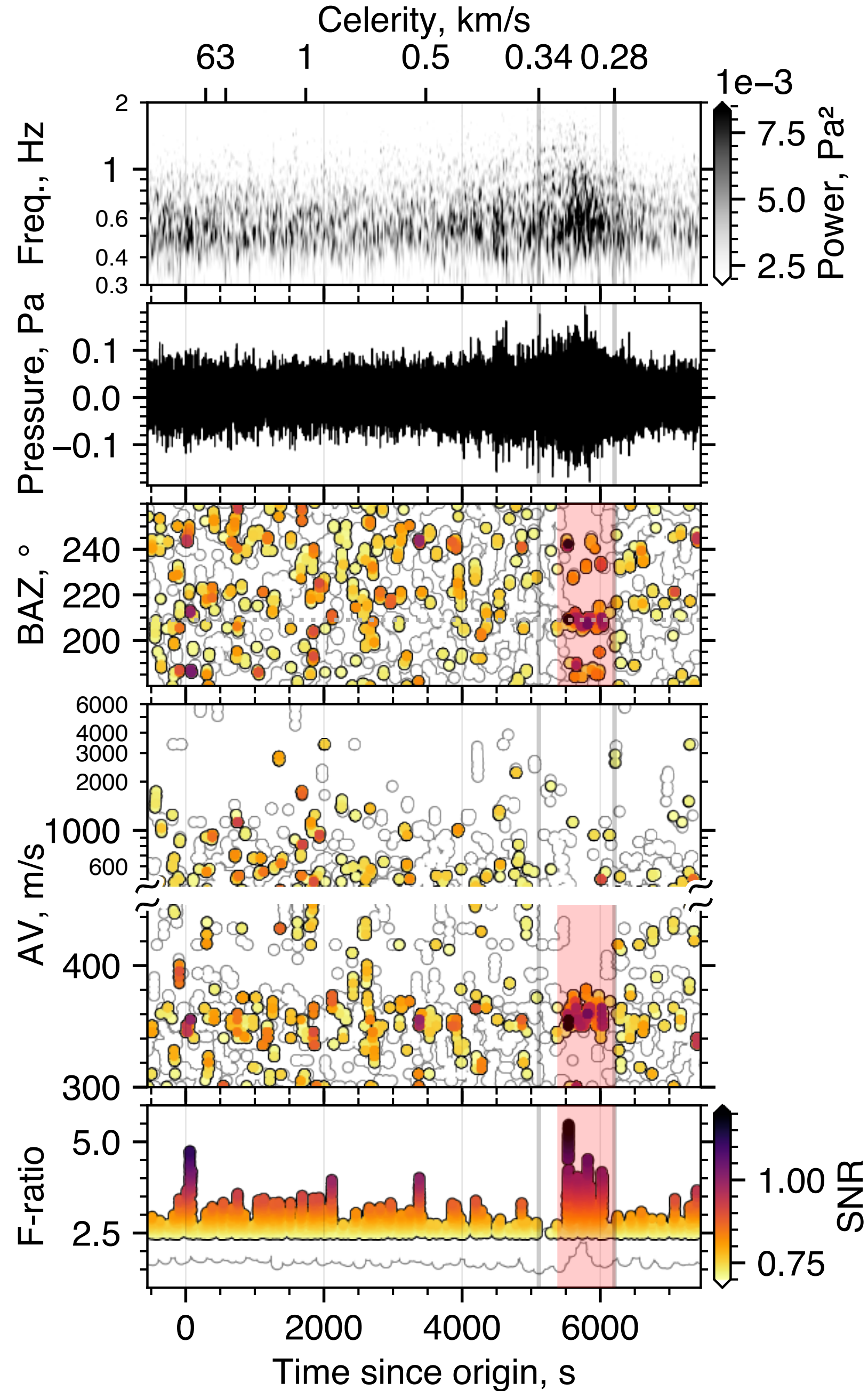


Data Processing

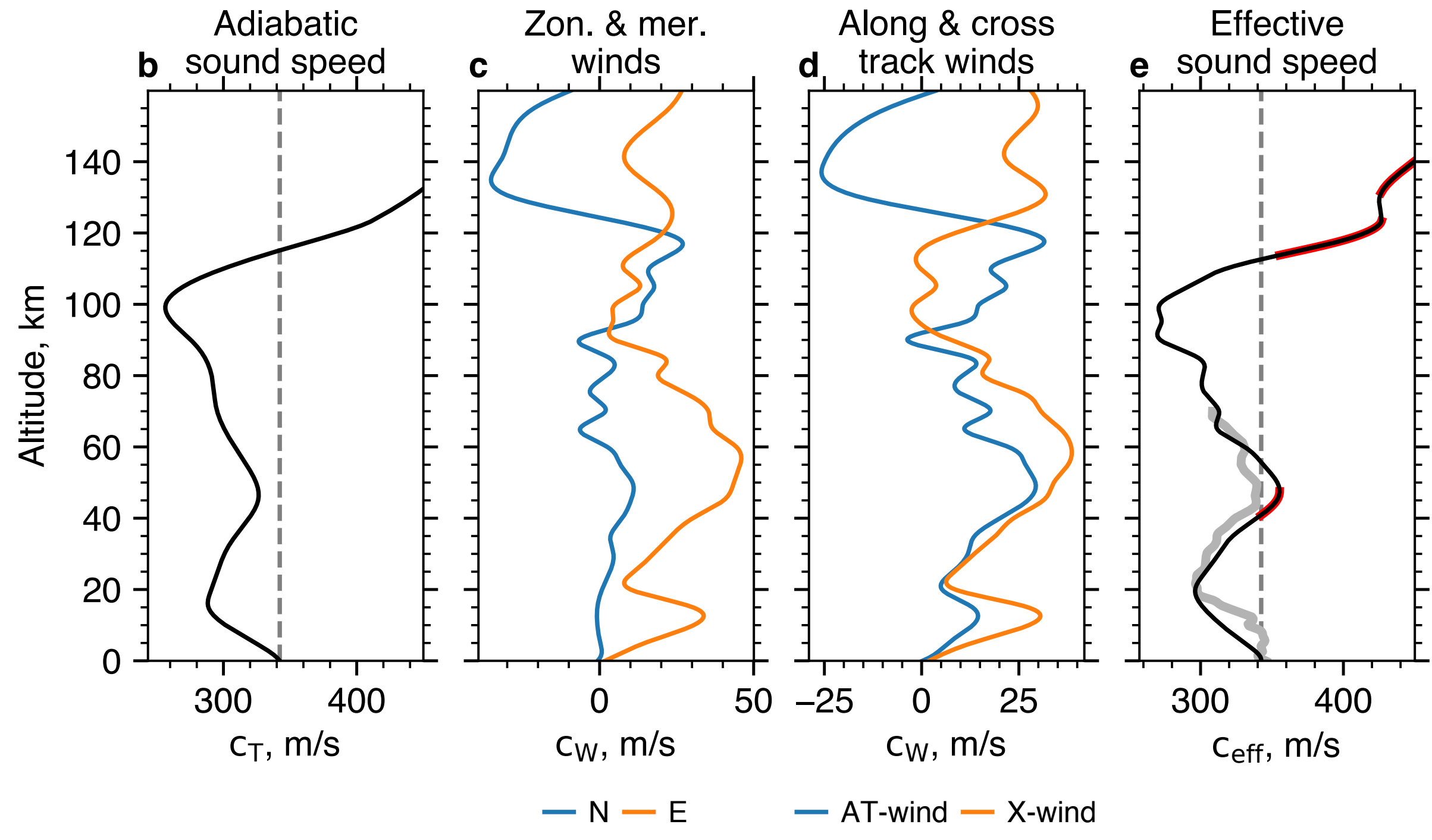
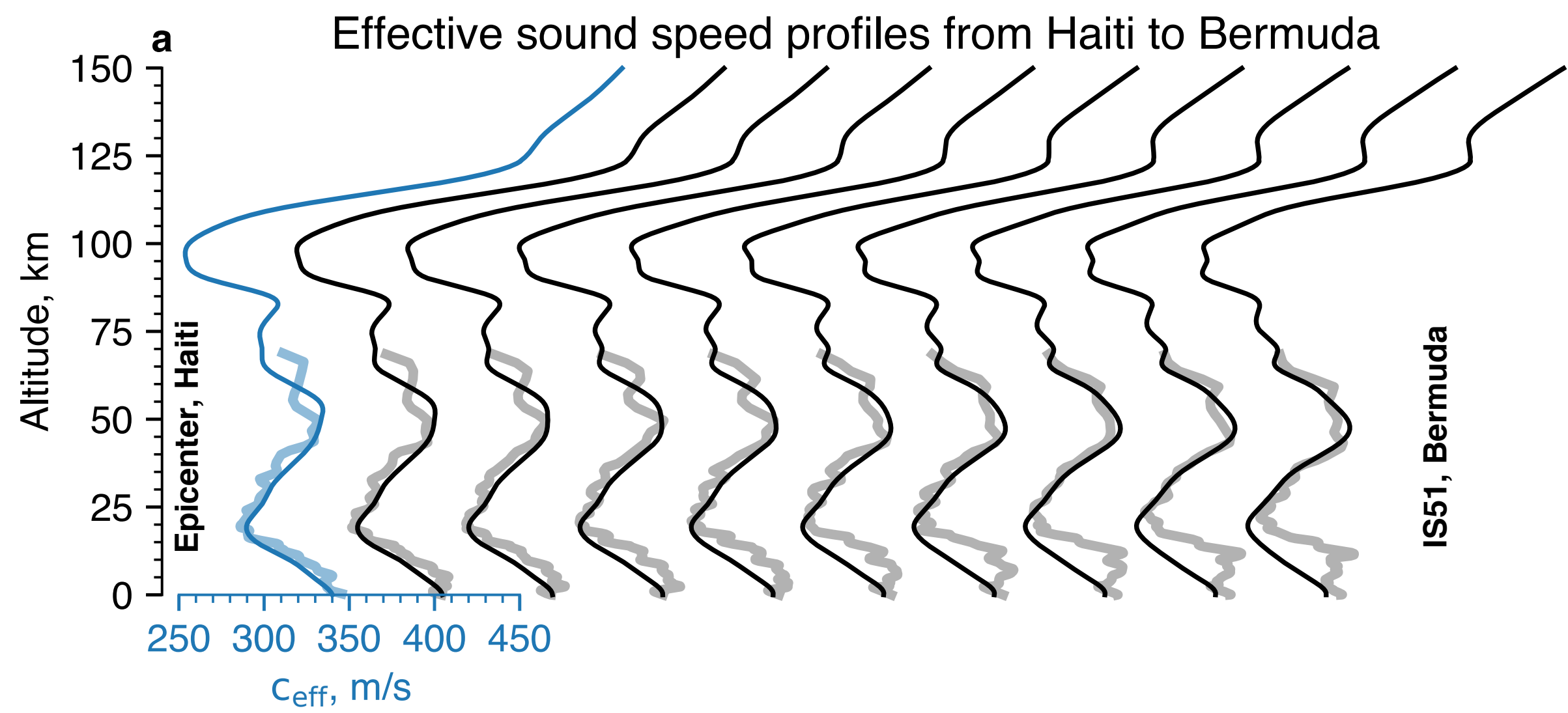
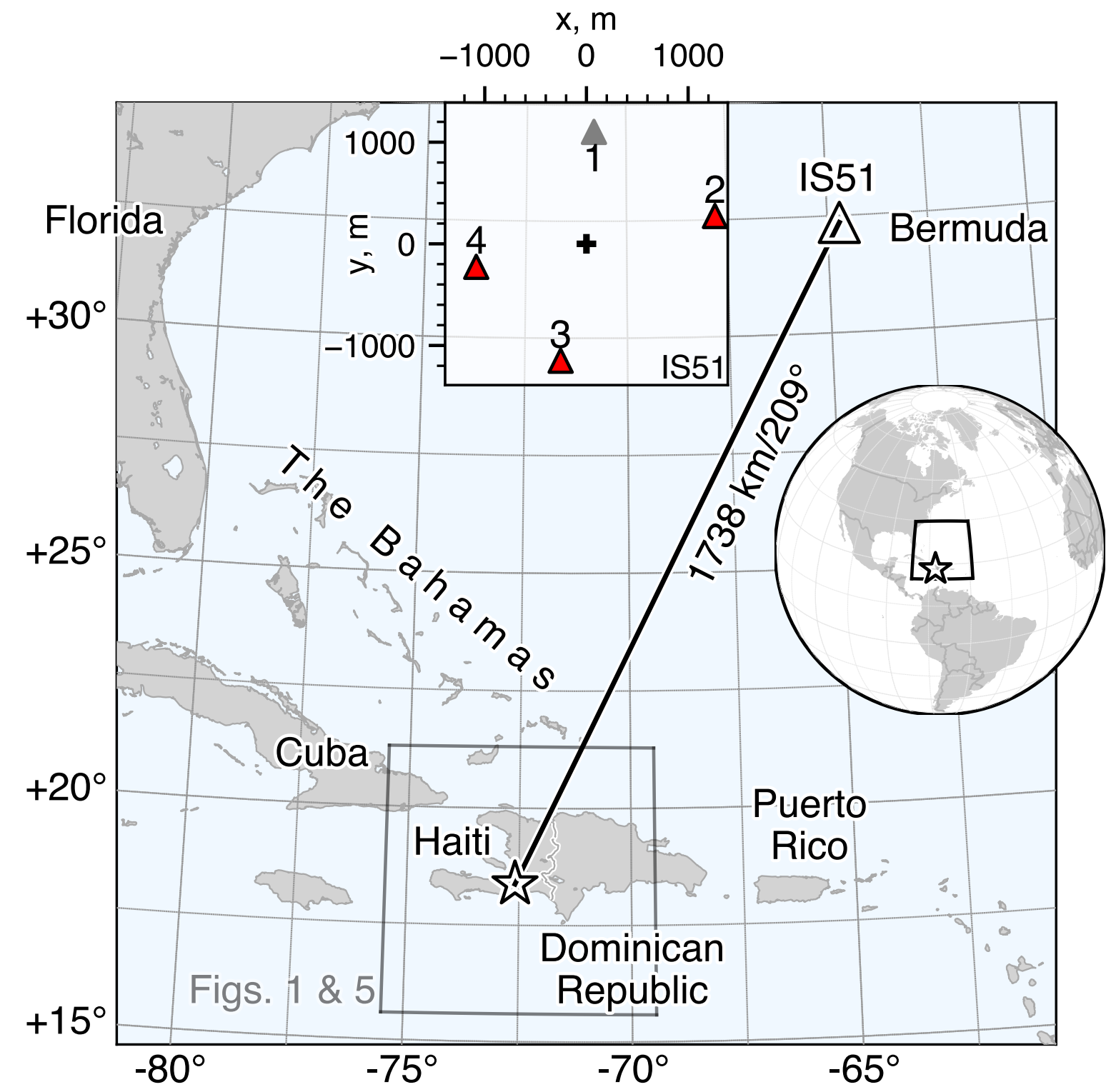


Array: IS51 | 3 elements | event range: 1738 km | event BAZ: 209°
 Beamforming: wlen: 30.00 s | overlap: 99% | freq: 0.45 --> 2.00 Hz
 Event origin time: 2010-01-12 21:53:10.410000 UTC

Array Processing

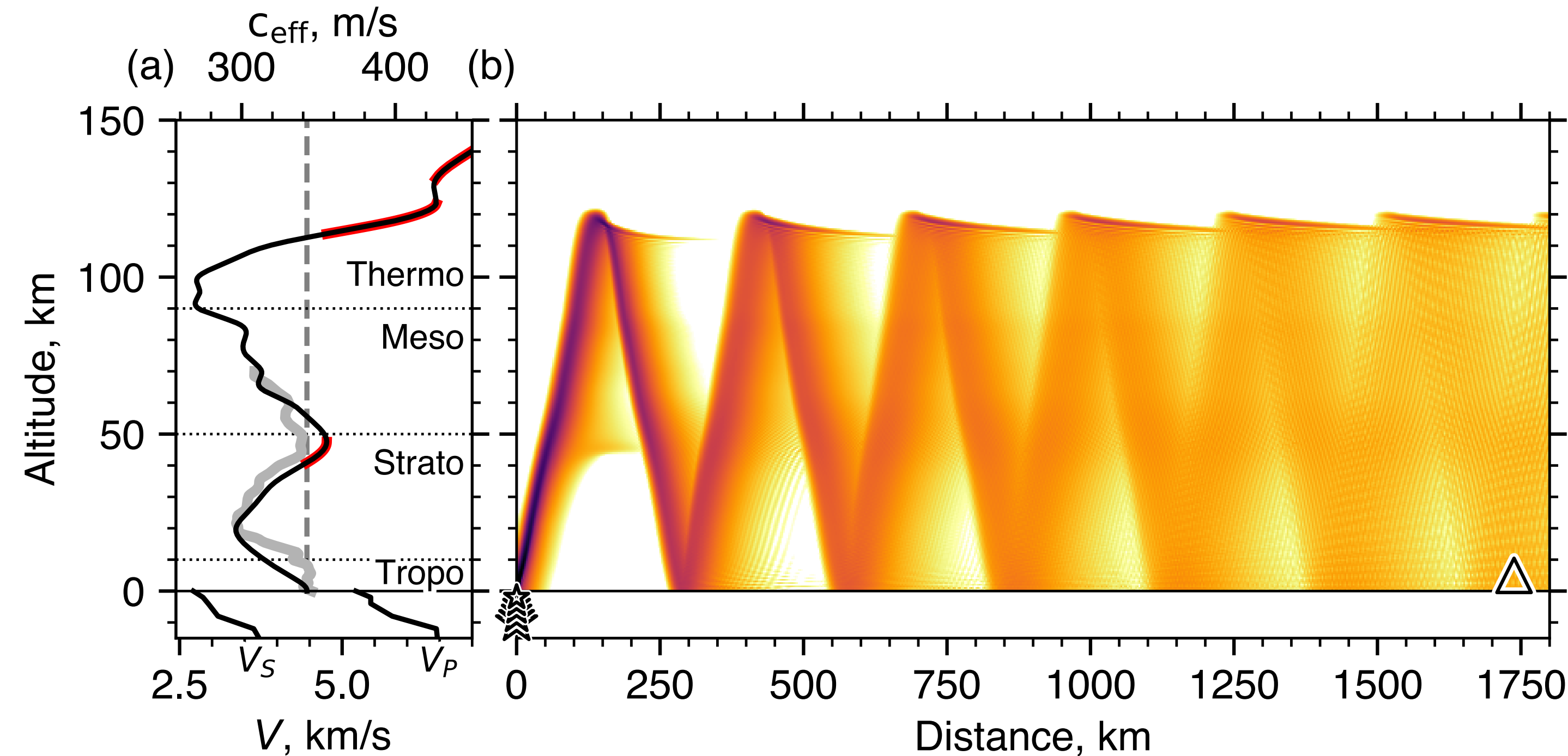


Seismo-Acoustic Coupling and Propagation Conditions



MSIS-00 & HWM14
ERA5 RA ECMWF

Seismo-Acoustic Coupling and Propagation Conditions



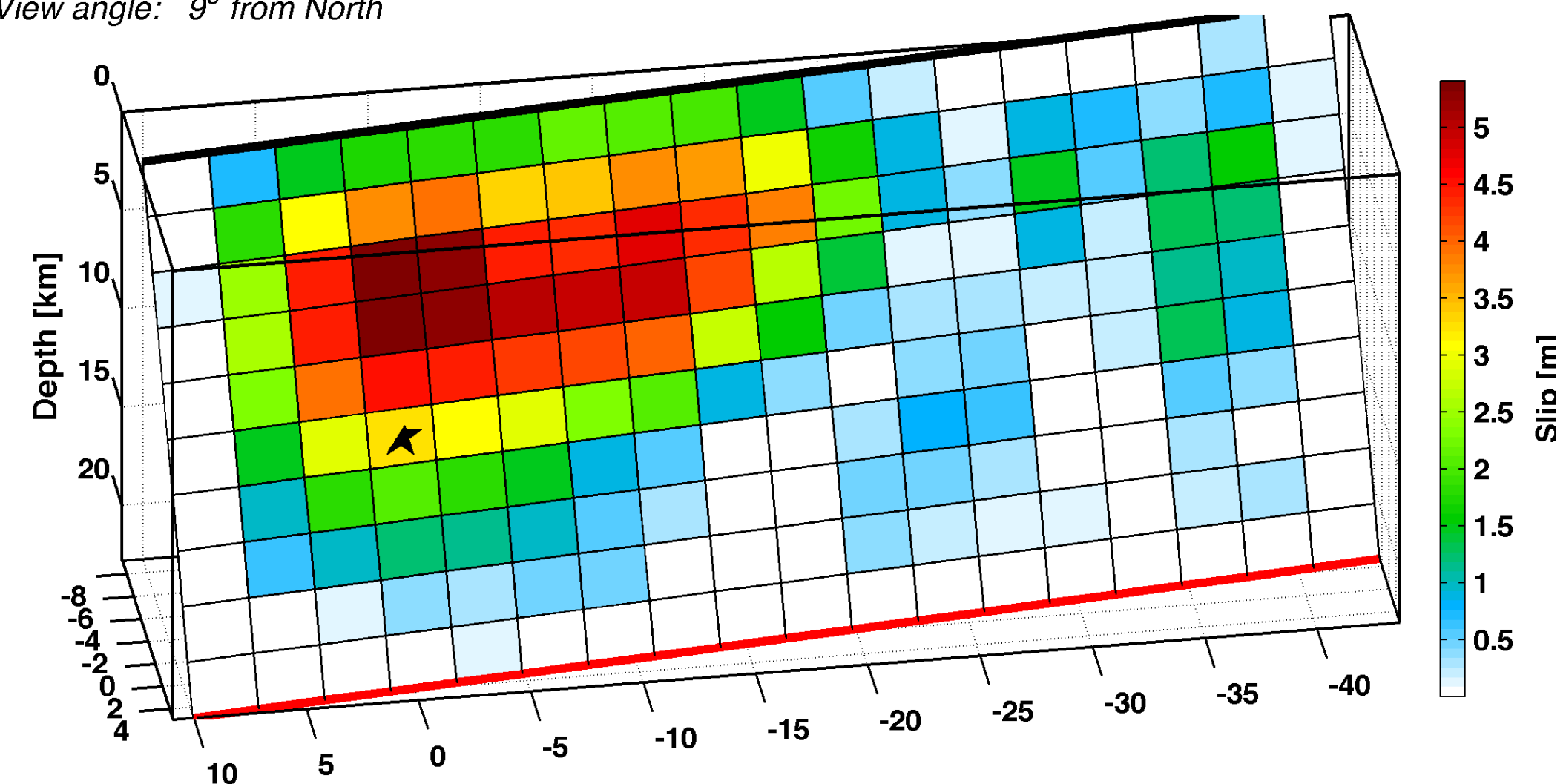
Fast Field Program (FFP) Setup:

- Sources at 10, 8, 6, 4, 2 km depth
- $f = 0.5$ Hz
- Modeled velocities: 300 - 450 m/s

Averbuch et al. ASA2020

s2010HAITIX02HAYE
 Mw 7.0 Mo 3.55e+19
 Lat/Lon/Dep: 18.51°, -72.53°, 12.2 km
 View angle: 9° from North

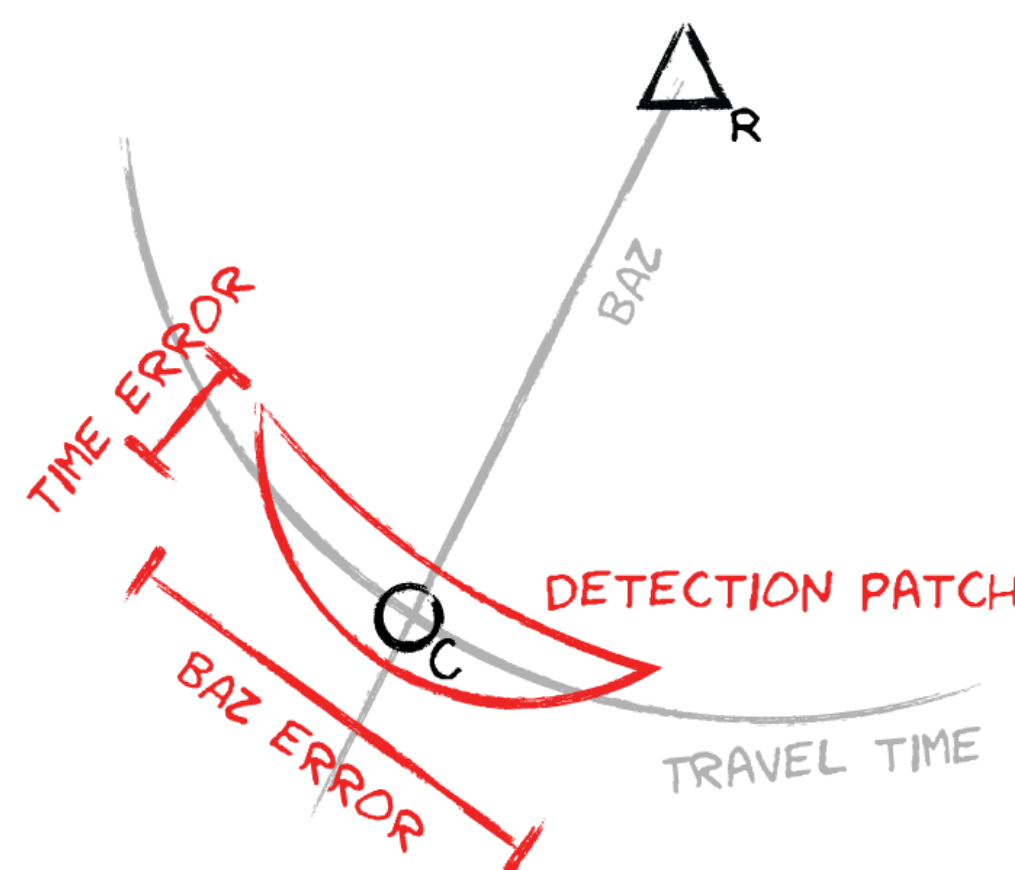
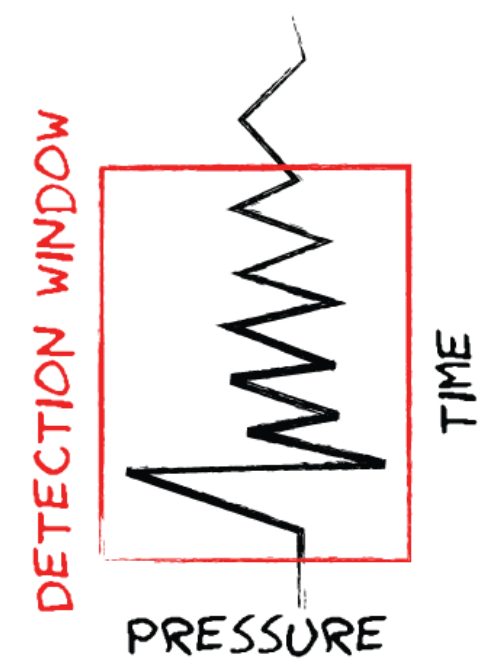
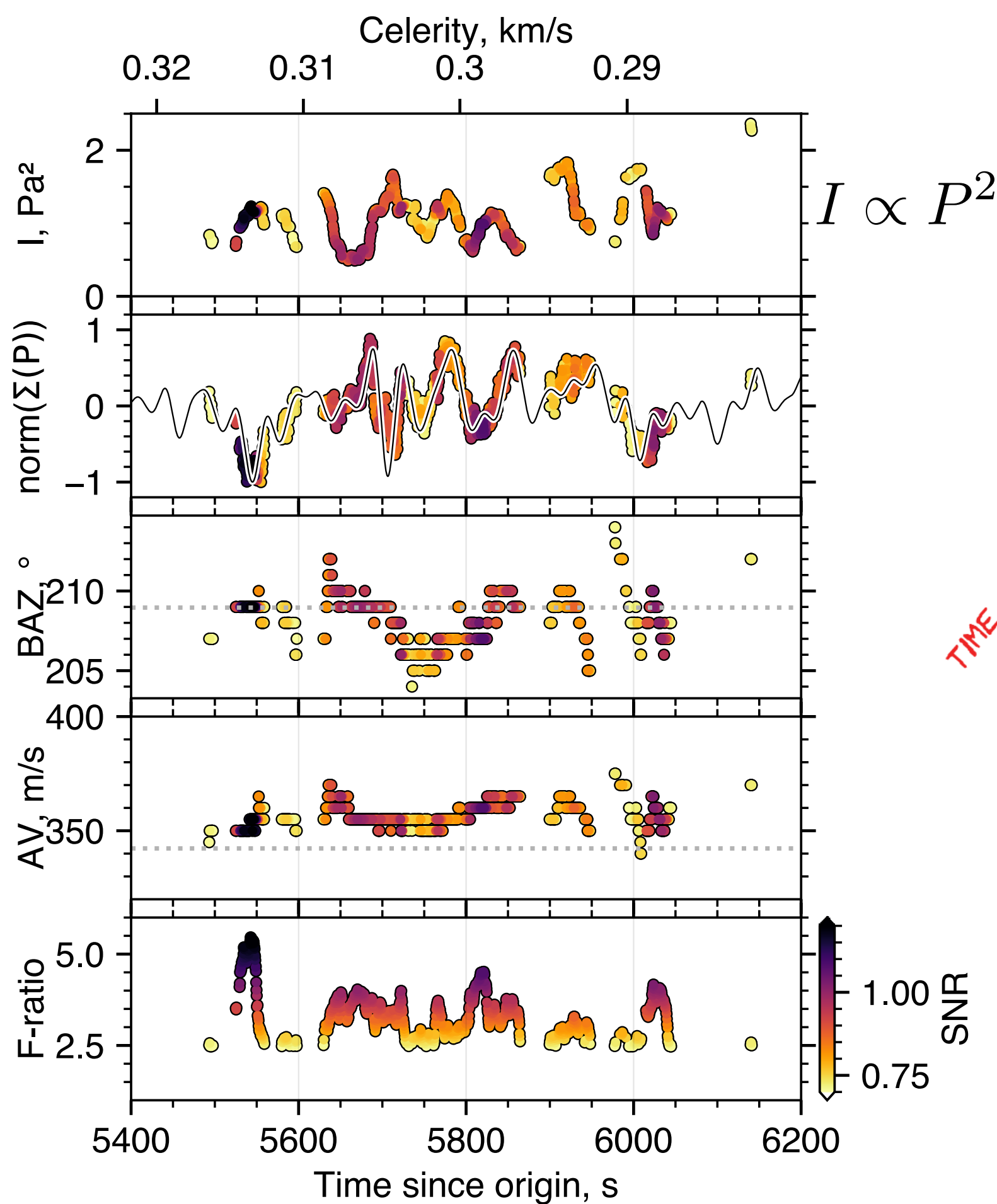
Haiti



Hayes et al. NG2010

Backprojections

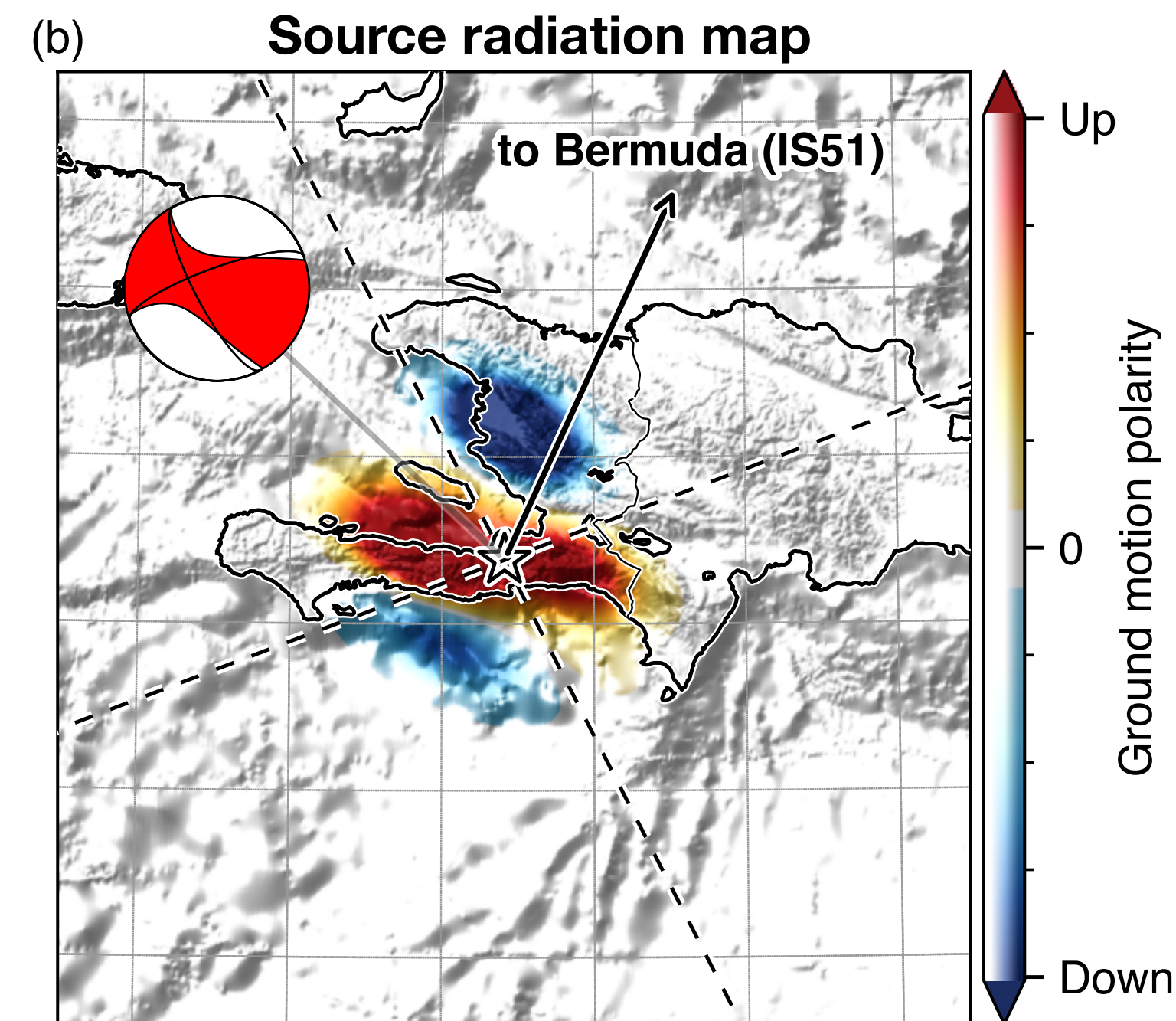
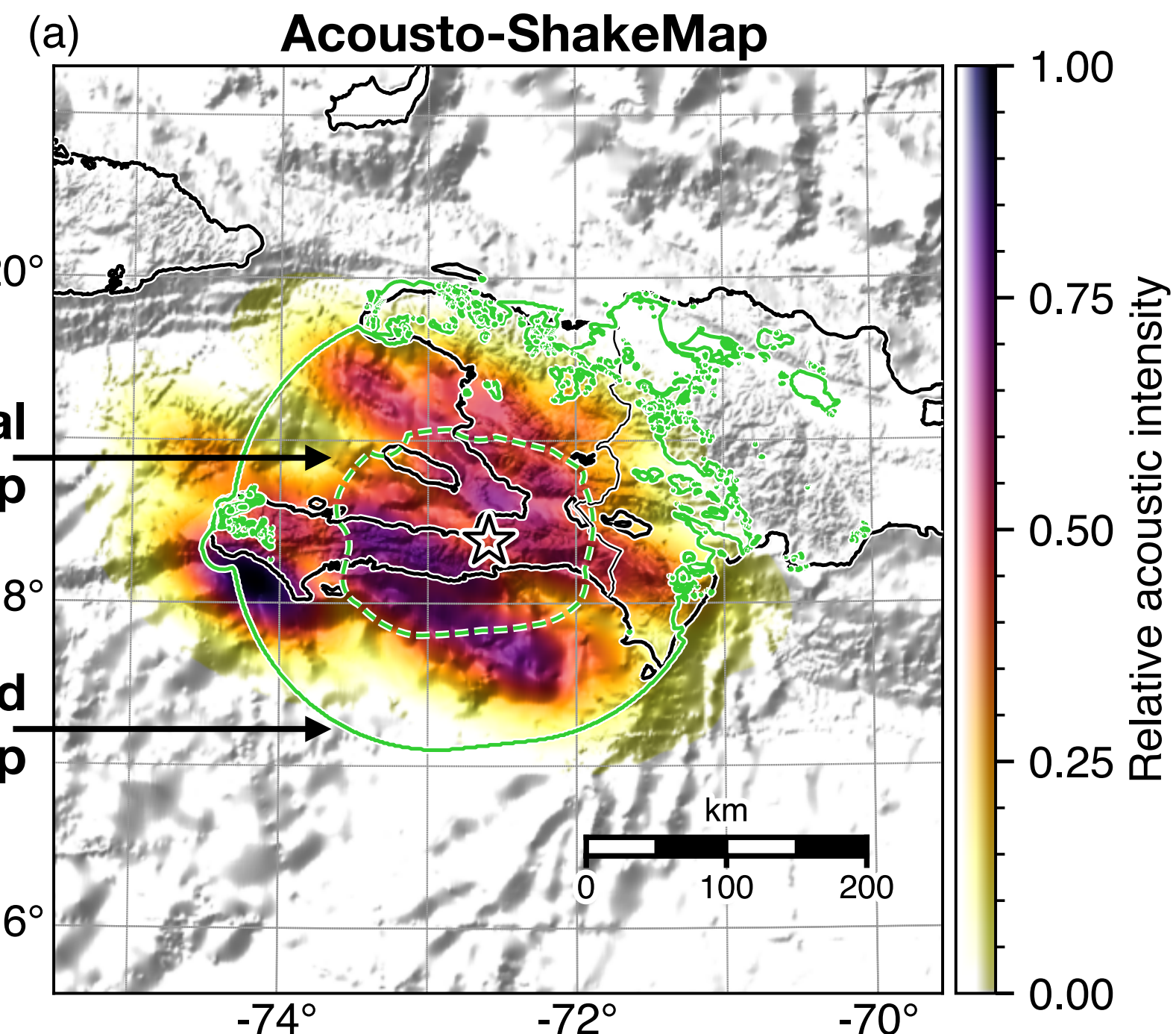
Epicentral infrasound detections



Backprojection accounts for:

- Horizontal advection effects from along- and cross-track wind
- Uncertainties with regard to the time of detection and back azimuth

Important: Back azimuth uncertainty grows with backprojection distance

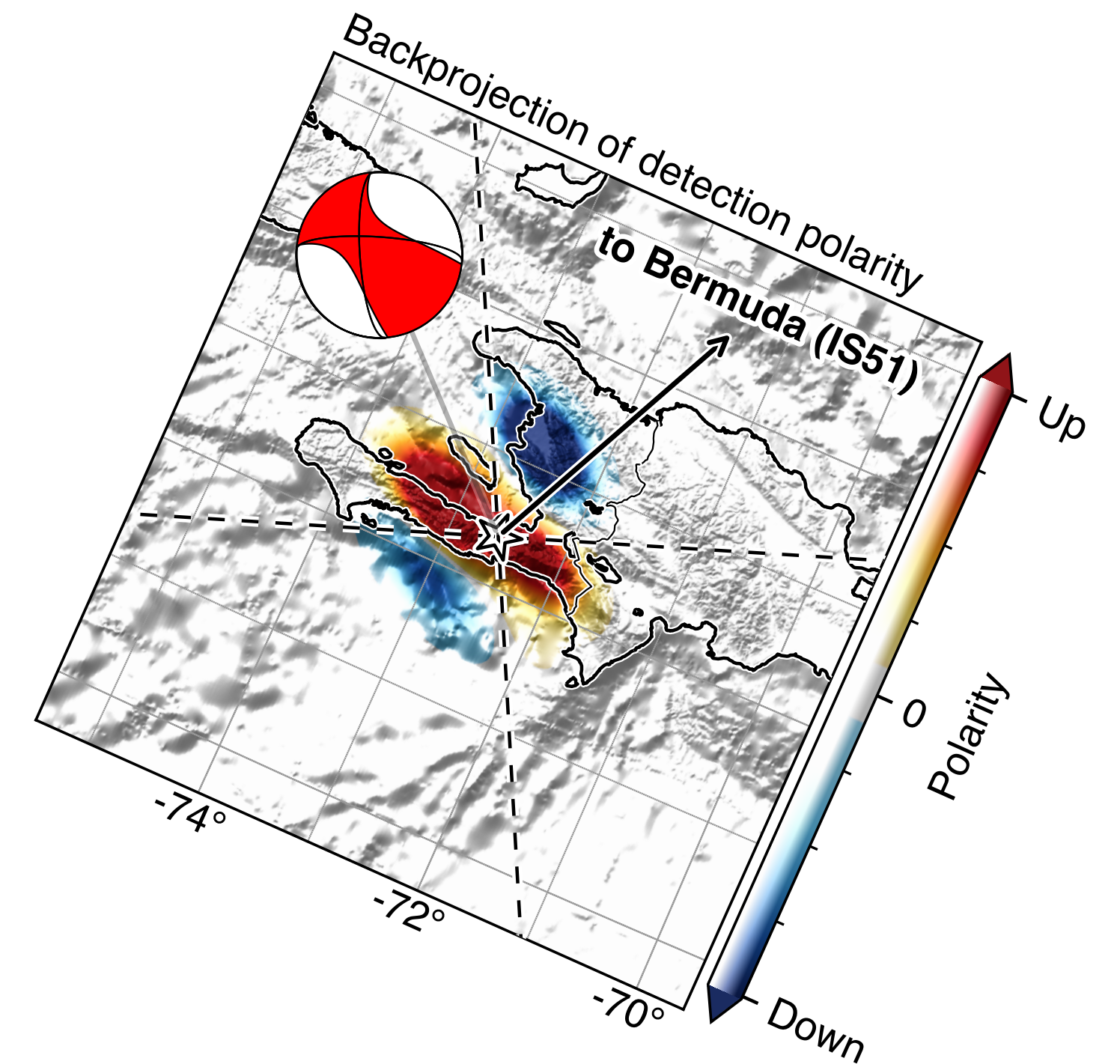
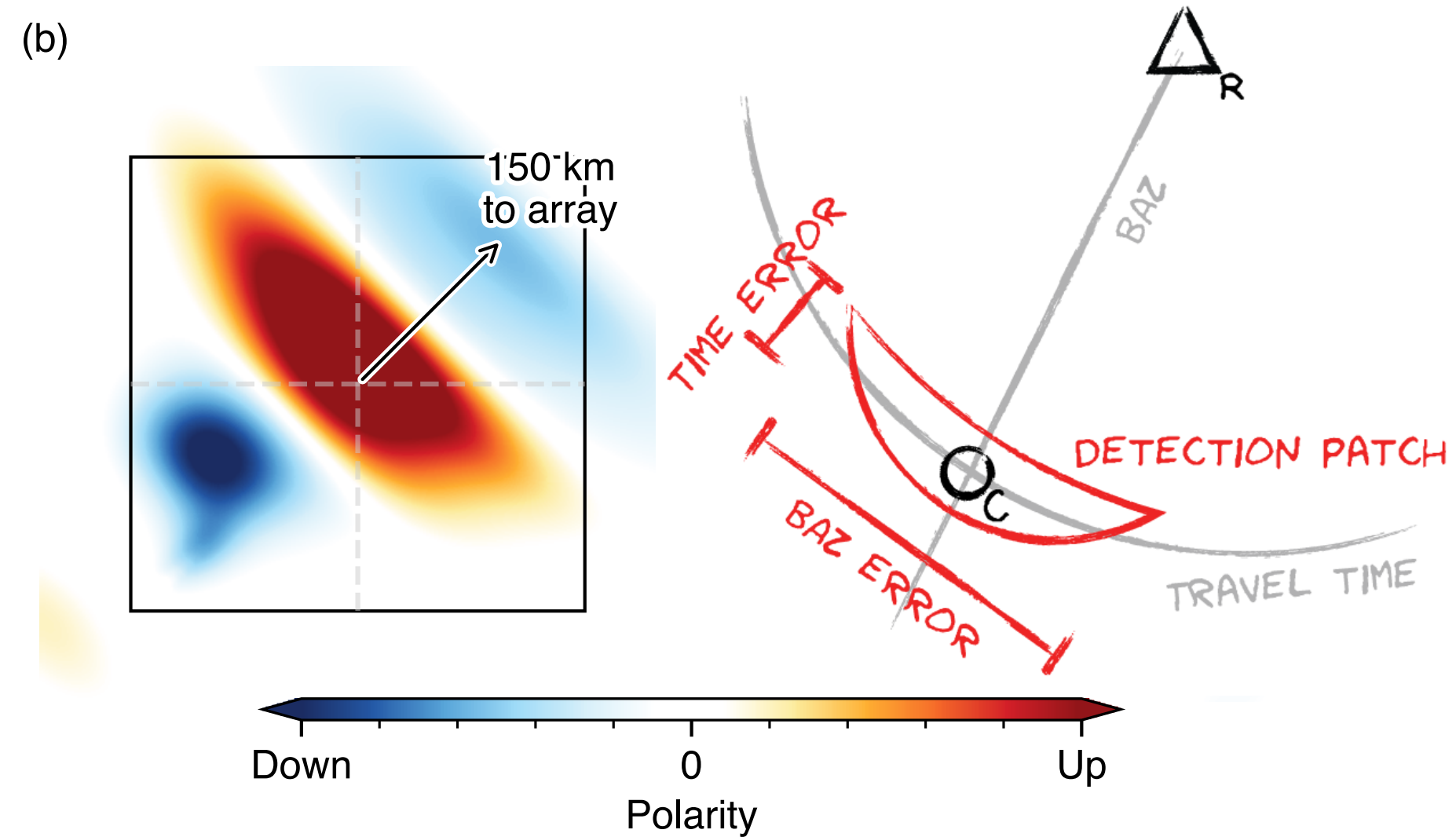
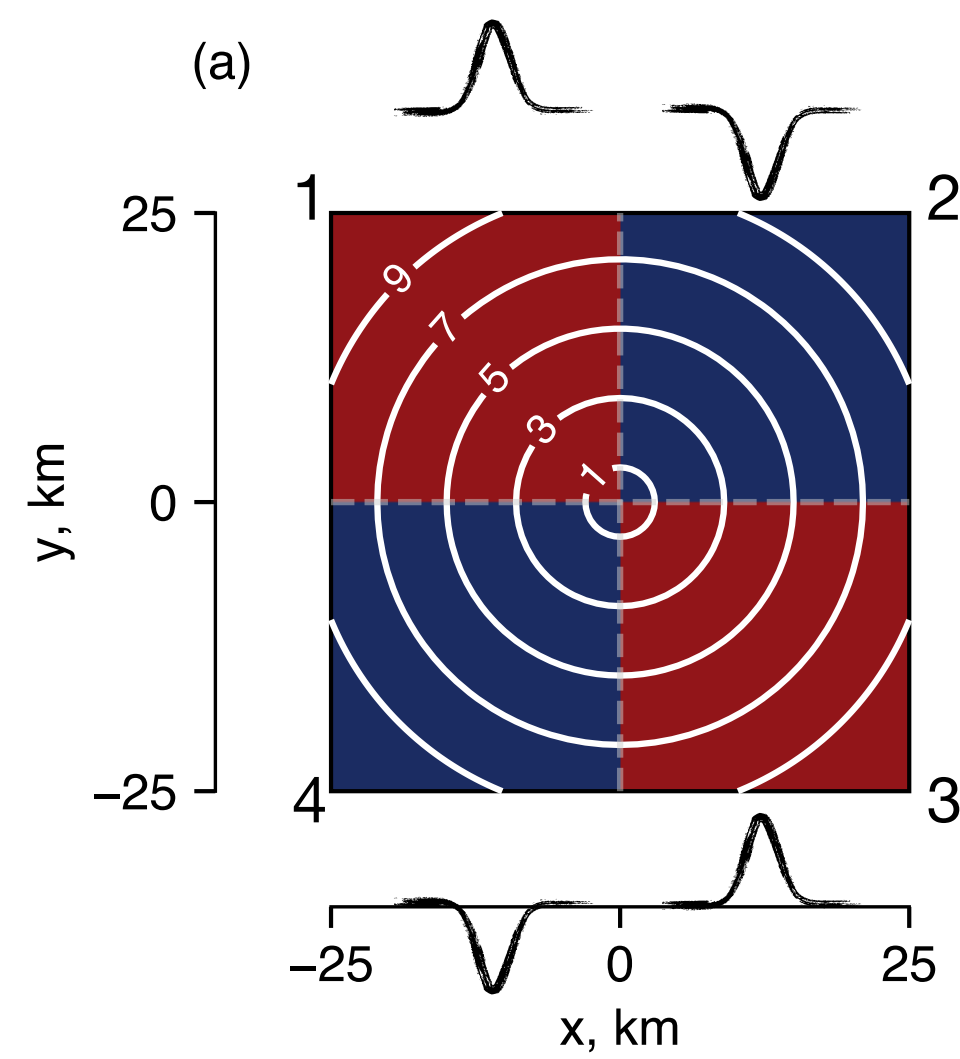
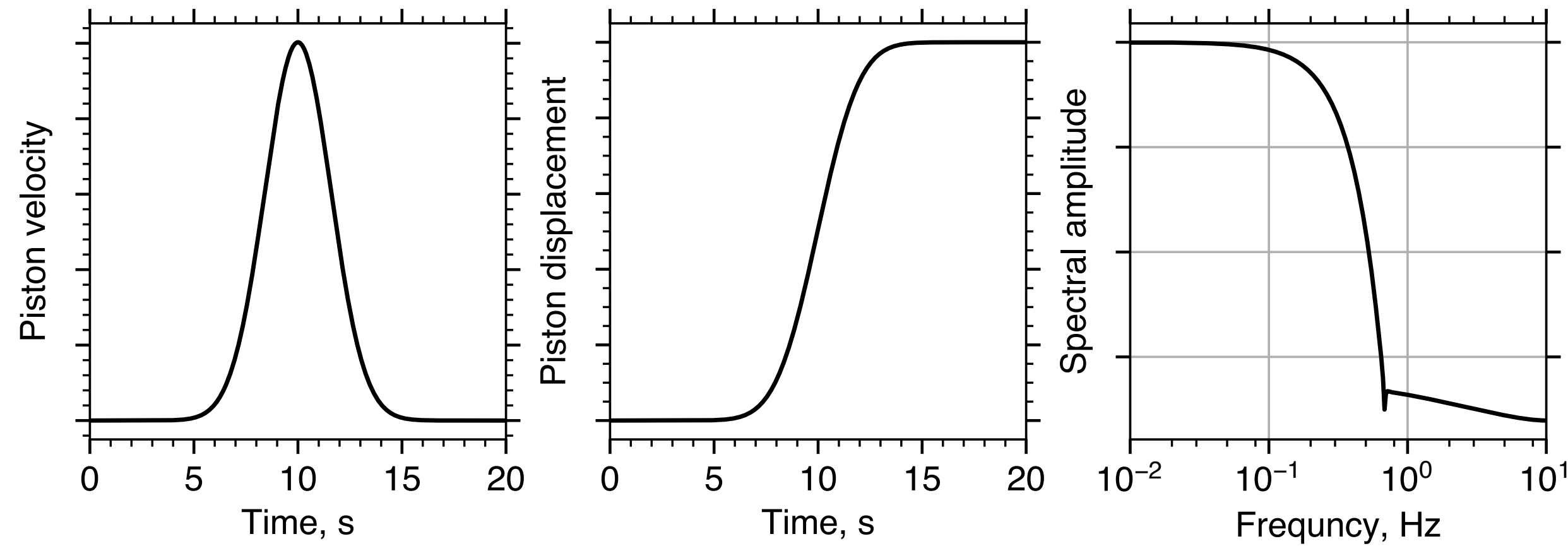


Proof of Concept with the help of Lord Rayleigh



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**50 km extended source
discretised to 400 m pistons
Gaussian STF, $f_c=0.1$ Hz
moveout velocity of 3 km/s**

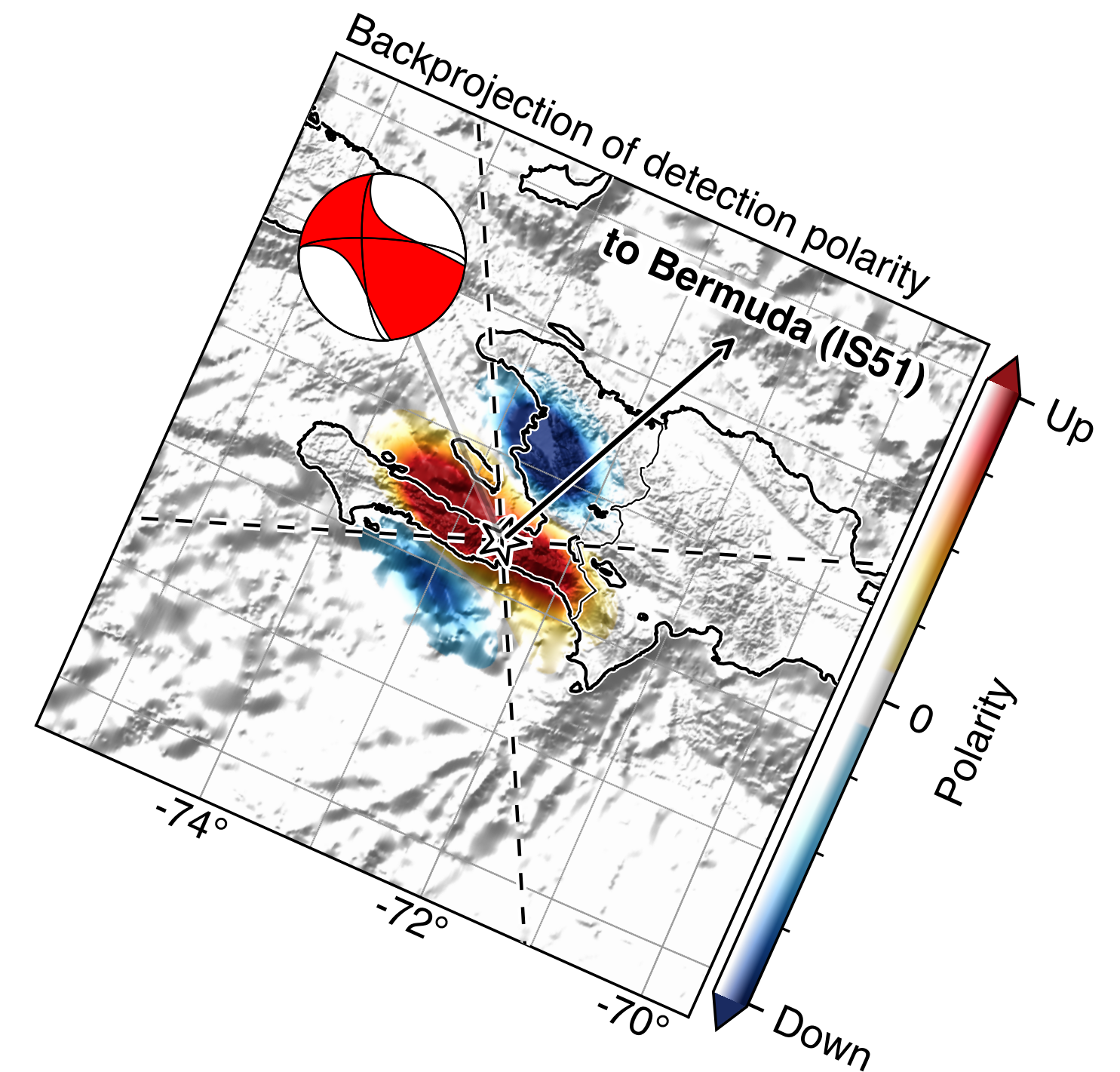
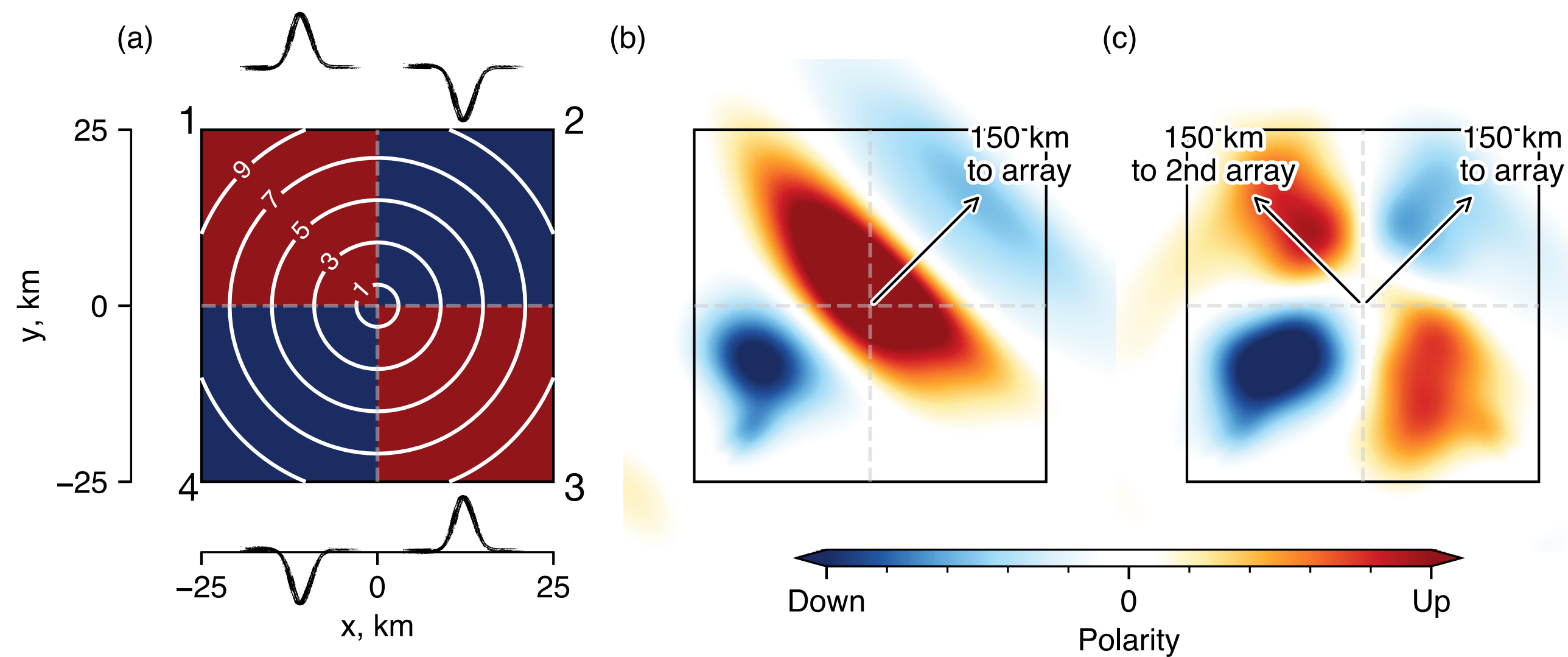
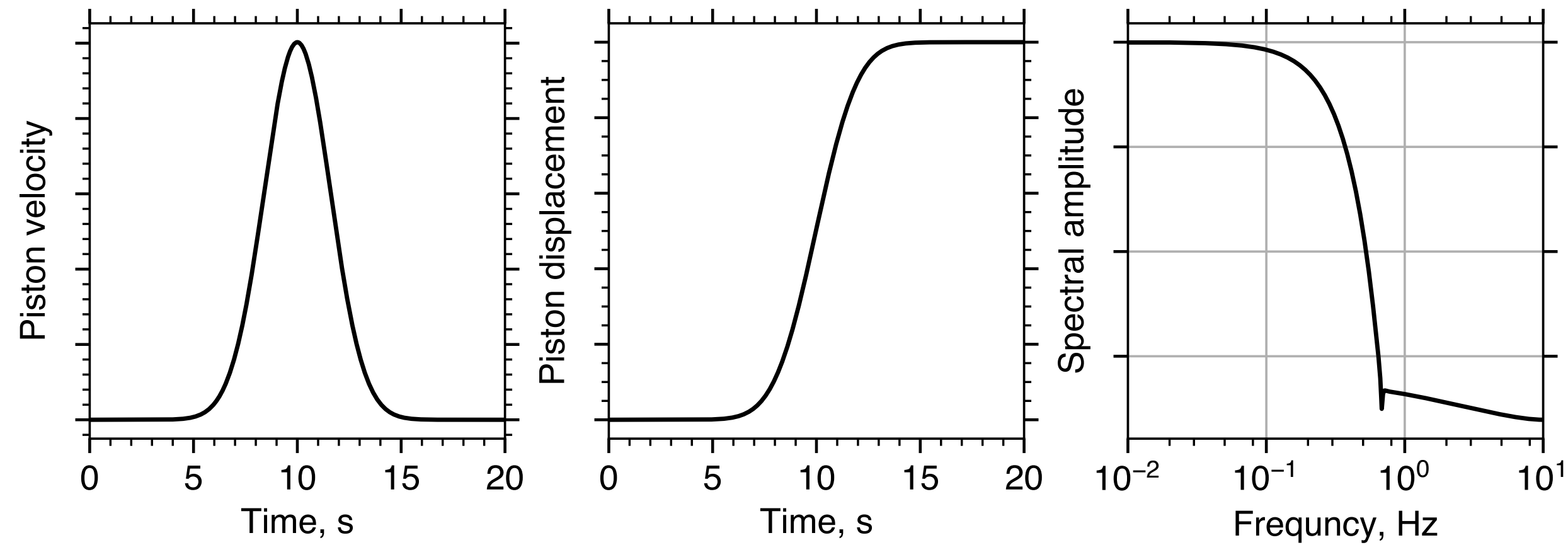


Proof of Concept with the help of Lord Rayleigh



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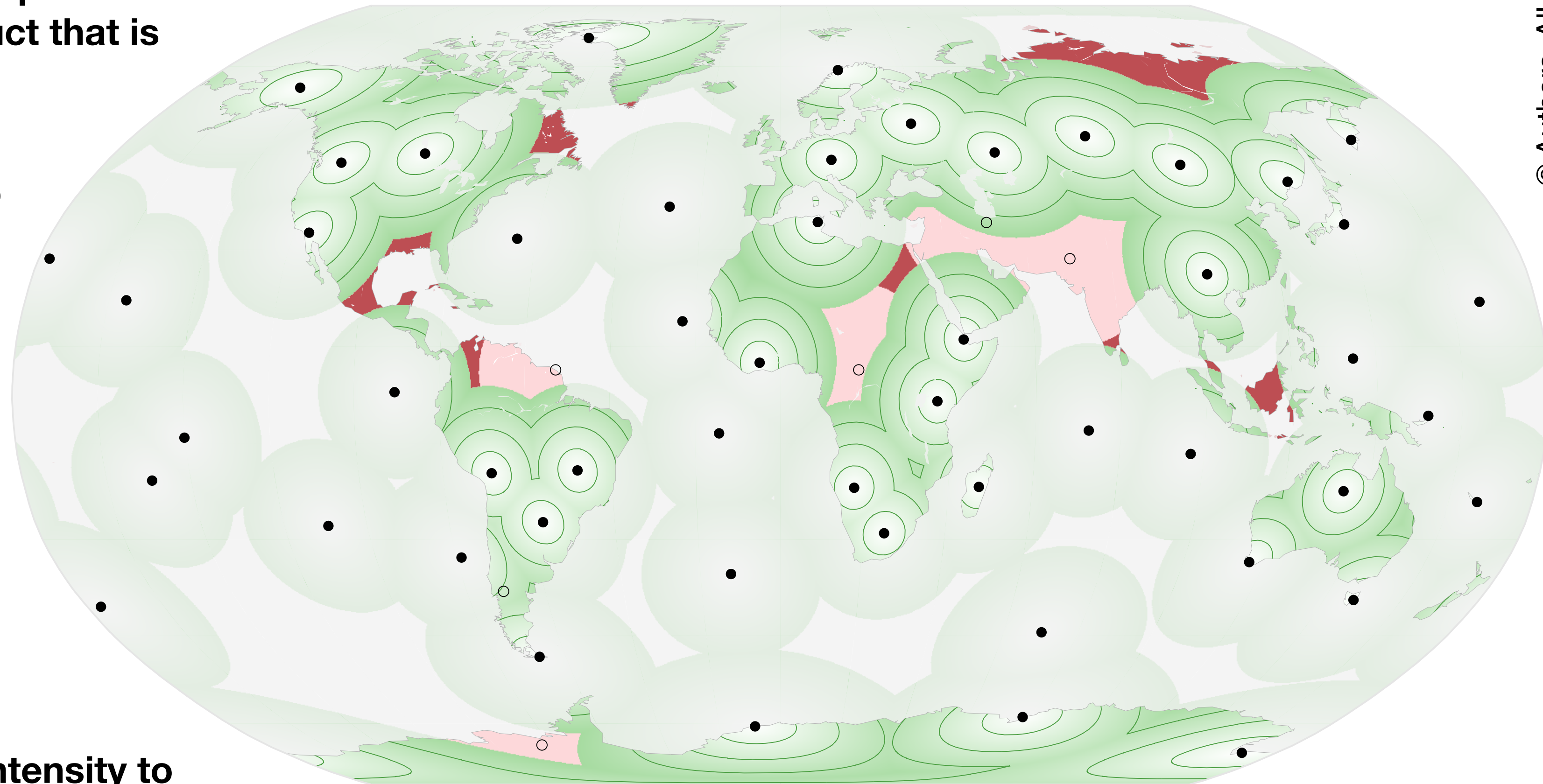
Discussion & Conclusions

- The temperature gradient in the thermosphere makes it so there is always a thermospheric duct that is effective to about 2000 km.
- This means that the existing infrasound stations of the IMS already provide 90% coverage of landmass.
- When the network is complete the coverage will increase to 97%.
- The remaining gaps can be easily covered with nationally operated stations.

Next steps:

3D and 1D raytracing...

Translating measured acoustic pressure intensity to ground motions at the epicenter...



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